



ELSEVIER

Available online at www.sciencedirect.com

SCIENCE @ DIRECT®

Physics and Chemistry of the Earth xxx (2005) xxx–xxx

**PHYSICS
and CHEMISTRY
of the EARTH**
www.elsevier.com/locate/pce

Estimation of small reservoir storage capacities in a semi-arid environment

A case study in the Upper East Region of Ghana

J. Liebe^{a,d,*}, N. van de Giesen^{b,d}, M. Andreini^{c,d}

^a Department of Biological and Environmental Engineering, Cornell University, 76 Riley-Robb Hall, Ithaca, NY 14853-5701, USA

^b Civil Engineering and Geosciences TU Delft, Stevinweg 1, 2628 CN Delft, The Netherlands

^c International Water Management Institute (IWMI) Ghana, PMB CT 112 Cantonments, Accra, Ghana

^d Center for Development Research (ZEF), University of Bonn, Walter-Flex-Str. 3, 53113 Bonn, Germany

11 Abstract

12 In semi-arid regions at the margins of the Sahel, large numbers of small reservoirs capture surface runoff during the rainy season,
13 making water available during the dry season. For the local population, small reservoirs are important water sources which help
14 them cope with droughts. The lack of knowledge of the number of existing reservoirs, their distribution, and their storage volumes
15 hinders efficient water management and reservoir planning. The authors have developed a simple method that allows the estimation
16 of reservoir storage volumes as a function of their surface areas. This function is based on an extensive bathymetrical survey that
17 was conducted in the Upper East Region of Ghana. In combination with satellite imagery, this function can be used determine and
18 monitor the storage volumes of large numbers of small reservoirs on a regional scale.

19 © 2005 Published by Elsevier Ltd.

20 *Keywords:* Small reservoirs; Drought mitigation; Water management; Dams; Ghana; Africa

22 1. Introduction

23 In the semi-arid regions of Northern Ghana, large
24 numbers of small reservoirs dot the landscape. Reser-
25 voirs capture surface runoff during the rainy season
26 making water available in the dry season. For the rural
27 population in environments such as the Upper East Re-
28 gion of Ghana, the presence of a small reservoir is an
29 important means of overcoming minor droughts. Effi-
30 cient water management and sound reservoir planning
31 are hindered by the lack of information about the func-
32 tioning of these reservoirs. The reservoirs were built at

different times by various agencies. Poor record keeping 33
and the lack of appropriate institutional support result 34
in deficiencies of information on the capacity, operation, 35
and maintenance of these structures. As a first step to- 36
wards understanding the impact these dams have on 37
the availability of water in this area, the authors devel- 38
oped a simple method for estimating and monitoring 39
the storage volumes of these reservoirs on the basis of 40
their surface areas. The use of satellite imagery allows 41
us to measure the reservoir surface areas and gives in- 42
sight into the statistical (e.g. size, and frequency) and 43
spatial distribution. The area based volume estimation 44
is made possible because this region is morphologically 45
and morphometrically regular. The reservoirs are 46
located in the stream channels, and the morphometry 47
of stream channels are a response to surface runoff char- 48
acteristic of this area (Windmeijer and Andriess, 1993). 49

* Corresponding author. Address: Department of Biological and Environmental Engineering, Cornell University, 76 Riley-Robb Hall, Ithaca, NY 14853-5701, USA. Tel.: +1 607 255 2463; fax: +1 607 255 4080.

E-mail address: jrl58@cornell.edu (J. Liebe).

50 Damming these streams results in characteristic rela-
51 tionships between volumes to surface areas.

52 2. The study area

53 The Upper East Region of Ghana is situated in the
54 center of the Volta Basin (Fig. 1, van de Giesen et al.,
55 2002). The Upper East is inhabited by approximately
56 one million people and has a population density of
57 96.5 inhabitants/km² (Asenso-Okyere et al., 2000). With
58 a poverty incidence of 88% in 1998/1999, the Upper East
59 has the largest portion of poor people of Ghana's ten re-
60 gions (Ghana Statistical Service, 2000). The residents in-
61 comes are generated from rainfed and some irrigated
62 agriculture. Population growth places pressure on scarce
63 land and water resources. The scarcity of usable water
64 resources is mainly due to the climate, especially the
65 mode of rainfall. The Upper East's semi-arid climate is
66 characterized by a three month, monomodal rainy sea-

67 son. Ninety percent of the Region's total rainfall
68 (986 mm) occurs as thunderstorms, originating from
69 squall lines (Eldridge, 1957; Hayward and Oguntoyinbo,
70 1987; Friesen, 2002). Rainfall intensities often exceed the
71 soil's infiltration rates causing surface runoff, without
72 replenishing soil moisture and groundwater. Small reser-
73 vvoirs help make better use of the rainfall by capturing
74 runoff. This water can be used for domestic purposes
75 and agricultural production. The small reservoirs' prox-
76 imity to places of demand is another advantage that
77 makes them an appropriate tool for drought mitigation.

78 3. Reservoir inventory with satellite imagery

79 Due to the lack of baseline data, our inventory of res-
80 ervoires was conducted by means of remote sensing. The
81 reservoirs were classified with four Landsat ETM
82 images. Three images were acquired at the end of the
83 rainy season in 1999 (194/052-053—November 7, 1999;

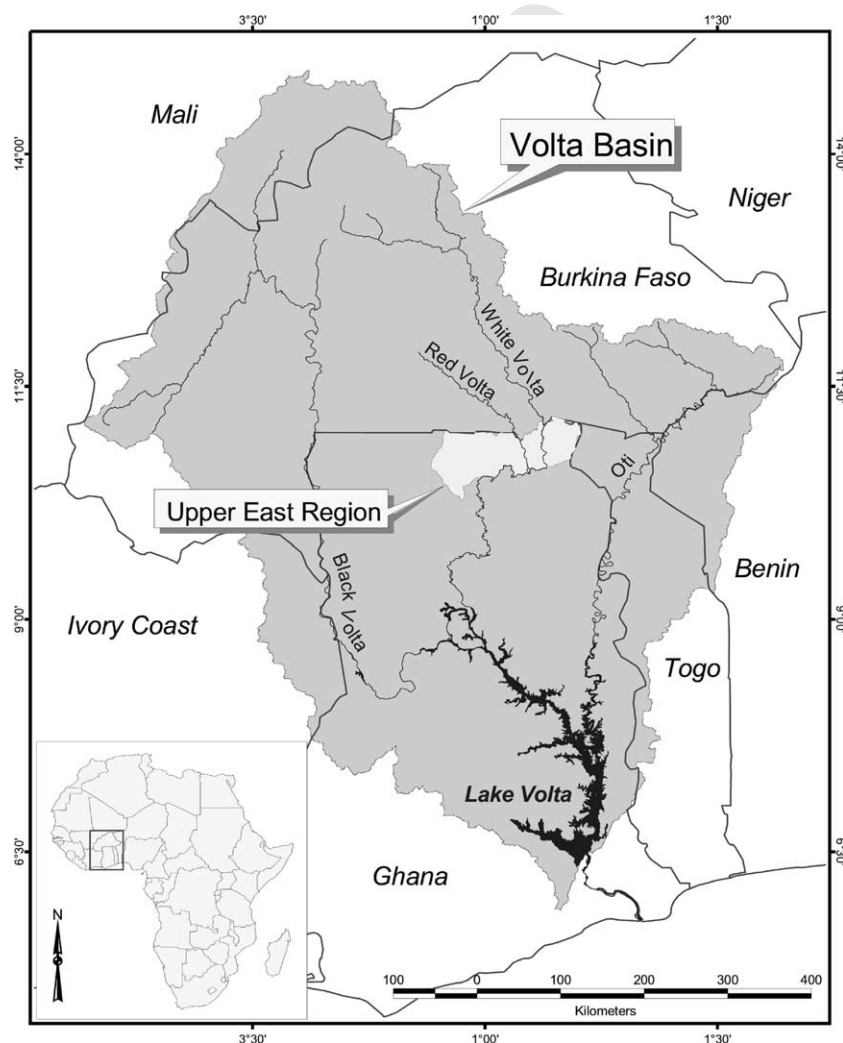


Fig. 1. The Upper East Region of Ghana within the Volta Basin, West Africa.

84 195/052—October 13, 1999), while the fourth image,
 85 which only covers a small part of the Region, was made
 86 during the following dry season (195/053 acquired Feb-
 87 ruary 2, 2000). The rainy season of 1999 was typical of a
 88 comparably wet year when the reservoirs were expected
 89 to be full, so the satellite images capture the reservoirs'
 90 maximum surface areas. The reservoirs were extracted
 91 through maximum likelihood classification (Mather,
 92 1999) of the Landsat ETM bands 3, 4 and 5. The visible
 93 red and the two infrared bands are well-suited for the
 94 detection and delineation of open water bodies. Gener-
 95 ally, the spectral reflectance curve of water shows a
 96 reduction in reflection with increasing wavelength
 97 (Mather, 1999). In bands 4 and 5, the spectral reflec-
 98 tance curve of water is low, whereas the reflectances of
 99 soils and healthy vegetation are higher in these bands.
 100 This contrast allows us to distinguish between open
 101 water and the surrounding land surface. The inclusion
 102 of the visible band enhances the contrast between water
 103 and soils (Kite and Pietroniro, 2000).

104 Open surface water is a land cover type that has a
 105 wide range of reflectance patterns (Meijering et al.,
 106 1994; Kondratyev and Filatov, 1999). This variety is
 107 caused by three processes, which are surface reflectance,
 108 volume reflectance, and bottom reflectance (Mather,
 109 1999). The scattered component of the surface reflection
 110 mainly consists of shorter wavelengths, particularly
 111 from the visible part of the spectrum. The infrared part
 112 of the spectrum is strongly absorbed. Certain sun-to-
 113 sensor constellations and the roughness of the water sur-
 114 face can cause sunglint, while under calm conditions the
 115 reflectance may be specular (Mather, 1999; Meijering
 116 et al., 1994). The presence of standing or floating vege-
 117 tation causes a steep ascent in reflectance from the visi-
 118 ble to the infrared part of the spectrum, which is
 119 characteristic for healthy vegetation (Horler et al.,
 120 1983). The signal of the volume reflection is influenced
 121 by turbidity, dissolved matter, the trophic status, and al-
 122 gae content of the water. The degree to which the vol-
 123 ume reflection contributes to the total reflectance
 124 signals of water bodies depends on the penetration
 125 depth of light, which decreases from 10 m at 0.5–
 126 0.6 μm to less than 10 cm in the range between 0.8 and
 127 1.1 μm (Meijering et al., 1994). The influence of bottom
 128 reflectance to the total signal is equally wavelength
 129 dependent and mainly originates from the deeper pene-
 130 trating, shorter wavelengths.

131 Taking into account the range of reflectance patterns
 132 of surface water, the classification was performed with
 133 various sub-classes of water, such as clear water, turbid
 134 water, water with algae, etc., that were later regrouped
 135 again. Freshly burned areas from bushfires and cloud
 136 shadows show a spectral overlap with the response pat-
 137 terns of water and are therefore a constraint for remote
 138 sensing of open water in the semi-arid tropics (Koutsias
 139 et al., 2000). To prevent bushfires and cloud shadows

140 from being misclassified as surface water, these two land
 141 cover types were also classified and later discarded
 142 (Liebe, 2002). A 3×3 median filter was applied to re-
 143 move single pixels and to fill single-pixel-holes. The
 144 remaining river segments were manually deleted. The
 145 obtained reservoir map was georeferenced in quadratic
 146 mode using 27 reference points, which were taken from
 147 1:50,000 topographic maps. Road intersections were
 148 preferred as reference points, but in less developed areas
 149 more variable features such as river mouths, and bridges
 150 intersecting rivers had to be used, leading to an RMS of
 151 30.2 m.

152 The classification returned a total of 504 reservoirs
 153 with a total acreage of 3408 ha (Table 1). Given the
 154 30 m resolution of the Landsat imagery, the likelihood
 155 of incorrectly identifying very small features is high.
 156 Three hundred and forty eight reservoirs with acreage
 157 of less than 1 ha were therefore deleted from the dataset.
 158 The two commercially operated reservoirs, Tono and
 159 Vea were also discarded. These are managed by the Irr-
 160 igation Company Of Upper Region Ltd., ICOUR, which
 161 monitors and records the fluctuations of their storage
 162 volumes. The remaining 154 reservoirs with a total acre-
 163 age of 999.54 ha are those herein referred to as small res-
 164 ervoirs, ranging from 1 to 35 ha. Because image
 165 acquisition took place in November 1999, and field
 166 work from December 2002 to February 2003, there
 167 was no direct comparison possible between reservoir
 168 outlines as found in the field and in the images. There-
 169 fore, the commonly used pixel based user and producer
 170 accuracy matrix was not compiled. Instead, an indicator
 171 of good user accuracy is the fact that 50% of the areas
 172 classified as reservoirs were visited and in each instance
 173 a reservoir was indeed found.

174 A second quality indicator of the classification is gi-
 175 ven by the comparison of the satellite derived and the
 176 ground based surface areas estimates. On average, the
 177 field measurements were found to be 13% smaller than
 178 the satellite based area estimates, with a standard devia-
 179 tion of 16%. This is not surprising given that field data
 180 collection took place later in the dry season. Reservoirs
 181 found to be significantly smaller than classified can be
 182 explained by intensive irrigation use. The few reservoirs
 183 that were found larger than classified showed extensive

Table 1
Results of the satellite based reservoir inventory

Reservoirs	Number	Total surface area (ha)	Percent of total area
All	504	3408	100
<i>Thereof</i>			
Tono	1	1894	56
Vea	1	435	13
“Small”	154	999.54	29
<1 ha	348	79.46	2

184 areas infested with dense herbaceous water plants in the
 185 shallow tail parts of the reservoirs, which were not clas-
 186 sified as water due to the dense biomass. The overall cor-
 187 relation of satellite to field measured reservoir surface
 188 area has an R^2 of 0.88.

189 4. Distribution of reservoirs

190 The distribution of reservoirs pertains both to the sta-
 191 tistical distribution of specific properties such as size,
 192 and to their distribution over space. Fig. 2 shows the dis-
 193 tribution of reservoir sizes. On a semi-log plot, reservoir
 194 sizes show a relatively smooth linear distribution. The
 195 absence of significant breaks in the distribution is an
 196 indicator for the uniformity of the Upper East's topog-
 197 raphy. Also interesting is the frequency distribution of
 198 reservoir sizes. The smallest reservoirs occur at the high-
 199 est frequency, and with increasing surface area their
 200 frequency decreases exponentially (Liebe, 2002). This
 201 distribution of reservoir sizes mirrors to some extent
 202 general laws in stream morphometry, such as Horton's
 203 law of stream numbers (Knighton, 1998). Horton's law
 204 refers to the expected number of streams of a certain
 205 stream order, where lower order stream segments are
 206 more frequent than those of higher orders. In contrast
 207 to the stream numbers in Horton's law, however, the
 208 reservoirs do not follow a power law distribution (Tar-
 209 boton et al., 1988).

210 The spatial distribution of reservoirs and sizes is
 211 shown in Fig. 3. It depicts the location of reservoirs in
 212 relation to the Upper East's topography, which was
 213 compiled from contour lines digitized from 1:50,000
 214 topographic map sheets. The reservoir sizes were catego-
 215 rized, based on a sampling frame that was designed for
 216 ground truthing purposes. The total of 154 reservoirs
 217 were divided into three approximately equal size groups.
 218 Category One has 51 reservoirs of 1–2.79 ha, Category
 219 Two has 53 of 2.88 to 6.93 ha, and Category Three
 220 has 50 reservoirs of 7.02–35 ha. In Fig. 3, the topogra-
 221 phy is shown as steps of one standard deviation
 222 (35.33 m) from the Region's mean elevation of 197 m

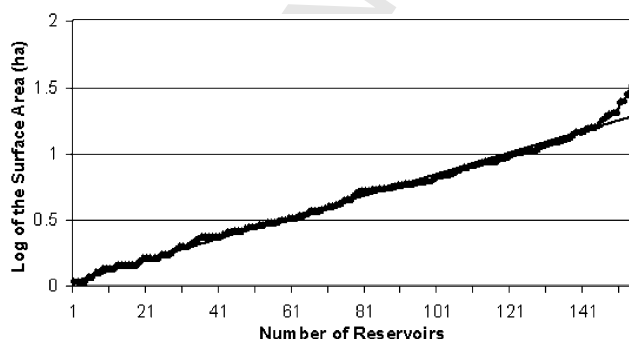


Fig. 2. Reservoir size distribution.

(a.s.l.). Most of the reservoirs are located above the
 mean elevation and yet close to the mean elevation or
 the fringe area to the next standard deviation step. Con-
 ditions favorable for dam construction are indicated by
 elevation level boundaries that wriggle back and forth,
 following coves and ledges. Noticeable is the two-sided-
 ness of the elevation distribution with the western half
 mainly occupying lower elevation ranges, whereas the
 eastern half lies mainly above the mean elevation. The
 lower elevations show a lack of reservoirs in comparison
 to the upper elevations. These lower areas are mainly
 occupied by the extensive floodplains of the Red Volta,
 the White Volta, and their tributaries. The occurrence of
 river blindness used to make the floodplains an unfavor-
 able habitat and they are much less populated than the
 higher ranges. The higher eastern half of the region
 has large areas well-suited for settlement and dam con-
 struction. Here we find most of the larger reservoirs.

Although drainage pattern and valley shape are
 important for dam construction, the proximity of a po-
 tential dam site to a road is also relevant. Easy access to
 transportation facilitates construction. Once a reservoir
 is in use, perishable produce can be brought to market
 quickly. Morphological regularities can also be deduced
 from the reservoir distribution. The higher frequency of
 lower order stream segments also indicates a greater
 number of potential dam sites and vice-versa. Forest re-
 serves pose a further constraint to dam construction and
 administrative boundaries also seem to play a role in
 reservoir distribution (Liebe, 2002). The occurrence of
 reservoirs can thus be termed “semi-natural”, as their
 existence relies to a great extent on the people's decisions
 as to where to make use of the available opportunities in
 the natural landscape.

5. Bathymetrical survey and derivation of area–volume relations

As it is not possible to determine a priori the mini-
 mum number of reservoirs to be surveyed to yield a ro-
 bust area–volume relation, a large sample of almost 40%
 of the total population was surveyed. This sample con-
 tained twenty reservoirs from size Categories One and
 Three, and 21 reservoirs from Category Two. The out-
 line of each reservoir was mapped using GPS. Bathy-
 metric maps were compiled on the basis of 25–171
 GPS-referenced depth measurements, which were taken
 with a telescopic stadia rod from a boat. In order to get
 accurate representations of the reservoirs' shapes, care
 was taken to capture the deepest point, as well as the
 submerged streambed. This was achieved by oversam-
 pling those parts of the reservoirs where the deepest
 points were expected to be found, taking into account
 the apparent depth distribution of the in situ measure-
 ments. Because a reservoir is a dammed stream, the

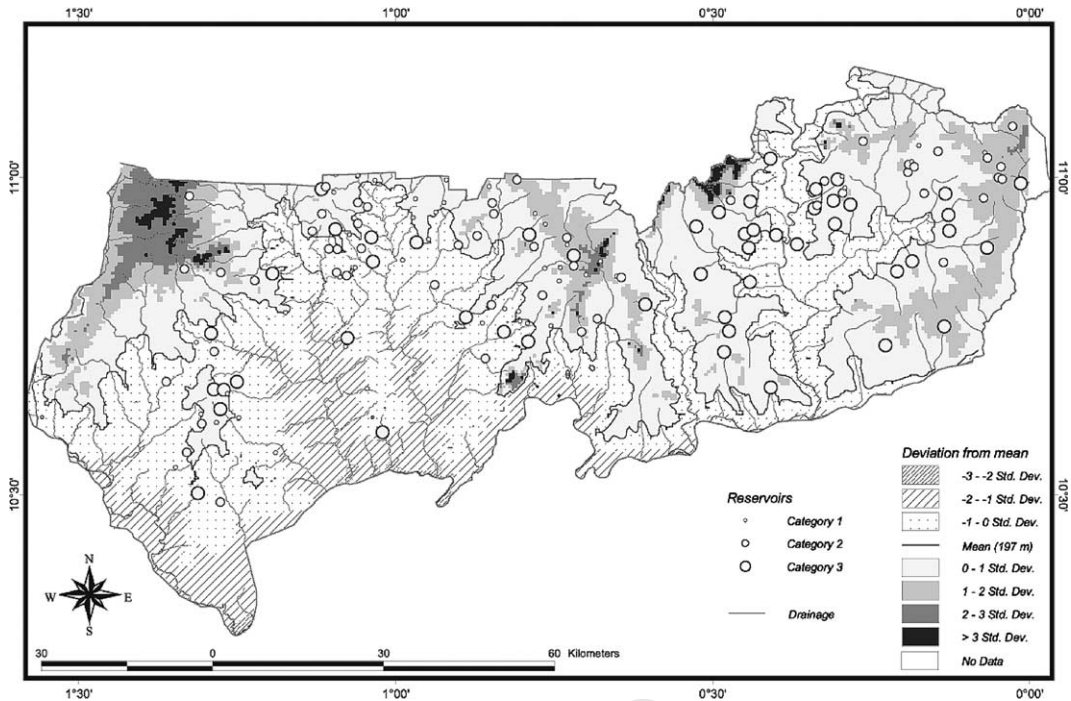


Fig. 3. Distribution of reservoirs with respect to the Upper East Region's topography.

329 second term of the equation, $1.4367 \cdot \log A$, is close to
 330 the expected $3/2 \log A$ in (4). Additional information is
 331 contained in the first term of the equation. Comparing
 332 (4) and (5) gives

$$334 \quad -2.067 = \log \frac{1}{6} + \log \frac{1}{f} = -0.778 + \log \frac{1}{f} \quad (6)$$

335 or

$$337 \quad f = 19.45. \quad (7)$$

338 This means that the depth d of a reservoir is $\approx 1/20$ of its
 339 characteristic length.

340 To use Eq. (5) properly for the determination of reser-
 341 voir volumes, the precision of the prediction needs to
 342 be known. The goodness of fit between measured and
 343 modeled volumes can be evaluated. The widely used
 344 model efficiency measure of Nash and Sutcliffe (1970)
 345 indicates that the model explains 97.5% of the measured
 346 variance. Despite the variety of reservoir shapes, the de-
 347 rived equations are highly consistent. The sizes of the
 348 measured reservoirs are finally summarized in the
 349 equation:

$$352 \quad \text{Volume} = 0.00857 \cdot \text{Area}^{1.4367} \text{ [m}^3\text{]}. \quad (8)$$

353 Eq. (8) and field measurements show that, at full capac-
 354 ity, the Upper East's 154 small reservoirs can capture up
 355 to $185 \times 10^6 \text{ m}^3$ of water.

356 6. Conclusions and outlook

357 This research begins to fill the gaps in our knowledge
 358 concerning the physical characteristics of small reser-
 359 voirs as found in the Upper East Region of Ghana. First
 360 we established that surface areas of small reservoirs can
 361 reliably be mapped on the basis of satellite images. Here,
 362 Landsat 7 satellite images were used which in the rainy
 363 season may be hindered by clouds. Given the ease with
 364 which open water can be mapped with RADAR images
 365 (Hess et al., 1995; van de Giesen, 2001), future applica-
 366 tions will involve the use of RADAR satellites that are
 367 not hindered by clouds and, thereby, allow year-round
 368 monitoring of reservoir surfaces.

369 We also found a relation between reservoir volumes
 370 and their surface areas. Eq. (8) relates reservoir areas di-
 371 rectly to storage volumes with high precision. Combin-
 372 ing Eq. (8) with periodical satellite-based reservoir
 373 area measurements would allow us to build a cost effec-
 374 tive monitoring system. Such a system is both valuable
 375 for hydrological research, as it provides indications of
 376 surface runoff in ungauged basins, and for water manag-
 377 ers who need estimates of water availability in the Upper
 378 East Region of Ghana.

379 Although their total storage capacity of $185 \times 10^6 \text{ m}^3$
 380 is modest, these reservoirs form a set of well-distributed
 381 and easily accessible water sources that can be used for

agriculture, domestic use, and livestock. Small dams 382
 help to reduce the people's vulnerability to drought 383
 and improve their livelihoods. Their modest size mini- 384
 mizes the negative environmental impacts often associ- 385
 ated with construction of larger dams. This research 386
 will facilitate management of these diffuse systems and 387
 ensure their sustainable use. 388

Acknowledgments

The presented research was carried out as part of the 390
 GLOWA Volta Project. We gratefully acknowledge 391
 financial support of the German Federal Ministry of 392
 Education and Research (BMBF) as main sponsor of 393
 the GLOWA project and the North Rhine-Westphalia's 394
 Ministry of Science and Research (MWF). The 395
 GLOWA-Volta project is de facto a research network 396
 and we express our thanks for the scientific inputs from 397
 our partners: IMK-IFU, Meteorology Department 398
 Wageningen University, Savanna Agricultural Research 399
 Institute (CSIR), Water Research Institute (CSIR), 400
 Institute for Statistical, Social and Economic Research 401
 (University of Ghana), Meteorological Services Depart- 402
 ment (Ghana), Institut de l'Environnement et des 403
 Recherches Agricoles (Burkina Faso). 404

References

- Asenso-Okyere, W.K., Twum-Baah, K.A., Kasanga, A., Anum, J., 406
 Pörtner, C., 2000. Ghana living standards survey. Report of the 407
 Fourth Round (GLSS 4), Ghana Statistical Service, Accra. 408
 Eldridge, R.H., 1957. A synoptic study of West African disturbance 409
 lines. *Quart. J. Roy. Meteorol. Soc.* 83, 303–314. 410
 Friesen, J., 2002. Spatio-temporal rainfall patterns in Northern Ghana. 411
 Diploma thesis, Department of Geography, University of Bonn, 412
 Bonn. 413
 Ghana Statistical Service, 2000. Poverty trends in Ghana in the 1990s. 414
 Ghana Statistical Service, Accra. 415
 Hayward, D., Oguntoyinbo, J., 1987. *Climatology of West Africa.* 416
 Hutchinson, London. 417
 Hess, L.L., Melack, J.M., Filoso, S., Wang, Y., 1995. Delineation of 418
 inundated area and vegetation along the Amazon floodplain with 419
 the SIR-C synthetic aperture radar. *IEEE Trans. Geosci. Remote* 420
Sensing 33, 896–903. 421
 Horler, D.N.H., Dockray, M., Barber, J., 1983. The red edge of plant 422
 leaf reflectance. *Int. J. Remote Sensing* 4 (2), 273–288. 423
 Kite, G., Pietroniro, A., 2000. Remote sensing of surface water. In: 424
 Schultz, G.A., Engman, E.T. (Eds.), *Remote Sensing in Hydrology* 425
and Water Management. Springer, Heidelberg, pp. 217–238. 426
 Knighton, D., 1998. *Fluvial Forms and Processes. A new Perspective.* 427
 Arnold, London. 428
 Kondratyev, K.Y., Filatov, N.N. (Eds.), 1999. *Limnology and Remote* 429
Sensing. A Contemporary Approach. Springer-Praxis Series in 430
Remote Sensing. Springer, Berlin. 431
 Koutsias, N., Karteris, M., Chuvieco, E., 2000. The use of intensity- 432
 hue-saturation transformation of Landsat-5 thematic mapper data 433
 for burned land mapping. *PE&RS* 66 (7), 829–839. 434
 Liebe, J., 2002. Estimation of water storage capacity and evaporation 435
 losses of small reservoirs in the Upper East Region of Ghana. 436

- 437 Diploma thesis, Department of Geography, University of Bonn,
438 Bonn.
- 439 Mather, P.M., 1999. Processing of Remotely-Sensed Images. An
440 Introduction. Wiley & Sons, Chichester.
- 441 Meijering, A.M.J., Brouwer, H.A.M., Mannaert, C.M., Valenzuela,
442 C.R., 1994. Introduction to the use of geographic information
443 systems for practical hydrology. UNESCO International Hydro-
444 logical Programme IHP-IV M 2.3, Publication No. 23. ITC,
445 Enschede.
- 446 Nash, J.E., Sutcliffe, J.V., 1970. River flow forecasting through
447 conceptual models. 1. A discussion of principles. *J. Hydrol.* 10,
448 282–290.
- Tarboton, D.G., Bras, R.L., Rodriguez-Iturbe, I., 1988. The fractal 449
nature of river networks. *Water Resour. Res.* 24 (8), 1317–1322. 450
- van de Giesen, N., Kunstmann, H., Jung, G., Liebe, J., Andreini, M., 451
Vlek, P.L.G., 2002. The GLOWA-Volta project: integrated assess- 452
ment of feedback mechanisms between climate, landuse, and 453
hydrology. *Adv. Global Change Res.* 10, 151–170. 454
- Windmeijer, P.N., Andriess, W. (Eds.), 1993. Inland Valleys in West 455
Africa: An Agro-Ecological Characterization of Rice-Growing 456
Environments, ILRI Publication 52. International Institute for 457
Land Reclamation and Improvement, Wageningen. 458
459

UNCORRECTED PROOF