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Retrievals of tropospheric ozone profiles from the synergism of AIRS and OMI: methodology and validation

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Abstract. The Tropospheric Emission Spectrometer (TES) on the A-Train Aura satellite was designed to profile tropospheric ozone and its precursors, taking measurements from 2004 to 2018. Starting in 2008, TES global sampling of tropospheric ozone was gradually reduced in latitude, with global coverage stopping in 2011. To extend the record of TES, this work presents a multispectral approach that will provide O₃ data products with vertical resolution and measurement error similar to TES by combining the single-footprint thermal infrared (TIR) hyperspectral radiances from the Aqua Atmospheric Infrared Sounder (AIRS) instrument and the ultraviolet (UV) channels from the Aura Ozone Monitoring Instrument (OMI). The joint AIRS+OMI O₃ retrievals are processed through the MULTI-SpEctra, MULTI-SpECies, MULTI-SEnsors (MUSES) retrieval algorithm. Comparisons of collocated joint AIRS+OMI and TES to ozonesonde measurements show that both systems have similar errors, with mean and standard deviation of the differences well within the estimated measurement error. AIRS+OMI and TES have slightly different biases (within 5 parts per billion) vs. the sondes. Both AIRS and OMI have wide swath widths (~1650 km for AIRS; ~2600 km for OMI) across satellite ground tracks. Consequently, the joint AIRS+OMI measurements have the potential to maintain TES vertical sensitivity while increasing coverage by 2

orders of magnitude, thus providing an unprecedented new data set with which to quantify the evolution of tropospheric ozone.

1 Introduction

Long-term records of the vertical distribution of ozone are essential for quantifying the impact of changes in tropospheric ozone on air quality and climate, driven recently by rapid industrialization in Asia concurrent with reductions in ozone precursor emissions in North America and Europe (Jacob et al., 1999; Wild and Akimoto, 2001; Akimoto, 2003; Worden et al., 2008, 2011; Fischer et al., 2011). The A-Train Aura satellite has played an important role in quantifying the atmospheric ozone and advancing our understanding of the processes controlling its distribution. The Dutch–Finnish Ozone Monitoring Instrument (OMI) measures ultraviolet (UV) radiances, which are used to infer a number of species including ozone profiles and columns (Levelt et al., 2006a, b, 2018; Liu et al., 2010a, b; Huang et al., 2017). These measurements have been used in a number of assimilation systems to constrain both stratospheric and tropospheric ozone distributions (Stajner et al., 2008; Pierce et al., 2009; Huang et al., 2013; Inness et al., 2013; Wargan et al., 2015; Olsen

et al., 2016). OMI ozone columns have been used to understand both tropical ozone variability (Chandra et al., 2007; Ziemke et al., 2007) and high-latitude ozone, including the unprecedented Arctic ozone loss in 2011 (Manney et al., 2011). The Aura Tropospheric Emission Spectrometer (TES) has a spectral resolution of 0.1 cm^{-1} , the highest infrared spectral resolution among any current nadir sounder, which enables estimation of tropospheric ozone profiles and precursors. TES has advanced a number of Aura science objectives, including detection of tropospheric ozone trends over Asia (Lamsal et al., 2011; Verstraeten et al., 2015), the influences of long-range pollution transport on surface ozone (Parrington et al., 2008, 2009), and the tropospheric ozone response to stratospheric circulation (Neu et al., 2014). The TES record has also played an important role in evaluating chemistry–climate model simulations of present-day ozone distributions and their ozone radiative forcing as part of the Intergovernmental Panel on Climate Change (IPCC) Fifth Assessment Report (AR5; Bowman et al., 2013; Shindell et al., 2013; Young et al., 2013; IPCC, 2014) and in providing constraints on the tropospheric chemistry through data assimilation (Miyazaki et al., 2012, 2014, 2015). TES global observations are limited to a roughly 5-year period (2005–2009) due to instrument aging. TES global sampling of tropospheric ozone was gradually reduced starting in 2008, with global observations ceasing altogether in 2011. Consequently, TES's well-validated global-survey record of tropospheric ozone (H. M. Worden et al., 2007; Nassar et al., 2008; Boxe et al., 2010; Verstraeten et al., 2013; Bella et al., 2015) ended in 2011.

The synergy of combining UV and ultra-spectral thermal infrared (TIR) radiances provides an approach to measuring lower-tropospheric ozone, a key objective of air quality remote sensing (J. Worden et al., 2007; Landgraf and Hasekamp, 2007; Costantino et al., 2017). This capability was demonstrated by Fu et al. (2013) for joint TES+OMI and Cuesta et al. (2013, 2018) for joint Infrared Atmospheric Sounding Interferometer (IASI) and Global Ozone Monitoring Experiment 2 (GOME-2). Ozone profiles from joint TES+OMI retrievals are a part of the standard Earth Observing System (EOS) Aura products from the time period 2005 to 2008, the time period when neither the degradation of TES instrument nor the row anomaly of OMI pixels (Huang et al., 2017; Schenkeveld et al., 2017; Levelt et al., 2018), which provide measurements collocated to TES measurements, played a role.

In this work, we demonstrate that joint Atmospheric Infrared Sounder (AIRS) and OMI retrievals can extend the Aura EOS TES standard Level 2 tropospheric ozone concentration vertical profile products. The retrieved ozone profiles harnessing the Level 1B radiances from AIRS and OMI measurements have vertical resolution and error characteristics similar to the TES instrument on Aura and the prospect of vastly increased spatial coverage.

2 TES, AIRS, OMI, and ozonesonde measurements

The NASA A-Train satellites (Aqua, Aura, Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observation (CALIPSO), CloudSat, Orbiting Carbon Observatory-2 (OCO-2)) are providing long-term global measurements of the land surface, biosphere, atmosphere, and oceans of the Earth in a near-polar, sun-synchronous, $\sim 700\text{ km}$ altitude orbit whose ascending node has an Equator-crossing time of around 13:30 local time. The measurements of three nadir-viewing instruments in the A-Train satellites – including Aura-TES, Aura-OMI, and Aqua-AIRS – play essential roles in quantifying atmospheric composition, including O_3 and a suite of trace gases, to advance understanding of air quality and climate science.

TES is a Fourier transform spectrometer (FTS) that measures the double-sided interferograms of TIR radiances emitted and absorbed by Earth's surface, gases, and particles in the atmosphere (Beer et al., 2001). Although TES has both the nadir and limb views, nadir has been the primary scanning geometry used to obtain full vertical and horizontal coverage of Earth's atmosphere. In nadir mode, TES measurements cover four optical filter bands ($650\text{--}900$, $950\text{--}1150$, $1100\text{--}1325$, and $1900\text{--}2250\text{ cm}^{-1}$) with a constant spectral resolution of 0.1 cm^{-1} and a ground pixel size of $5.3 \times 8.5\text{ km}^2$. The $950\text{--}1150\text{ cm}^{-1}$ spectral region includes high-density absorption features of the ozone ν_3 band (the strongest fundamental band) and minor absorption from interfering species. The ν_3 band has been exploited in the tropospheric O_3 soundings by a suite of TIR satellite-borne, nadir-viewing instruments, including AIRS (Susskind et al., 2003, 2014; Wei et al., 2010), Cross-track Infrared Sounder (CrIS) (Gambacorta et al., 2013), and IASI (Boynard et al., 2009, 2016; Dufour et al., 2012; Oetjen et al., 2014, 2016), as well as the solar occultation satellite-borne (Bernath et al., 2005; Bernath, 2017), balloon-borne (Toon, 1991; Fu et al., 2007a), and ground-based (Hannigan et al., 2011) FTSs that quantify the stratospheric ozone layer and the species playing an essential role in the stratospheric ozone chemistry (Fu et al., 2007b, 2009, 2011; Sung et al., 2007; Wunch et al., 2007; Allen, 2009; Boone, 2013; Nassar, 2013; Griffin et al., 2017). The spectral resolution of TES (resolving power (RP): 10 500) is significantly higher than the existing TIR, including AIRS (RP: 1200), CrIS (RP: 816), and IASI (RP: 5250). Benefiting from the Aura afternoon orbit, TES takes measurements around local noontime when the atmosphere–land thermal contrast is typically higher than other times of the day. Taking the spectral coverage, spectral resolution, and noise performance into account, the vertical sensitivity of TES and other satellite sensors (AIRS alone, OMI alone) is quantified in Sect. 3.2. It shows that TES has the sensitivity to distinguish between the upper- and lower-tropospheric O_3 .

AIRS is a grating spectrometer that measures the Earth's TIR emission in the spectral range of $650\text{--}2665\text{ cm}^{-1}$ (Aumann et al., 2003). It is a cross-track scanning instrument

providing measurements with daily global coverage. AIRS atmospheric measurements in the ozone ν_3 band provide sensitivity for estimating atmospheric ozone column density. The currently operational AIRS version 6 retrieval algorithm (Susskind et al., 2003, 2014) estimates the temperature, humidity, and atmospheric composition products using the 45 km resolution Level 2 cloud-cleared radiance products for weather prediction and environmental monitoring. In order to fully exploit the spatial resolution of AIRS measurements, our joint AIRS+OMI ozone retrievals use single-footprint (i.e., non-cloud-cleared) Level 1b AIRS infrared radiances with a spatial resolution of ~ 13.5 km nadir horizontal resolution.

OMI is a nadir-viewing push broom ultraviolet–visible (UV-VIS) imaging spectrograph that measures backscattered radiances covering the 270–500 nm wavelength range (Levelt et al., 2006a, b) and captures the absorption features of the ozone Hartley and Huggins bands that are clearly present in the 270–310 nm (mainly for stratospheric ozone information) and 310–330 nm (mainly for tropospheric ozone information) spectral regions. The ground pixel size of OMI measurements at nadir position is about 13 km (along the ground track of spacecraft) \times 24 km (across the track) when using the spectral radiances 310–330 nm. Since 2009, row anomaly and stray-light issues have affected the quality of some OMI pixels (Huang et al., 2017; Schenkeveld et al., 2017; Levelt et al., 2018). Following 2009, for retrieval, the MUlti-SpEctra, MUlti-SpEcies, MUlti-SEnsors (MUSES) algorithm uses the measured radiances from the quality-assured OMI off-nadir pixels and the corresponding collocated AIRS measurements.

The World Ozone and Ultraviolet radiation Data Centre (WOUDC, <http://www.woudc.org>, 4 October 2018) ozonesonde measurements provide in situ data from the surface to the stratosphere (about 35 km) with vertical resolution of ~ 150 m and accuracy of 5 % (Witte et al., 2017, 2018; WMO/GAW, 2017). These data fill a critical need for the validation of ozone profiles measured by spaceborne remote-sensing instruments (Thompson et al., 2017). The ozonesonde sensor has a dilute solution of potassium iodide to produce a weak electrical current proportional to the ozone concentration of the sampled air (Komhyr et al., 1995). To examine the performances of remote-sensing measurements, we applied the following coincidence criteria to determine sonde–AIRS+OMI: (1) mean cloud optical depth < 2.0 , (2) cloud fraction within OMI field of view < 30 %, (3) both satellite ground pixel–sonde distances < 300 km, (4) solar zenith angle $< 80^\circ$, and (5) daytime measurements with a time difference < 4 h. In order to determine the sonde–TES pairs, we applied the criteria (1), (3), (4), and (5), and exclude criterion (2) because the TES retrieval does not use information from OMI measurements. As a result, for the 2006 time frame, we obtained 424 sonde–AIRS–OMI triads and 556 sonde–TES measurement pairs.

3 Retrieval algorithms and retrieval characteristics

The joint AIRS+OMI ozone profile is produced from the MUSES retrieval algorithm, crafted to accommodate multiple instruments, including joint TES+OMI O₃ retrievals (Fu et al., 2013); joint CrIS+TROPOMI (TROPOspheric Monitoring Instrument) carbon monoxide (CO) profiling (Fu et al., 2016); joint TES+Microwave Limb Sounder (MLS) CO retrievals (Luo et al., 2013); and AIRS CH₄, HDO, H₂O, and CO retrievals (Worden et al., 2018; Kulawik et al., 2018). These atmospheric composition products, with characteristics of vertical resolution and error similar to TES standard Level 2 data, have the potential to extend the Aura-TES atmospheric composition Earth science data records (ESDRs), continuing the climate and air quality science enabled by TES measurements. The development of the MUSES algorithm leverages a suite of existing atmospheric composition retrieval algorithms, especially forward radiative transfer models, including the Earth Limb and Nadir Operational Retrieval (ELANOR) of the TES operational algorithm (Worden et al., 2004; Clough et al., 2006; Kulawik et al., 2006a, b; Bowman et al., 2006; Eldering et al., 2008) for simulation of TIR radiances and Jacobians (Fu et al., 2013, 2016); the U.S. Smithsonian Astrophysical Observatory (SAO) OMI OZone PROFile (RROFOZ) algorithm (Liu et al., 2010a, b) for simulation of UV radiances and Jacobians of Hartley and Huggins bands (Fu et al., 2013; Worden et al., 2013); and the full-physics OCO-2 algorithm (O’Dell et al., 2012, 2018; Connor et al., 2016; Crisp et al., 2012, 2017; Eldering et al., 2017) for simulation of short-wavelength infrared radiances and Jacobians (Fu et al., 2016).

3.1 Joint AIRS+OMI ozone profile retrievals

The retrieval methodology is based on the optimal-estimation (OE) method (Rodgers, 2000), which minimizes the differences between observed and measured radiances subject to a priori knowledge, i.e., mean and covariation of the atmospheric–cloud–surface state, to infer the “optimal” or maximum a posterior (MAP). Numerically, the MAP state vector $\hat{\mathbf{x}}$, which represents the concentration of atmospheric trace gases and ancillary parameters, is computed by minimizing the following cost function with respect to \mathbf{x} :

$$C(\mathbf{x}) = \|\mathbf{x} - \mathbf{x}_a\|_{\mathbf{S}_a}^2 + \|\mathbf{L}_{\text{obs}} - \mathbf{L}_{\text{sim}}\|_{\mathbf{S}_e}^2. \quad (1)$$

Equation (1) is a sum of quadratic functions representing a weighted Euclidean norm ($\|b\|_a^2 = b^T a b$), with the first term accounting for the difference between the retrieval vector \mathbf{x} and a priori state \mathbf{x}_a , inversely weighted by the a priori covariance matrix \mathbf{S}_a , and with the second term representing the difference between the observed \mathbf{L}_{obs} and simulated \mathbf{L}_{sim} radiance spectra inversely weighted by the measurement error covariance matrix \mathbf{S}_e .

Under the assumption that measurement error between AIRS and OMI is uncorrelated, Eq. (1) can be written as

$$C(\mathbf{x}) = \|\mathbf{x} - \mathbf{x}_a\|_{\mathbf{S}_a^{-1}}^2 + \underbrace{\|\mathbf{L}_{\text{obs_AIRS}} - \mathbf{L}_{\text{sim_AIRS}}\|_{\mathbf{S}_{\epsilon_AIRS}^{-1}}^2}_{\text{AIRS}} + \underbrace{\|\mathbf{L}_{\text{obs_OMI}} - \mathbf{L}_{\text{sim_OMI}}\|_{\mathbf{S}_{\epsilon_OMI}^{-1}}^2}_{\text{OMI}}. \quad (2)$$

The joint retrieval algorithm iteratively updates the state vector based upon a trust-region Levenberg–Marquardt (LM) optimization algorithm (Moré, 1977; Bowman et al., 2006) to minimize the cost function in Eq. (2):

$$\mathbf{x}_{i+1} = \mathbf{x}_i + \left[\gamma_i \mathbf{W}^T \mathbf{W} + \mathbf{S}_a^{-1} + \underbrace{\mathbf{K}_{\text{AIRS}}^T \mathbf{S}_{\epsilon_AIRS}^{-1} \mathbf{K}_{\text{AIRS}}}_{\text{AIRS}} + \underbrace{\mathbf{K}_{\text{OMI}}^T \mathbf{S}_{\epsilon_OMI}^{-1} \mathbf{K}_{\text{OMI}}}_{\text{OMI}} \right]^{-1} \times \left[\mathbf{S}_a^{-1} (\mathbf{x}_a - \mathbf{x}_i) + \underbrace{\mathbf{K}_{\text{AIRS}}^T \mathbf{S}_{\epsilon_AIRS}^{-1} \Delta \mathbf{L}_{\text{AIRS}}}_{\text{AIRS}} + \underbrace{\mathbf{K}_{\text{OMI}}^T \mathbf{S}_{\epsilon_OMI}^{-1} \Delta \mathbf{L}_{\text{OMI}}}_{\text{OMI}} \right], \quad (3)$$

where the parameter γ_i is called the LM parameter, \mathbf{W} is a nonzero scaling matrix, $\mathbf{K}_{\text{instrument}}$ is the Jacobian matrix representing instrument sensitivity of spectral radiances to the atmospheric state, and $\Delta \mathbf{L}$ is the difference between observed and simulated spectral radiances. The computation of the γ_i value and \mathbf{W} follow Sects. 5.5 and 6.3 of Moré (1977), utilizing the fitting residuals and \mathbf{K} from the space instruments as input parameters. The $\gamma_i \mathbf{W}^T \mathbf{W}$ term, the core of the trust-region LM optimization algorithm, plays the crucial role in balancing the convergence speed and robustness. Under large γ_i , the step size computation is similar to a steepest-descent algorithm, which has a lower convergence rate, and under low γ_i the step computation is towards a Gauss–Newton approach.

To simulate TIR spectral radiances \mathbf{L} and Jacobians \mathbf{K} in TIR and UV spectral regions (Table 1), the joint AIRS+OMI retrieval adopts the forward models of the joint TES+OMI retrievals (Fu et al., 2013) with necessary revisions to incorporate the AIRS specifications (spectral range, signal-to-noise ratios (SNRs), viewing geometry, and spectral response function) (Pagano et al., 2003; Strow et al., 2003).

The joint AIRS+OMI retrievals start with the list of the fitting parameters, a priori values, and a priori variance shown in Table 2. In addition to the initial guess for the trace gas concentration (O_3 , H_2O , and CO_2), the initial guess for auxiliary parameters used in the simulation of AIRS radiances (including temperature profile, surface temperature and emissivity, and cloud extinction and cloud top pressure) are also retrieved from AIRS radiances in order to take into account their spectral signatures in the O_3 spectral regions. The joint AIRS+OMI algorithm incorporated a suite of treatments in order to optimize the spatial resolution, retrieval stability,

data throughput, and consistency to TES data products (version 6): (1) when the clouds travel across its field of view, a space sensor for atmospheric composition measurements often faces the challenge of obtaining high-precision and high-accuracy measurements of the trace gas vertical distribution due to the interference among retrieval parameters, and MUSES algorithm uses single-footprint AIRS Level 1B radiances in the retrievals (Irion et al., 2018), which leads to a footprint 9 times smaller in area than the AIRS version 6 operational algorithm (Susskind et al., 2003, 2014), mitigating the chance of the impacts of cloud interference on the trace gas retrievals; (2) global infrared land surface emissivity database from the University of Wisconsin–Madison (UOW–M) (Seemann et al., 2007), which improves clear land throughput by 4.5 %; (3) an initial-guess refinement step of cloud fraction prior to the step of joint AIRS+OMI ozone retrievals; (4) a priori constraint vector and matrix identical to the TES version 6 operational algorithm to obtain error estimates consistent with TES data products; (5) an updated a priori and initial-guess information of atmospheric temperature profiles taken from the near-real-time Goddard Earth Observing System Model, Version 5 (GEOS-5) (Rienecker et al., 2008) model data for AIRS TIR temperature profile retrievals; (6) updated a priori ozone built from the Model for Ozone and Related chemical Tracers (MOZART)-4 (Emmons et al., 2010) as offline climatology; (7) High-resolution TRANsmision (HITRAN) 2012 (Rothman et al., 2013) spectroscopic parameters and a priori information of water vapor, the primary interfering species in TIR ozone measurements jointly retrieved with ozone; and (8) labeling the target scenes with a retrieved cloud fraction less than 30 % within the AIRS+OMI field of view as quality-assured, in order to minimize the impacts of cloud interference on ozone data quality. The throughput of AIRS+OMI data processing over the globe is about 30 %.

3.2 Retrieval characteristics of TES, AIRS, OMI, and joint AIRS+OMI

For moderately nonlinear problems, the estimated state can be written as the linear expression (H. M. Worden et al., 2007)

$$\hat{\mathbf{x}} = \mathbf{x}_a + \mathbf{A}[\mathbf{x}_{\text{true}} - \mathbf{x}_a] + \mathbf{G}\epsilon + \delta_{\text{CS}}, \quad (4)$$

where \mathbf{x}_a is the a priori constraint vector; \mathbf{A} is the averaging kernel matrix, whose rows represent the sensitivity of the retrieval to the true state; \mathbf{x}_{true} is the true state vector; ϵ is the spectral noise of satellite instruments; and \mathbf{G} is the gain matrix, which can be written as $\mathbf{G} = (\mathbf{K}^T \mathbf{S}_{\epsilon}^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1} \mathbf{K}^T \mathbf{S}_{\epsilon}^{-1}$. The “cross-state” error, $\delta_{\text{CS}} = \mathbf{A}_{\text{CS}}[\mathbf{x}_{\text{CS}} - \mathbf{x}_{\text{CS}_a \text{ priori}}]$ (H. M. Worden et al., 2007), is incurred from retrieving \mathbf{x}_{CS} , which contains multiple parameters (e.g., water vapor, surface temperature, cloud extinction and cloud top pressure in TIR, cloud fraction in UV, surface albedo, and wavelength-shifting parameters).

Table 1. Spectral regions used in ozone retrievals.

Case selection ^a	Spectral data	Frequency		Resolving power	Atmospheric species
		Start (cm ⁻¹)	End (cm ⁻¹)		
AIRS+OMI, AIRS	AIRS Level 1B version 5 data ^b	985.10	1031.24	1200	H ₂ O, O ₃ , CO ₂
AIRS+OMI, AIRS		1042.76	1048.58		H ₂ O, O ₃ , CO ₂
AIRS+OMI, AIRS		1068.98	1071.38		H ₂ O, O ₃ , CO ₂
AIRS+OMI, AIRS		1108.88	1112.06		H ₂ O, O ₃ , CO ₂
AIRS+OMI, AIRS		1224.10	1227.88		H ₂ O, HDO, O ₃ , CO ₂ , CH ₄ , N ₂ O
AIRS+OMI, AIRS		1259.38	1261.42		H ₂ O, HDO, O ₃ , CO ₂ , CH ₄ , N ₂ O
AIRS+OMI, AIRS		1265.92	1267.06		H ₂ O, HDO, O ₃ , CO ₂ , CH ₄ , N ₂ O
AIRS+OMI, AIRS		1269.46	1270.54		H ₂ O, HDO, O ₃ , CO ₂ , CH ₄ , N ₂ O
AIRS+OMI, AIRS		1311.70	1315.36		H ₂ O, HDO, O ₃ , CO ₂ , CH ₄ , N ₂ O
AIRS+OMI, AIRS		1315.72	1317.82		H ₂ O, HDO, O ₃ , CO ₂ , CH ₄ , N ₂ O
AIRS+OMI, OMI	OMI Level 1B version 3 data ^c	270.00	310.00	460	O ₃
AIRS+OMI, OMI		310.00	330.00	800	

^a The parameters are included in the retrievals for different cases (AIRS only, OMI only, and joint AIRS+OMI). ^b AIRS single-footprint infrared geolocated and calibrated radiance data (Aumann et al., 2003) are used directly rather than Level 2 cloud-cleared spectra, which are calculated using nine adjacent AIRS infrared footprints. Using single-footprint spectra improves the performance of horizontal resolution of the AIRS retrieval from ~ 45 to ~ 13.5 km at nadir, leading to improved representation of horizontal details (Irion et al., 2018). ^c Retrievals' normalized radiances (i.e., $I_{\text{Earthshine}}/I_{\text{solar_irradiance}}$) were used in the retrievals. OMI Level 1B global geolocated earthshine radiance ($I_{\text{Earthshine}}$) and solar irradiances ($I_{\text{solar_irradiance}}$) (Dobber et al., 2006a, b; Van den Oord et al., 2006).

Table 2. List of parameters in state vector.

Case selection ^a	Fitting parameters	Number of parameters	A priori	A priori error
AIRS+OMI, AIRS, OMI	O ₃ at each pressure level	25	MOZART-4 ^b	MOZART-3 ~ 10–40 %
AIRS+OMI, AIRS	H ₂ O at each pressure level	16	GEOS-5 ^c	NCEP ^d ~ 30 %
AIRS+OMI, AIRS	Surface temperature	1	GEOS-5	0.5 K
AIRS+OMI, AIRS	Surface emissivity ^e	23	UOW-M ^f	~ 0.006
AIRS+OMI, AIRS	Cloud extinction ^g	11	Initial BT difference	300 %
AIRS+OMI, AIRS	Cloud top pressure ^g	1	500 mbar	100 %
AIRS+OMI, OMI	UV1 surface albedo	1	OMI climatology ^h	0.05
AIRS+OMI, OMI	UV2 surface albedo (zeroth order) ⁱ	2	OMI climatology	0.05
AIRS+OMI, OMI	UV2 surface albedo (first order) ⁱ	0		0.01
AIRS+OMI, OMI	UV1, UV2 ring scaling factors	2	1.9	1.0
AIRS+OMI, OMI	UV1, UV2 radiance/irradiance wavelength shifts	2	0	0.02 nm
AIRS+OMI, OMI	UV1, UV2 radiance/O ₃ cross-section wavelength shifts	2	0	0.02 nm
AIRS+OMI, OMI	Cloud fraction ^j	1	Derived from 347 nm	0.05

^a The parameters are included in the retrievals for different cases (AIRS only, OMI only, and joint AIRS+OMI). ^b Model for OZone and Related chemical Tracers (MOZART)-4 (Emmons et al., 2010). ^c Goddard Earth Observing System, Version 5 (GEOS-5) (Rienecker et al., 2008). ^d National Center for Environmental Prediction (NCEP) reanalysis (Kalnay et al., 1996). ^e Retrievals over land; spectral surface emissivity is factored in. ^f Global infrared land surface emissivity database at University of Wisconsin-Madison (UOW-M) (Seemann et al., 2007). ^g For cloud treatment in TIR spectral region, we adopt the approach used in the TES Level 2 full-physics retrieval algorithm (Kulawik et al., 2006b; Eldering et al., 2008). Gaussian parameters represent the total optical depth, peak altitude, and profile width. ^h The surface reflectance climatology was constructed using 3 years of OMI measurements obtained between 2004 and 2007 (Kleipool et al., 2008). ⁱ The surface is assumed to be Lambertian with a variable slope in wavelength to the albedo, such that the albedo can vary linearly across the spectral band. ^j For cloud treatment in UV spectral region, we adopt the approach used in the TES+OMI retrieval algorithm (Fu et al., 2013) by adding in an initial-guess refinement step for retrieving the cloud fraction prior to joint AIRS+OMI ozone retrievals.

The use of OE in the MUSES algorithm also provides the averaging kernel and error matrices for each sounding needed for trend analysis, climate model evaluation, and data assimilation. Based on optimal-estimation theory, the averaging kernel matrix (**A**) and total error covariance matrix (**S**) can be calculated as follows:

$$\mathbf{A} = \mathbf{G}\mathbf{K}, \tag{5}$$

$$\mathbf{S} = \underbrace{(\mathbf{I} - \mathbf{A})\mathbf{S}_a(\mathbf{I} - \mathbf{A}^T)}_{\text{smoothing error}} + \underbrace{\mathbf{G}\mathbf{S}_\epsilon\mathbf{G}^T}_{\text{satellite instrument measurement error}} + \underbrace{\mathbf{A}_{cs}\mathbf{S}_{cs}\mathbf{A}_{cs}^T}_{\text{cross-state error}}, \tag{6}$$

satellite instrument observation error

where **I** is the identity matrix, **S_a** is the a priori covariance matrix of the full retrieved state containing both atmospheric

and auxiliary parameters, and S_e is the measurement noise covariance of both TIR and UV radiances. The error variance represented by the diagonal elements in the S_e matrix is computed from the square of spectral noise values obtained from Level 1 data products of AIRS and OMI missions, while the off-diagonal elements are equal to zero. A_{cs} is the submatrix of the averaging kernel for the full-state vector of all jointly retrieved parameters that relates the sensitivity of x (the vector of cross-state parameters) to x_{CS} . The diagonal elements of S_{cs} contain the a priori covariance for the other jointly retrieved parameters, including water vapor, surface temperature, surface emissivity, cloud parameters in infrared (extinction and cloud top pressure), surface albedo in UV, wavelength shifting in UV, and cloud parameter in UV (cloud fraction) parameters, while the off-diagonal elements are equal to zero. It is worth noting that the retrieval scheme does not include the radiative transfer model error, which is negligible since (1) both the ELANOR for the TIR and Vector Linearized Discrete Ordinate Radiative Transfer (VLIDORT) for the UV (Spurr, 2006, 2008) are full-physics radiative transfer models that have high accuracy and (2) the comparisons of satellite–ozonesonde presented in Sect. 4.2 show that agreement of the collocated ozonesonde–satellite measurements is within the expected ranges.

The trace of the averaging kernel matrix (A) gives the number of independent pieces of information in the vertical profile, or the degrees of freedom for signal (DOFS) (Rodgers, 2000). A larger DOFS value indicates a better vertical sensitivity. Figure 1 shows sample averaging kernel matrices for TES, AIRS, OMI, and joint AIRS+OMI transect observations over the western United States on 23 August 2006. The joint AIRS+OMI and TES retrievals show similar capability for resolving the lower/upper troposphere (tropospheric DOFS: 1.64 for TES, 1.55 for joint AIRS+OMI). Both AIRS and OMI tropospheric DOFS are ~ 1 – capable of estimating the tropospheric columns but lacking vertical sensitivity in the troposphere.

4 Validation of joint AIRS+OMI data

An initial comparison between TES, AIRS, OMI, and AIRS+OMI is shown by a transect from $\sim 6^\circ$ N to 55° N taken on 23 August 2006 (Fig. 2a) and processed through the MUSES algorithm. The tropospheric ozone concentration profiles of joint AIRS+OMI retrievals show better agreement with TES data (Fig. 2g, green curve; mean differences $< 2\%$ from surface to 400 hPa, and $< 5\%$ from 400 to 100 hPa) than the retrievals for both AIRS and OMI alone (Fig. 2g, blue curve for AIRS, purple curve for OMI). The joint retrievals improve the agreement due to the increased vertical sensitivity in comparison to each instrument alone since the multispectral retrievals have the advantage of obtaining the vertical distribution information of atmospheric composition from multiple physical regimes, including the

atmospheric thermal emissions, pressure- and temperature-dependent spectral line broadenings and absorption cross sections via both TIR and UV radiances, and wavelength- and altitude-dependent atmospheric scattering events via UV radiances.

Further evaluation of the joint AIRS+OMI O_3 retrievals is shown in two modes: global survey (GS) and regional mapping (RE). The GS mode provides profile data at nadir position along the satellite ground track, i.e., a spatiotemporal sampling identical to TES GS, while RE mode processes all available AIRS+OMI measurements over a region; specifically in this case we have considered the Korean Peninsula during the 2016 Korea–United States Air Quality (KORUS-AQ) campaign (Miyazaki et al., 2018). The global joint AIRS+OMI retrievals have been compared to the well-validated TES data (Sect. 4.1) and high accuracy in situ global ozonesonde measurements (Sect. 4.2) to quantify the performance of this multispectral tropospheric ozone profile data product. These comparisons were made using measurements in 2006, when neither the TES instrument degradation nor OMI row anomaly played a role.

4.1 Comparison to the TES data

Joint AIRS+OMI ozone retrievals apply only to daytime scenes, since OMI measurements depend on the sunlight, though the MUSES algorithm processes both daytime and nighttime TIR space measurements. The “species retrieval quality” flag of joint AIRS+OMI data, a master quality flag available in the Level 2 product files, was determined by evaluating a suite of retrieval characteristics including the spectral fitting residuals, cloud fraction within field of view (when effective cloud fraction in OMI $> 30\%$), and the lapse rate of tropospheric ozone vertical distribution. The retrieval scheme processes the AIRS+OMI measurements over all sky conditions, though only the scenes of the cloud fraction within field of view less than 30% were flagged as quality-assured. The retrieval acceptance rate of joint AIRS+OMI ozone in 2006 is about 30% .

Both TES and joint AIRS+OMI 2006 ozone profile data were screened prior to the comparison using (1) the species retrieval quality, (2) the retrieved cloud effective TIR optical depth (removed when $OD > 2.0$), and (3) solar zenith angle (SZA; excluded when $SZA > 80^\circ$, i.e., daytime only). We excluded profiles with thick clouds in the field of view because these obscure the infrared emission from the lower troposphere, which greatly reduces the satellite sensitivity of both TIR and UV radiances. For cloud treatment, we adopt the approach used in the joint TES+OMI retrieval algorithm (Fu et al., 2013) by adding in an initial-guess refinement step for retrieving the cloud fraction within OMI field of view, prior to joint AIRS+OMI ozone retrievals. The impacts of cloud and surface properties have been taken into account in the retrievals, since the MUSES algorithm simultaneously retrieves both the trace gases profiles and the cloud/surface

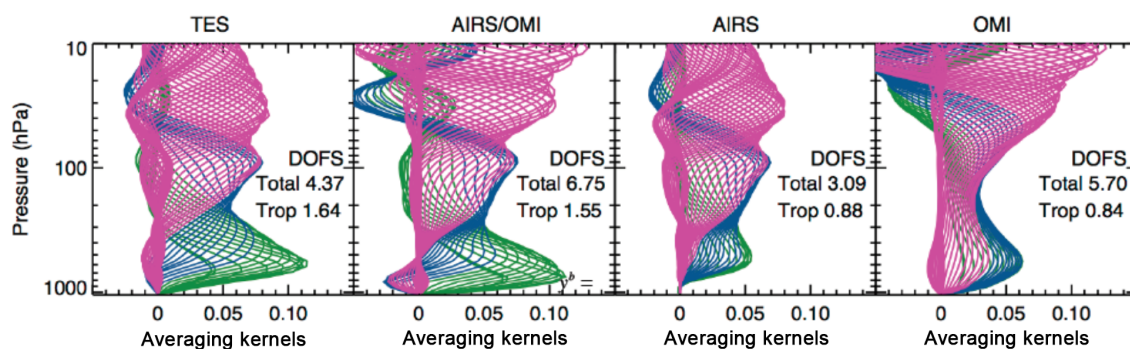


Figure 1. Averaging kernels of collocated measurements of TES (version 6), joint AIRS+OMI, AIRS alone, and OMI alone over California, USA, on 23 August 2006. The green, blue, and magenta curves in four panels indicate the averaging kernels in the pressure range of surface–400 hPa, 400–100 hPa, and above 100 hPa, respectively.

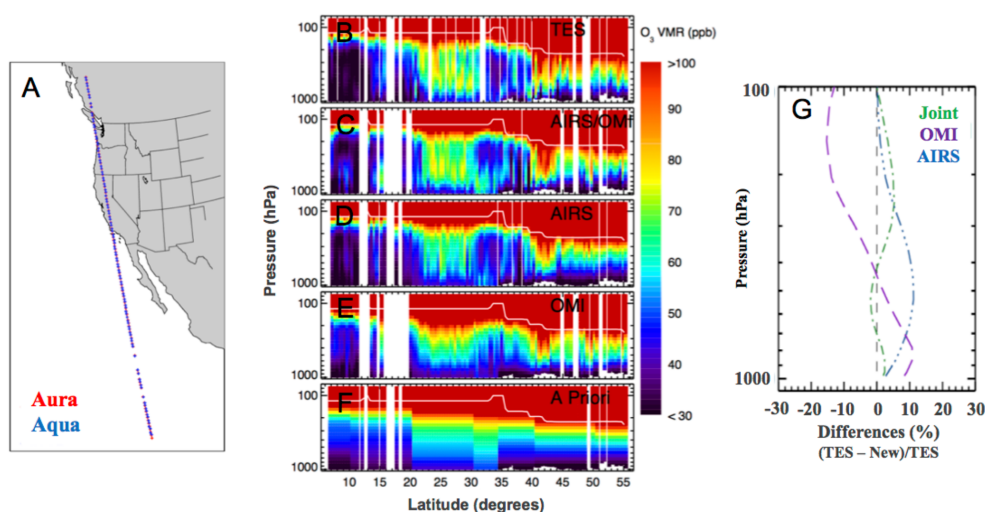


Figure 2. Collocated ozone (O_3) measurements from A-Train nadir-viewing spectrometers over the western United States on 23 August 2006. (A) Geolocation of 110 TES–AIRS–OMI triads (spatiotemporal differences ~ 8 km, ~ 16 min); (B) vertical profile of TES O_3 volume mixing ratio (VMR) data (version 6) in units of parts per billion (ppb); (C) joint AIRS+OMI retrievals; (D) AIRS alone; (E) OMI alone; (F) a priori used in retrievals; and (G) averaged percentage differences of retrieved O_3 profiles in comparison to TES O_3 data (version 6): TES vs. joint AIRS+OMI (green dash-dotted line), TES vs. AIRS alone (blue), and TES vs. OMI alone (purple dashed line). The white curves in the panels of (B–F) indicate the tropopause pressure taken from the Goddard Earth Observing System Model, Version 5.

parameters. The retrieved values and estimated errors of the TIR cloud effective optical depth and cloud height, UV cloud fraction within the field of view, and cloud top height are provided in the joint AIRS+OMI data product files.

Joint AIRS+OMI global tropospheric O_3 retrievals (Fig. 3A1–3, August 2006 monthly mean data) show good agreement with TES data, as shown in Fig. 3B1–3. Both data sets are significantly different from the a priori and capture the synoptic ozone patterns such as the midlatitude Atlantic and the biomass burning events (e.g., southern Africa). Results for the remaining months of 2006 are available in Figs. S1–S11 in the Supplement. The correlation coefficients of joint AIRS+OMI and TES version 6 data (Table 3) are greater than 0.71 and up to 0.92 for all months across the troposphere, where the mean and root mean square (rms) of the

differences of two data sets (Table 3) are well within the estimated total error. The period of September–November coincides with the slight drop of the Pearson correlation coefficient values. For September 2006 data, the different spatiotemporal sampling between TES and joint AIRS+OMI data is the reason for the slight drop. In September 2006, TES and joint AIRS+OMI data deliver nine and 15 global surveys, respectively (bottom row of Table 3). TES did not deliver measurements from 1 to 9 September. In support of the TEXAQS II flight campaign, TES delivered additional special observations by reducing the number of global surveys in the end of September. For October and November 2006 data, the slight drop of the correlation coefficients might relate to the slight difference of measurement sensitivity between TES

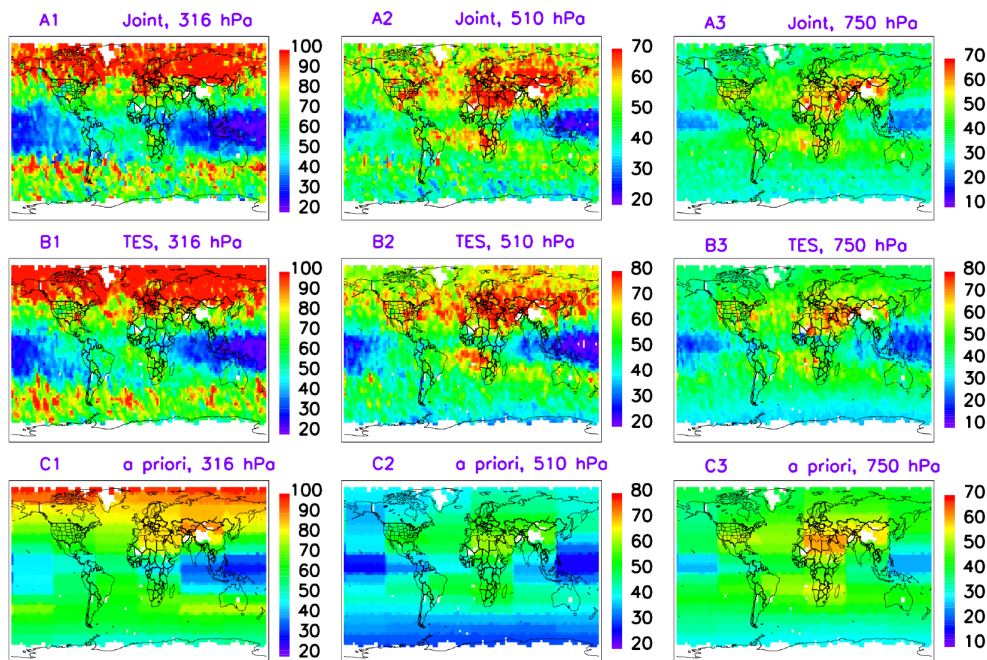


Figure 3. Global maps of monthly averaged ozone (O_3) volume mixing ratio (VMR) in units of ppb. The A-Train measurements in August 2006 were used in creating these global maps. Comparison of joint AIRS+OMI (A), TES (B), and a priori (C) ozone VMR for the pressure level of 316, 510, and 750 hPa (columns left, middle, right), respectively. All data have been gridded to $2.5^\circ \times 2.5^\circ$ cells. Results for the remaining months of 2006 are available in Figs. S1–S11 in the Supplement.

and joint AIRS+OMI, as shown in Figs. S20 and S21 in the Supplement.

The characteristics of the joint AIRS+OMI retrievals, in terms of vertical sensitivity and estimated error characteristics, are similar to those of TES data. The DOFS, which quantify the vertical sensitivity of global tropospheric ozone retrievals, show distributions similar to TES data (Fig. 4 panels A2 and B2 for August 2006). Figures S12–S22 in the Supplement present the DOFS for the remaining months of 2006. Both the estimated observation and total errors of joint AIRS+OMI retrievals (black curves of Fig. 5) show peaks and widths equivalent to those of TES data products (green curves of Fig. 5) across troposphere over the globe. Figures S23–S33 in the Supplement present the estimated errors for the remaining months of 2006. The peak of the estimated observation errors, which are the sum of second and third terms in Eq. (6), resides in the range of 6 %–8 % (or ~ 3 ppb) for the joint AIRS+OMI retrievals – equivalent to the observation error of 5 %–7 % (or ~ 2 –3 ppb) from TES data across the troposphere. Finally, the joint AIRS+OMI retrievals have total errors within 3 % agreement over the globe – equivalent to TES data.

4.2 Comparison to ozonesonde measurements

We identified 424 sonde–AIRS–OMI triads and 556 sonde–TES pairs following the coincidence criteria in Sect. 2. Following H. M. Worden et al. (2007), satellite observation oper-

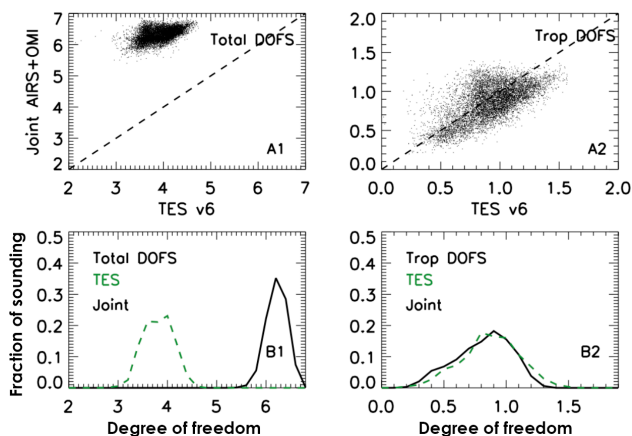


Figure 4. DOFS for O_3 over globe shown in Fig. 3. Here, we used the A-Train measurements from August 2006. Results for the remaining months of 2006 are available in Figs. S12–S22 in the Supplement. (A1) Total DOFS; (A2) tropospheric DOFS; (B1) histogram of total DOFS: joint AIRS+OMI (black line) and TES version 6 (green dashed line); and (B2) histogram of tropospheric DOFS joint AIRS+OMI (black line) and TES version 6 (green dashed line).

ators $\mathbf{H}(x_a, \mathbf{A})$ defined in the equation for joint AIRS+OMI and TES were applied to the in situ ozonesonde profiles accounting for known bias and precision. As a result, the expected covariance matrix of the differences between the

Table 3. Comparisons between joint AIRS+OMI and TES gridded (2.5° × 2.5°) global survey measurements of ozone concentration at three pressure levels (316, 510, and 750 hPa) for the year 2006.

316 hPa		Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.
Pearson correlation coefficient		0.83	0.84	0.85	0.84	0.84	0.84	0.84	0.82	0.74	0.74	0.71	0.78
Differences (TES–AIRS+OMI)	Mean (ppb)	8.3	8.5	7.3	6.9	8.1	6.0	4.8	2.8	1.4	1.9	3.4	5.4
	rms (ppb)	16.3	20.5	21.5	21.6	22.6	19.8	17.8	15.6	16.2	14.9	13.4	13.2
	Mean (%)	12.9	11.6	9.8	7.3	7.3	5.0	3.6	2.3	0.1	1.9	4.3	7.7
	rms (%)	24.1	26.5	24.2	25.7	24.7	23.8	22.1	22.2	23.2	23.7	21.2	20.8
Total error	AIRS+OMI O ₃ (%)	28.8	28.8	28.6	28.9	28.5	28.0	27.9	27.5	27.8	28.2	28.8	28.9
	TES V6 O ₃ (%)	22.7	22.6	22.5	23.0	22.9	22.1	22.1	22.2	22.5	22.3	22.8	22.9
510 hPa		Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.
Pearson correlation coefficient		0.81	0.84	0.87	0.88	0.89	0.86	0.86	0.82	0.74	0.74	0.71	0.79
Differences (TES–AIRS+OMI)	Mean (ppb)	3.3	2.6	2.9	3.3	3.6	4.1	4.1	4.2	3.9	3.2	3.1	3.5
	rms (ppb)	7.7	8.3	8.6	8.9	9.2	9.5	8.7	8.2	8.7	8.2	7.4	7.1
	Mean (%)	6.5	3.8	4.9	4.2	4.2	4.5	4.6	5.6	4.7	4.2	5.2	6.6
	rms (%)	16.3	18.2	17.3	18.2	16.4	17.0	16.5	15.0	16.5	16.6	15.4	15.4
Total error	AIRS+OMI O ₃ (%)	22.5	22.4	22.5	22.8	23.0	22.8	22.7	22.6	22.4	22.3	22.4	22.5
	TES V6 O ₃ (%)	20.4	20.4	20.1	20.1	20.1	19.5	19.6	19.7	19.6	19.4	19.9	20.3
750 hPa		Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sep.	Oct.	Nov.	Dec.
Pearson correlation coefficient		0.92	0.89	0.90	0.90	0.90	0.83	0.82	0.80	0.74	0.76	0.87	0.92
Differences (TES–AIRS+OMI)	Mean (ppb)	0.4	−0.8	0.4	1.2	1.6	2.2	2.3	3.2	3.3	2.0	1.5	1.4
	rms (ppb)	5.2	6.4	6.7	7.0	6.4	7.1	6.4	6.3	6.9	6.1	5.2	4.8
	Mean (%)	0.2	−4.1	0.6	0.3	1.3	2.3	2.9	5.4	5.4	2.9	2.9	3.0
	rms (%)	14.7	19.3	19.3	19.8	15.9	17.4	16.0	14.6	16.8	16.1	14.3	13.4
Total error	AIRS+OMI O ₃ (%)	23.9	22.9	22.4	22.9	24.1	24.7	24.6	24.2	23.9	24.2	23.9	24.1
	TES V6 (%)	24.2	23.8	23.1	23.3	24.0	24.0	24.0	23.6	23.4	23.4	23.5	23.8
Number of global survey	AIRS+OMI	16	13	14	15	16	15	15	16	16	15	15	16
	TES	15	13	16	14	15	15	16	14	9	16	15	15

satellite retrievals and ozonesonde measurements smoothed by instrument averaging kernels can be written similarly to Eq. (6) (H. M. Worden et al., 2007; Fu et al., 2013):

$$\begin{aligned}
 E \left[(\hat{x} - \hat{x}_{\text{sonde}}) (\hat{x} - \hat{x}_{\text{sonde}})^T \right] &= \underbrace{\mathbf{A} \mathbf{S}_{\text{sonde}} \mathbf{A}^T}_{\text{ozonesonde measurement error}} \\
 &+ \underbrace{\mathbf{G} \mathbf{S}_{\varepsilon} \mathbf{G}^T}_{\text{satellite instrument measurement error}} + \underbrace{\mathbf{A}_{\text{cs}} \mathbf{S}_{\text{cs}} \mathbf{A}_{\text{cs}}^T}_{\text{cross-state error}} + \underbrace{\mathbf{G} \mathbf{S}_{\varepsilon r} \mathbf{G}^T}_{\text{remaining radiance calibration error}} \\
 &+ \underbrace{\mathbf{S}_{\text{SS}}}_{\text{sonde-satellite temporal spatial sampling}}. \tag{7}
 \end{aligned}$$

Equation (7) indicates that the error covariance matrix is not biased by the a priori \mathbf{x}_a , and the biases of O₃ retrievals relative to ozonesondes are due to the errors of the sonde measurements \mathbf{S}_s , the random spectral noise \mathbf{S}_{ε} , the interfering parameters in retrieval state vector \mathbf{S}_{cs} , the remaining radiometric calibration errors $\mathbf{S}_{\varepsilon r}$, or sonde–satellite spatiotemporal samplings \mathbf{S}_{SS} .

Figure 6 and Table 4 illustrate that both joint AIRS+OMI and TES data are in good agreement with ozonesonde measurements across seasonal variations in the troposphere.

Here, the biases of ozone from remote-sensing measurements are within 3, 2, and 5 ppb for joint AIRS+OMI at three pressure levels (316, 510, and 750 hPa, respectively) and within 6, 4, and 3 ppb, respectively, for TES version 6 data. The biases of these satellite data show an improvement for all seasons when compared to a high bias of 3–10 ppb estimated for the TES tropospheric ozone data prior to version 6 via validation using ozonesonde measurements (Nassar et al., 2008; Boxe et al., 2010). Additionally, the rms’s of the differences are 10–17, 8–11, and 7–9 ppb for the tropospheric ozone of joint AIRS+OMI retrievals and 12–22, 8–15, and 7–13 ppb for TES version 6 data, consistent with those reported by the existing TES validations. Overall comparisons of AIRS+OMI to ozonesondes (with observation operator applied to account for sensitivity) yield similar biases and errors to matching comparisons between TES and sondes. Note that Fig. 6 and Table 4 report that single band retrievals (AIRS-alone and OMI-alone data) have larger bias in comparison to the joint AIRS+OMI data. Table 5 shows comparisons to the original ozonesonde measurements (i.e., without satellite observation operator applied). These direct comparisons are often used for comparing instruments of differing sensitivities, because more sensitive instruments are expected to show better agreement to the ozonesondes. The

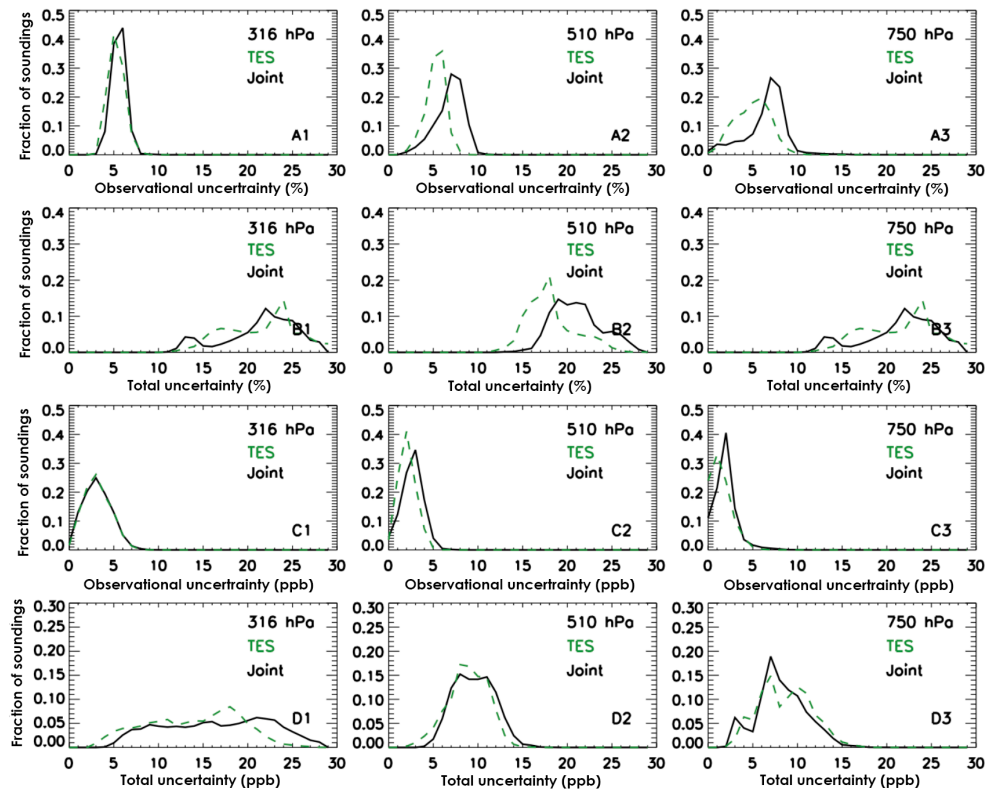


Figure 5. Estimated (predicted) error of retrieved global O₃ concentration shown in Fig. 3. Here, we used the A-Train measurements from August 2006. Results for the remaining months of 2006 are available in Figs. S23–S33 in the Supplement. (A1–A3) Observational error; (B1–B3) total error; (C1–C3) observational error in ppb; and (D1–D3) total error in ppb. Joint AIRS+OMI data are shown as a black line, and TES version 6 data are shown as a green dashed line.

joint AIRS+OMI performs best, as seen in the reduction of measurement bias at three pressure levels and improved rms at the 750 hPa level.

5 Conclusions

We have shown multispectral retrievals using both AIRS TIR and OMI UV measured radiances for tropospheric O₃ profiling. This technique enables the continuation of the TES capability to distinguish between upper- and lower-tropospheric ozone abundances. The global-scale comparisons between joint AIRS+OMI (version 1) and TES (version 6) O₃ profile products across four seasons in the troposphere on a global scale show that these two data products are comparable for a wide variety of geophysical conditions: correlation coefficients are 0.7–0.9 at three pressure levels (316, 510, and 750 hPa), and both the mean (0.8–4.2 ppb) and rms differences (± 4.8 –23 ppb) are within the estimated total errors. The joint TIR+UV retrieval provides equivalent vertical sensitivity and error characteristics of high-spectral-resolution TES measurements, which have a spectral resolution that is ~ 8 –12 times higher than AIRS and OMI measurements, though about 3-times-lower SNR. Comparisons

of collocated joint AIRS+OMI, TES, and ozonesonde measurements show that both mean and standard deviation of the differences are within the estimated measurement error of these space sensors. The joint AIRS+OMI ozone products have a high bias of 2–5 ppb similar to TES data (3–6 ppb). Consequently, the similarities of the retrieved concentration, vertical sensitivity, and error characteristics between joint AIRS+OMI and TES ozone data demonstrate that combining the measurements of the existing TIR and UV hyperspectral imaging spectrometers can extend the well-validated NASA EOS high-spectral-resolution TES tropospheric ozone profile data products.

Both AIRS and OMI have wide swath widths (AIRS: 1650 km; OMI: 2600 km) across satellites' ground tracks; consequently, the joint AIRS+OMI measurements promise to extend and even improve the number of available observations by over 100 times that of TES. The product files of the joint AIRS+OMI 2006 ozone global survey retrievals, including a validation report and a reader program, are available via the Aura Validation Data Center (AVDC) website (https://avdc.gsfc.nasa.gov/pub/data/satellite/Aura/TES/AIRS_OMI-version0.1Beta/, last access: 4 October 2018). The GS and RE modes of joint AIRS+OMI data from March

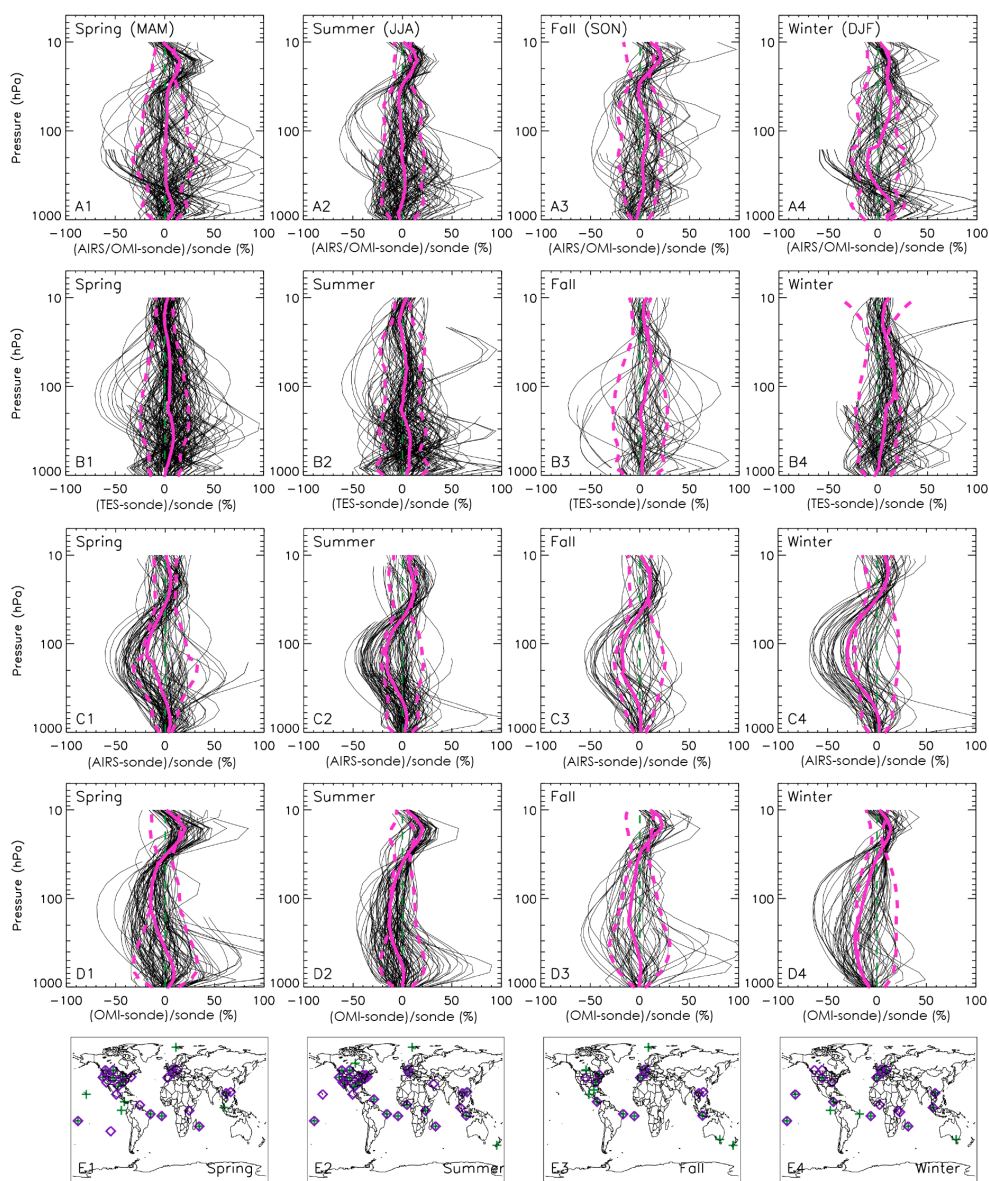


Figure 6. Joint AIRS+OMI–sonde (**A1–A4**), TES–sonde (**B1–B4**), AIRS–sonde (**C1–C4**), and OMI–sonde (**D1–D4**) percentage differences of measured ozone concentration for the four seasons (months abbreviated in parentheses) on a global scale. Individual profiles are shown in black, and the mean and 1σ standard deviation range are overlaid in solid magenta (mean) and as dashed magenta lines. The profiles were plotted after removing cloudy scenes and flagged satellite (joint AIRS+OMI and TES) data. (**A1–A4**) Joint AIRS+OMI vs. ozonesonde; (**B1–B4**) TES data (version 6) vs. ozonesonde; (**C1–C4**) WOUDC sonde location that have coincident measurements with joint AIRS+OMI (green plus signs) and TES (purple diamonds).

to June 2016 in support of KORUS-AQ are also available on the same website. These results were also applied in the post-flight data analysis by Miyazaki et al. (2018) that showed great error reductions in the tropospheric ozone analysis, especially in the middle troposphere, through assimilation of joint AIRS+OMI data. Overall comparisons of AIRS+OMI to ozonesondes and aircraft for the year 2016 yield similar biases and errors to matching comparisons for the year 2006. Using the MUSES algorithm, the AIRS+OMI global survey

mode data (2004 to present) with a footprint size of about 15 by 24 km is being processed using the facilities within the JPL TES Science Investigator-led Processing (SIP) system to build up a decadal record of tropospheric ozone products.

The current spatial coverage of AIRS+OMI is sufficient to extend the TES ozone record beyond 2010, when TES ceased the global survey mode measurements. The combined AIRS+OMI product can provide a record of tropospheric and total ozone spanning the full Aura satellite time peri-

Table 4. Comparisons between satellite remote-sensing and ozonesonde in situ measurements for 2006 at three pressure levels (316, 510, and 750 hPa), with satellite observation operators applied to the ozone measurements in order to account for sensitivity.

	316 hPa		Spring		Summer		Fall		Winter			
			AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES		
Mean (ppb)	2.8/	-2.2/-1.0	6.1	0.7/	-4.1/-6.1	4.2	1.1/	-5.9/-2.2	-1.6	-2.5/	-9.2/-11.9	2.9
Mean (%)	1.3/	-4.1/-2.6	8.6	2.2/	-7.0/-9.6	6.6	2.9/	-9.6/-4.4	3.3	-7.7/	-15.8/-22.0	6.5
rms (ppb)	17.1/	17.8/21.6	19.2	13.4/	10.7/11.6	17.0	12.6/	16.3/24.5	21.7	10.0/	10.7/12.2	12.4
rms (%)	25.6/	24.0/26.7	23.7	20.4/	16.5/17.4	23.8	19.0/	21.7/29.5	26.9	20.8/	18.8/19.5	20.5
	510 hPa		Spring		Summer		Fall		Winter			
			AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES		
Mean (ppb)	1.3/	1.1/3.8	3.6	-0.8/	1.2/-0.8	3.5	0.4/	-0.4/0.4	0.2	1.8/	-1.7/-9.2	1.4
Mean (%)	3.8/	2.3/5.8	7.0	1.6/	2.0/-1.7	7.3	2.5/	-1.0/0.3	3.5	5.9/	-2.8/-18.9	3.2
rms (ppb)	7.6/	7.6/18.7	9.2	10.9/	7.8/12.2	10.6	8.6/	9.4/16.1	14.5	7.5/	7.8/9.7	8.0
rms (%)	17.2/	14.4/32.0	17.4	20.4/	14.6/22.3	17.9	16.7/	17.1/28.6	21.8	19.1/	15.7/18.5	17.7
	750 hPa		Spring		Summer		Fall		Winter			
			AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES	AIRS+OMI/ AIRS/OMI	TES		
Mean (ppb)	2.4/	2.1/4.5	1.7	-2.2/	2.0/0.6	2.6	-1.2/	0.3/0.1	0.3	4.6/	0.5/-6.1	0.3
Mean (%)	8.0/	5.0/8.1	3.4	-2.0/	4.3/1.1	6.6	-1.3/	0.9/0.2	1.9	14.4/	1.7/-14.0	0.9
rms (ppb)	7.6/	5.3/14.7	6.9	8.6/	7.3/8.6	12.5	6.3/	6.0/10.1	11.2	8.5/	6.1/7.0	7.8
rms (%)	21.1/	11.6/28.3	16.2	18.8/	16.4/18.9	25.3	13.2/	12.5/21.0	23.9	24.8/	14.9/14.8	20.0
Number of WOUDC sonde sites	20		25	27	30	16		12	16		19	
Number of satellite/ sonde coincidences	131		197	134	171	72		60	87		128	

Table 5. Comparisons between satellite remote-sensing and ozonesonde in situ measurements for 2006 at three pressure levels (316, 510, and 750 hPa), without the satellite observation operators applied to the ozonesonde measurements.

	Mean						rms					
	ppb			%			ppb			%		
	316 (hPa)	510 (hPa)	750 (hPa)	316 (hPa)	510 (hPa)	750 (hPa)	316 (hPa)	510 (hPa)	750 (hPa)	316 (hPa)	510 (hPa)	750 (hPa)
AIRS+OMI	-2.8	-0.5	0.3	-7.0	-1.0	1.1	20.6	8.9	5.8	29.9	16.2	12.4
AIRS	-4.9	0.3	1.5	-8.4	0.6	3.5	14.0	7.9	6.2	20.7	15.1	14.0
OMI	-5.2	-1.0	0.4	-9.3	-2.8	-0.1	17.5	15.2	11.2	24.2	27.3	22.8
TES	1.8	1.7	0.9	1.7	3.2	4.4	27.5	15.4	16.2	36.5	27.9	37.5

ods (2005–current). However, the daily global coverage of OMI measurements has been decreasing since 2009 due to the OMI row anomaly (Schenkeveld et al., 2017; Huang et al., 2017; Levelt et al., 2018). Looking to the future and as a way to further increase science return, we have investigated the feasibility of constructing an additional multiple-decade-long tropospheric ozone profile data set using a MUSES-based multispectral approach that combines the radiance measured by the CrIS and Ozone Mapping Profiler Suite (OMPS) instruments. This additional data set has the potential to fill the spatial gaps in the joint AIRS+OMI data record since 2012. Both the CrIS and OMPS instruments are on the Suomi National Polar-orbiting Partnership (NPP) satel-

lite, which launched in 28 October 2011. The spectral characteristics of the CrIS instrument (Han et al., 2013; Strow et al., 2013) are similar to the AIRS instrument, and those for OMPS (Flynn et al., 2006, 2014; Kramarova et al., 2014; Pan et al., 2017) are similar to the OMI instrument. Hence, as expected, joint CrIS+OMPS retrievals present characteristics (Fig. 7) similar to the joint AIRS+OMI retrievals (Fu et al., 2017).

It is worth noting that the second set of CrIS and OMPS instruments on board the Joint Polar Satellite System-1 (JPSS-1, also known as NOAA-20) satellite were successfully launched to space on 18 November 2017. The JPSS-2 (also known as NOAA-21) satellite, which is the platform of the

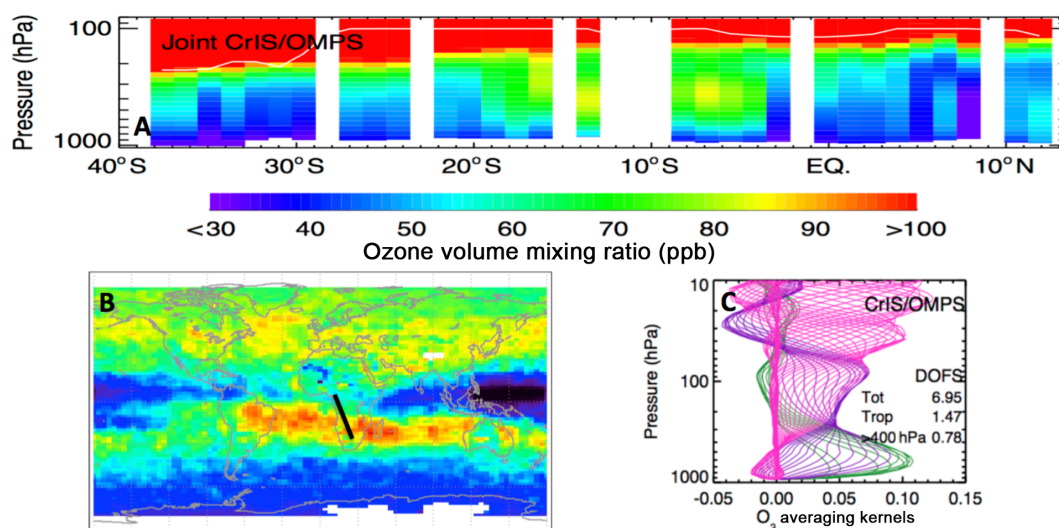


Figure 7. Joint CrIS+OMPS ozone profile retrievals over Africa on 21 October 2013. The elevated ozone concentrations between 2 and 20° S are associated with biomass burning. (A) The retrieved ozone concentration profiles along the transect measurements. The white curve indicates the tropopause pressure reported by GEOS-5. (B) TES monthly mean ozone concentration at 510 hPa. The black line indicates the joint CrIS+OMPS measurement location. (C) The averaging kernels of joint CrIS+OMPS measurements.

third set of CrIS and OMPS instruments, is scheduled to launch in 2022. The NOAA-20/JPSS-1 OMPS Nadir Mapper products' resolution has improved from $50 \times 50 \text{ km}^2$ field of view to $12 \times 17 \text{ km}^2$ (JPSS-1) and will further improve to $10 \times 10 \text{ km}^2$ (JPSS-2) in the operational NOAA processing (Lawrence E. Flynn, personal communication, 2018). The NASA Goddard Space Flight Center (GSFC) Level 1 products of the JPSS-1 OMPS Nadir Mapper will have a spatial resolution of $10 \times 10 \text{ km}^2$ to help detect sources of sulfur dioxide, including volcanoes and coal-burning power plants (press release via <https://spacenews.com/> by Glen Jaross, last access: 4 October 2018). As a result, the joint CrIS+OMPS retrievals, with characteristics similar to AIRS+OMI retrievals but with improved spatial coverage, illustrate the potentials of extending the tropospheric ozone profile data record to the next decades using the measurements from the Suomi-NPP, JPSS-1, and JPSS-2 satellites. The TROPOMI instrument (Veefkind et al., 2012) on board the sentinel-5 Precursor (S5P) satellite was successfully deployed into its orbit on 13 October 2017 and formed a new satellite constellation with Suomi-NPP, currently 5 min apart and with the plan of reducing to 3 min time difference in the future. The spatial resolution of TROPOMI is an unprecedented 3.5×7.0 and $7.0 \times 7.0 \text{ km}^2$ in the UV-VIS and shortwave IR (SWIR) spectral bands, respectively, providing another opportunity to obtain the high-resolution tropospheric ozone ESDR via the multispectral retrieval technique, which combines CrIS and TROPOMI measurements.

Data availability. The joint AIRS+OMI ozone data and Woudc sonde data used in the data analysis can be freely downloaded from the websites of AVDC (https://avdc.gsfc.nasa.gov/pub/data/satellite/Aura/TES/AIRS_OMI-version0.1Beta/ last access: 4 October 2018) and Woudc (<http://www.woudc.org>, last access: 4 October 2018) accordingly.

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Author contributions. DF and SK developed the joint AIRS+OMI retrieval algorithm; KM, KB, JW, AE, and NL helped in the estimation of joint AIRS+OMI measurement uncertainty; KM, RH, GO, AT, and ML helped in the validation and quality flagging of joint AIRS+OMI data products; XL and PL shared knowledge of the OMI Level 1B data and helped in the UV radiative transfer modeling; and JT and FI shared knowledge of the AIRS Level 1B data. All authors participated in writing the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

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