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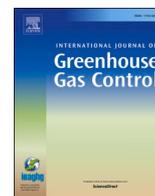
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Ultrasonic response of a brine-saturated reservoir rock during coupled stress and fluid perturbation during liquid-CO₂ injection

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ABSTRACT

CO₂ injection into porous sandstone reservoirs offers a promising pathway to curb anthropogenic carbon emissions, but poses risks of leakage and induced seismicity from stress perturbations and fault reactivation without meticulous monitoring. Here, we present a time-lapse monitoring approach based on laboratory measurements of ultrasonic V_p, V_s and corresponding peak amplitudes in critically stressed, partially saturated North Sea sandstones (porosity 9–23%). Our experiments show that V_p and V_s exhibit higher sensitivity (4–15%) to stress changes compared to fluid saturation changes (0.8–1%), whereas amplitudes are more responsive (30–500%) to saturation, showing staggered change when brine is displaced by CO₂. Under pure stress perturbation, amplitude variations are smaller (10–50%). During elastic deformation, the V_p/V_s ratio decreases while the ratio of their corresponding amplitudes increases, underscoring the need for both P- and S-wave measurements. Velocity and amplitude changes are more pronounced in high-porosity rocks. In a critically stressed state (beyond yield/before failure), the rise in pore fluid density from CO₂ injection boosts shear wave amplitudes, offsetting attenuation from inelastic deformation. Knowing the pre-injection stress state enables these velocity and amplitude trends to serve as robust indicators of reservoir conditions during and after CO₂ injection. This cost-effective approach can be adapted to reservoir-scale monitoring and extends beyond CCS, supporting enhanced detection of stress and fluid-induced changes in subsurface formations.

1. Introduction

Subsurface CO₂ storage (CCS) is gaining accelerated growth worldwide as governments and industry solidify their commitment to atmospheric CO₂ abatement (Bui et al., 2018; Y. Zhang et al., 2022). Depleted oil and gas reservoirs are relatively well studied due to their history of hydrocarbon production and are therefore identified as primary targets for storing CO₂. However, CO₂ injection operation is complicated and comes with its own risks including possibility of leakage, wayward plume movement, risk of fault reactivation etc. to name a few. Therefore, extensive studies are performed both in laboratory and reservoir scale prior to injection to ensure safe containment of CO₂ for a considerably long period. The short term and long-term effects of CO₂ on mechanical properties of sandstones along with its velocity signature have been explored in recent years (Cerasi et al., 2018; Hangx et al., 2013; Hu et al., 2017; Kim and Makhnenko, 2022; Ranjith and Perera, 2011; Tarokh et al., 2020; Van Stappen et al., 2018; Xing et al., 2019;

Zhang et al., 2021; Zheng et al., 2015), which report permanent alteration in mineralogy, porosity, permeability and fracture aperture (for fractured rocks) depending on the experimental timescale, pressure, temperature and phase of CO₂ (gas, liquid or supercritical). Monitoring tools used during and post injection also ensure if the CO₂ plume and reservoir both are behaving in an expected manner (Feng, 2017; Furre et al., 2017; Will et al., 2014).

Evolution of stress-state is a key parameter to monitor during a CO₂ injection operation, which is directly affected by continuous build-up of CO₂ pressure in the reservoir. In the subsurface direct measurement of stress is tricky and often other proxies are used to infer stress state in the reservoir such as InSAR data, strainmeter, tiltmeter etc., which provide regional or local reservoir or ground deformation (Furst et al., 2020; Murdoch et al., 2020; T. Zhang et al., 2022). Geophysical monitoring using seismic interferometry and 3D seismic imaging, using conventional tools and optic fibers are periodically performed to track CO₂ plume movement in the subsurface (Bachrach et al., 2023; Feng, 2017;

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Furre et al., 2017; Gasperikova et al., 2022), while microseismicity monitoring is performed to track any indications of induced seismic activity (Will et al., 2014). However, these monitoring techniques are expensive, which leads to an increased project expense. Compressional and shear wave velocity measurement have proven to be sensitive to both stress change (Barnhoorn et al., 2018; Chandra et al., 2024; Dutta et al., 2020; Sang et al., 2020; Sayers and Kachanov, 1995; Veltmeijer et al., 2024; Zotz-Wilson et al., 2019) and CO₂ plume movement in the subsurface (Agofack et al., 2018; Falcon-Suarez et al., 2020; Fortin et al., 2005; Guéguen and Schubnel, 2003; Hamada and Joseph, 2020), which can be leveraged to effectively monitor safe disposal of CO₂ into the subsurface. Lab scale experiments on the effect of fluid saturation and stress change on ultrasonic and seismic velocity have been explored before, which mostly govern the relationship between degree of saturation and velocity, but limited studies have been conducted to explore the coupled impact of stress and saturation change in porous reservoirs (Hughes and Kelly, 1952; Janssen et al., 2021; King and Marsden, 2012; Li et al., 2022; Meglis et al., 2005; Scott et al., 1993; Stanchits et al., 2003; Yamabe et al., 2016; Zhu et al., 2022; Zhubayev et al., 2016).

Irreversible deformation of reservoir rocks, which is detrimental for safe operation, cause reduction of V_p and V_s. However, such velocity reduction can also be caused by displacement of reservoir brine by CO₂ due to the difference in density of both fluids (Gassman, 1951; Nur and Simmons, 1969), which is an integral part of CO₂ injection. Therefore, without proper benchmarking, change in wave attributes can raise false alarm during CO₂ injection. Previous studies so far have illustrated the effect of CO₂ saturation and plume movement on wave propagation both from field scale and laboratory observations. Alemu et al., (2013) found that at low CO₂ saturation, P wave amplitude is more sensitive and as saturation increases, the change in wave amplitude is minimal due to energy loss between interfaces of brine and CO₂ saturated pores. Although the Gassman equation and its derivatives can predict the change in velocity and elastic modulus as a function of fluid saturation (Berryman, 2012; Diallo et al., 2003; Mavko and Nolen-Hoeksema, 2012a; Sevostianov, 2020), this doesn't account for the associated stress change in the reservoir.

In this study we simulate a CO₂ injection operation in a critically stressed (i.e beyond yield stress and before failure) brine-saturated sandstone coreplug to study how the ultrasonic wave attributes respond to stress change, fluid displacement and a coupling of both these factors. The experimental protocol adopted for this study emulates a situation where CO₂ is injected in a saturated and critically stressed saline aquifer. Since the principles discussed in this study will be somewhat similar for depleted oil and gas reservoirs as well, we will use the term 'reservoirs' going further to represent a broader array of potential CO₂ reservoirs. Understanding the change in elastic wave properties during the lifecycle of CO₂ injection is the primary objective of this study, where both fluid displacement and stress perturbation influences the ultrasonic wave attributes. Our aim is to find out how the contribution of stress and fluid invasion can be isolated by looking at wave attributes in a simplistic CO₂ injection case in order to establish elastic wave properties as an effective monitoring tool in a field scale scenario.

2. Methodology

2.1. Sample preparation

The core samples used in this work (Table 1) are sourced from different depths of the Aramis CO₂ storage license area in the North Sea (<https://www.aramis-ccs.com>). This is a depleted gas field where the reservoir rocks are primarily sandstone. The sandstone coreplugs were provided by Shell through TNO (Nederlandse Organisatie voor Toegepast Natuurwetenschappelijk Onderzoek). The geology of the area is well studied (De Jager and Geluk, 2007; Geluk and Röhling, 1997) and the geological maps as well as well logs are openly available through NLOG (<https://www.nlog.nl/>). Air permeability of studied depth intervals (or close) measured at 400 bars are reported from NLOG data repository. The collected samples belong largely to Upper and Lower Slochteren and middle Solling formation from offshore Netherlands. The middle Solling Formation consists of fine to medium well sorted sand grains formed by Eolean dune deposit. It has a high porosity, permeability and low heterogeneity (Filomena et al., 2012). The upper Slochteren with dominant sabkha facies on the other hand has impaired areas due to low permeability and a high heterogeneity (Van Hulst, 2006). The Lower Slochteren Formation has high degree of burial diagenesis and salt deposition in some depth intervals, causing heterogeneity in porosity and permeability (Van Hulst, 2006). The collected samples were first trimmed and polished to maintain parallel end faces with a final diameter of 1±0.1 inch and length of 2±0.1 inches. The finished samples were further dried at 60°C in an oven for 24 hrs and then their porosity and grain density were measured with a helium pycnometer. The samples were then vacuum saturated in a low salinity brine solution (80 g/L NaCl) for 24 hrs to allow maximum brine saturation before the experiments.

2.2. Experimental protocol

The experiment was planned in a way to closely represent the stress change in a critically stressed depleted saline aquifer during CO₂ injection. In the case of a saline aquifer, when CO₂ is injected into the reservoir, it displaces the brine from the pore spaces and increases the pore pressure simultaneously. When the reservoir is critically stressed, too much pore pressure buildup can initiate plastic deformation in the reservoir rock, which is often identified by a deviation from linearity in the stress-strain profile, resulting in irreversible deformation within the reservoir preceding failure. In this stage, microcracks start appearing, weakening or damaging the reservoir. The sandstone coreplugs saturated in brine (section 1.1) were loaded into a custom-made Hoek cell which is able to apply confining pressure onto the coreplug. The Hoek cell was then connected to a 500 kN uniaxial loading frame, which can apply axial load (Fig. 1). For these experiments, load was applied by constant upward movement of the bottom platen of the load frame with a slow rate of 0.2 µm/s. The rate of movement was controlled by two Linear Variable Differential Transformers (LVDTs) with a maximum displacement of 2 mm connected to both ends of the bottom platen (Fig. 1a). The stainless-steel pistons connected to the top and bottom of

Table 1

Samples used for this study along with their origin, connected porosity and grain density. The permeability values are extracted from NLOG.

Well ID	Formation	Sample name	Depth (m)	Porosity	Permeability* (mD)	Grain density (g/cc)
K15-12	Upper Slochteren	K15-12-5V	3930.05	0.11	0.32	2.71
K15-FG-102	Upper Slochteren	K15-FG-102-3V	4176.8	0.17	3.2	2.69
L09-10	Solling	L09-10-7V	3172.2	0.23	933	2.55
L09-10	Solling	L09-10-16V	3212.2	0.13	69	2.62
K15-15A	Lower Slochteren	K15-15A-7V	4236.18	0.11	72.9	2.75
K15-15A	Lower Slochteren	K15-15A-8V	4236.81	0.09	66.1	2.73
K15-15A	Lower Slochteren	K15-15A-12V	4249.3	0.14	0.18	2.73

* Permeability for exact same depth intervals were not present in most cases, therefore corresponding permeability of closest possible depth having similar porosity was reported.

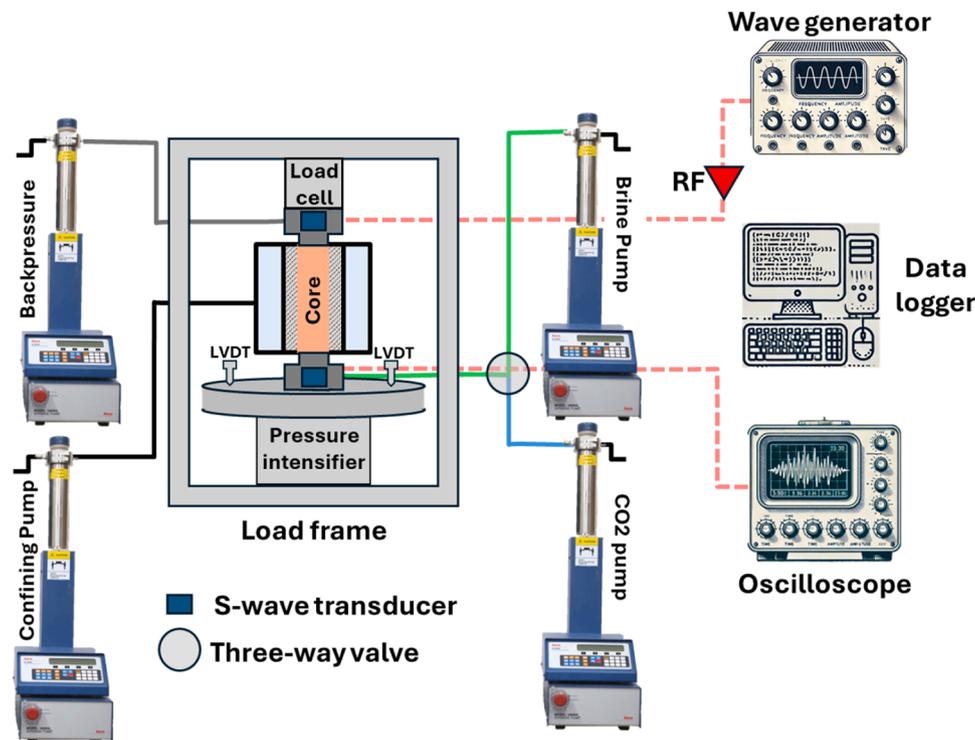


Fig. 1. Experimental setup used for this study containing the loading, flow and sonic measurement system.

the sample was also custom-made and each of them houses a 1 MHz S wave transducers of 5 mm diameter. The pistons also allow fluid flow through them without affecting wave propagation. Porous meshes were used between the sample and piston in both ends to ensure even distribution of fluid on each end of the coreplug. One of the S wave transducers was connected to the wave generator through an amplifier, which acted as the source and the transducer on the other end of the coreplug acted as a receiver which was connected to an oscilloscope and datalogger to automate saving of waveforms. A small component of the shear wave energy converts to compressional wave at the contact between anisotropic media (in this case between transducer-piston-specimen), resulting in small vertical displacement. This converted mode is recorded at the receiver and attributed as the P wave arrival time for the experiments. Technical specifications of the acoustic setup can be found in (Veltmeijer et al., 2024). It is worthwhile to mention that the ultrasonic waveforms were recorded every 10 s for 500 μ s and averaged over 128 stacks for an enhanced signal-to-noise ratio. A sampling frequency of 2 MHz was used for recording the waveforms, which resulted in a 0.5 μ s picking error in P and S wave arrival times for each experiments. Waveforms collected during the experiments were processed with Radex Pro software to pick V_p , V_s , P wave amplitude (Pamp) and S wave amplitude (Samp). Pamp and Samp represent the highest amplitude of a 10 μ s symmetric window around the P and S wave arrival time. The V_p and V_s were calculated after deducting the travel time of the pistons. A total of four ISCO 100DM syringe pumps were used to apply fluid pressure in the system. One of the pumps was used for building confining pressure. The outlets of two pumps (one for 80 g/L NaCl brine and another for pure liquid CO₂) were connected with a three way valve, which was further connected to the bottom piston for pore fluid injection so that only one fluid can be injected at once. One pump was connected to the top piston for applying backpressure and also to monitor pore fluid pressure buildup in the sample.

All pore fluid lines were made from 1/4 inch stainless steel tubes with a cumulative dead volume of 7 ml. To represent the stress condition of an intermediate-depth CO₂ storage reservoir (1–2 km deep), a hydrostatic stress of 40 MPa is selected as the starting point of the experiment. The

experiments were conducted in the following steps (as indicated in Fig. 2a, b):

- We built a 40 MPa hydrostatic load at 10 MPa/min rate. 20 pore volume (PV) of brine was injected into the sample at 5 ml/min and was allowed to bleed from the downstream line. This step ensured maximum escape of air (if any) from the flow lines or sample. The top piston was connected to the backpressure pump to build a constant nominal pore pressure of 0.1 MPa. A thin layer of viscous ultrasound gel was applied between the piston-sensor and piston-sample contact (in both pistons) for uniform transmission of waves. The parallel end faces of the coreplugs ensured even load distribution and optimum contact between sample and pistons.
- Timelapse ultrasonic velocity measurement was started along with axial stress buildup by constant upward movement of the bottom platen at 0.2 μ m/s. The stress-strain profile was closely monitored till the stress profile started changing its slope (start of plastic deformation). At this point, the loading mechanism was changed from deformation controlled to load controlled. This allowed the machine to maintain a constant axial stress and adjust piston movement accordingly.
- The brine injection pressure was raised up to 6 MPa at 1 MPa/min rate. The pressure was selected to represent a depleted aquifer and also since liquid CO₂ (which is the displacing fluid) is unstable below 6 MPa at laboratory temperature. The laboratory was centrally heated to maintain a constant temperature of 19°C, which is below the boiling point of pure CO₂ (21.6°C) at 6 MPa. The experimental setup was covered with a thick insulating cover to minimize influence of sudden temperature changes. A room thermometer was placed close to the sample enclosure to monitor any changes. While CO₂ stays in supercritical phase within the reservoir, due to lack of heating module in our experimental setup and to avoid the freezing effect due to the liquid-supercritical transition of CO₂, we selected liquid CO₂ as our pore fluid. We discuss later, how the observations might vary when supercritical CO₂ is used as pore fluid instead of liquid CO₂. It is worthwhile to mention that the samples tested are

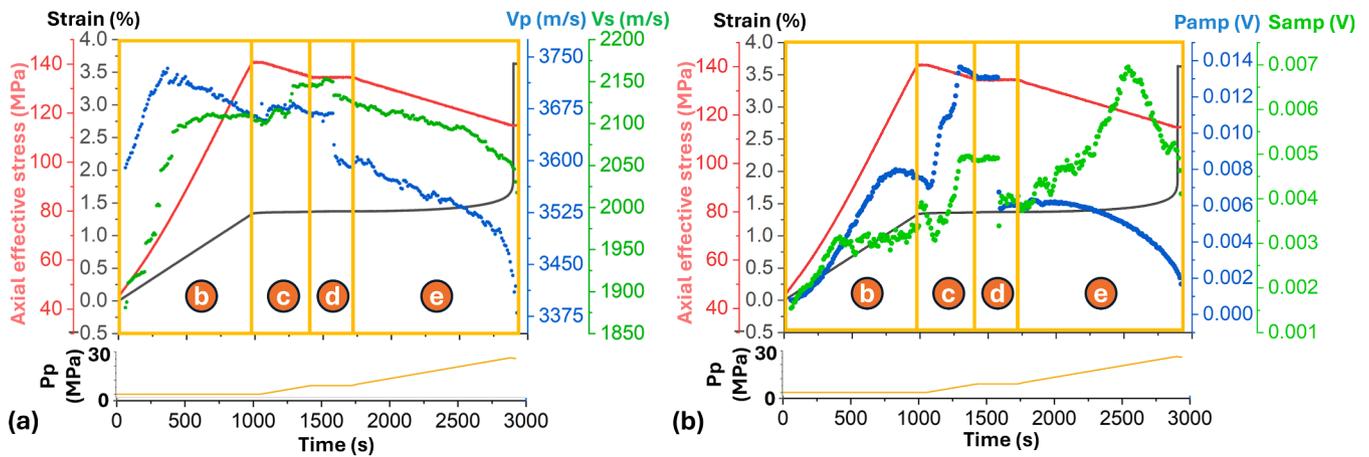


Fig. 2. An experiment with different stages indicated in circles (description in text) along with (a) Vp and Vs measurement and (b) Pamp and Samp measurement. The mini plots below each figure indicates pore pressure development over time. Pp indicates pore pressure.

- highly porous and permeable which ensured that the pressure difference between upstream and downstream pump never exceed 0.5 MPa. Once desired pressure is achieved, wait for pressure to equilibrate.
- d. A liquid CO₂ pressure of 6 MPa was built up to the three-way valve, and further allowed to flow into the plug by turning the valve. A drawdown rate of 5 ml/min was applied in the backpressure pump causing a slow drop in line pressure and allowing CO₂ to flow into the sample. After drawdown of 30 PV, we assumed that the coreplug is mostly saturated with CO₂ and some brine (similar to reservoir scenario).

- e. The drawdown was stopped from backpressure pump and CO₂ inlet pressure started increasing at 1 MPa/min and continued until the sample fails.

3. Results

In stage B (for nomenclature of the stages refer Fig. 2) as the axial stress is increased, both Vp and Vs increase uniformly during elastic deformation stage, followed by a nearly identical rate of decrease since the onset of plastic deformation (Fig. 3).

Interestingly, Vp of the sandstones from L09 well doesn't show

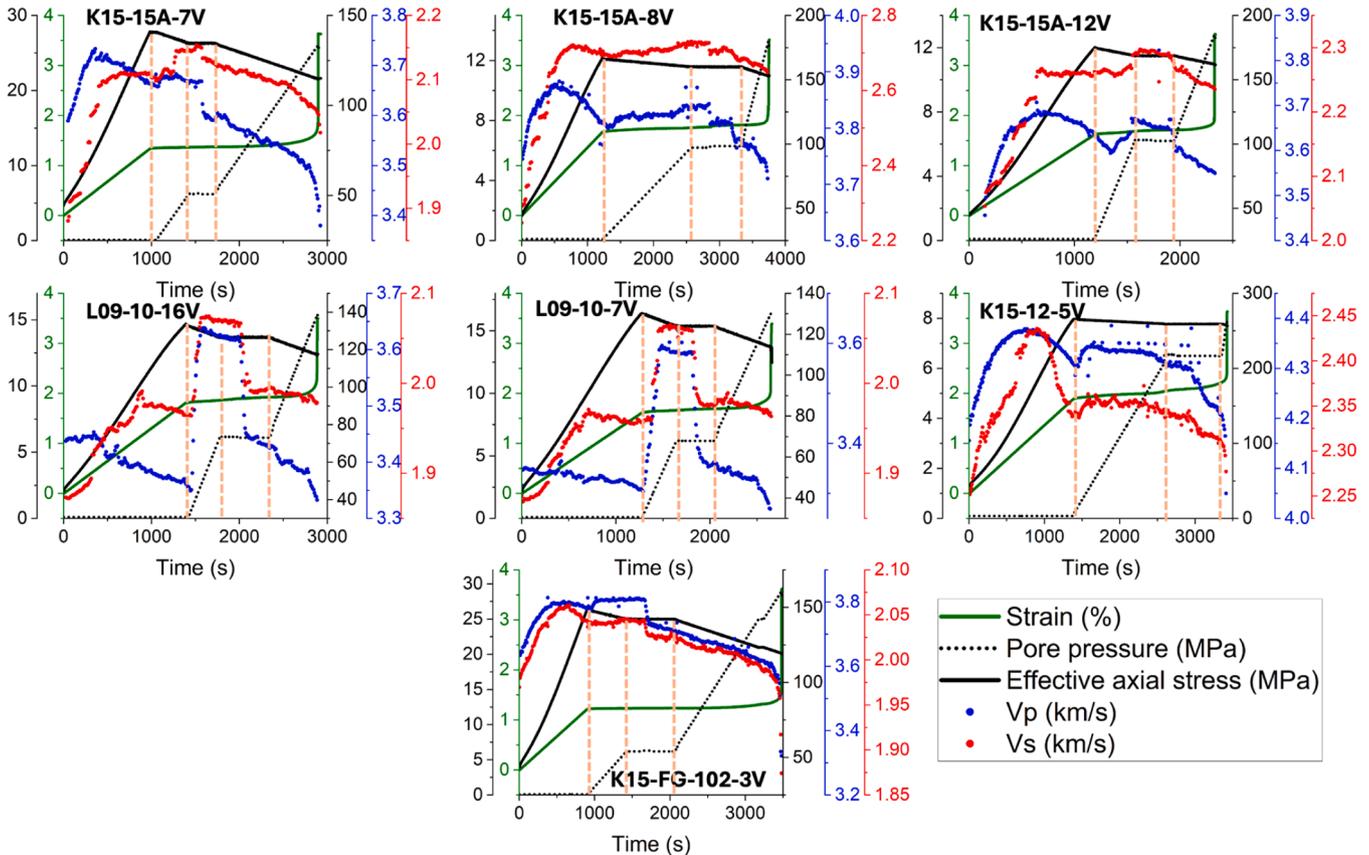


Fig. 3. Change in Vp and Vs as a function of stress change and fluid displacement until the point of failure during the triaxial tests. The dotted orange vertical lines indicate transition stages 'b' through 'e' of each experiment.

significant increase during the loading stage, which follows the Mori-Tanaka relationship and experimental study by [David and Zimmerman, \(2012\)](#) and is commonly observed in saturated high-porosity rocks. Since the start of brine injection in stage c, both V_p and V_s show minimal change in the beginning and only after the fluid pressure reaches 1–2 MPa, there is a sudden and significant increase in the velocities. It is to be noted that for low porosity samples (e.g K15-12-5 V and K15-15A-8 V), despite maintaining a constant rate of pressure increase, a longer time was allowed for the differential pressure across the coreplug to reach close to 0.1 MPa to ensure pressure equilibrium in the whole specimen. After the pore pressure reaching 6 MPa, liquid CO₂ is flushed through the sample (stage d) which at first causes minor drop in velocity when the coreplug is partially saturated with CO₂, followed by a rapid drop in velocities due to lower density of liquid CO₂ compared to brine ([Fig. 3](#)). In the last stage of the experiment (stage e), liquid CO₂ pressure is increased at a uniform rate, which causes a decrease in velocities as the pore pressure starts opening microcracks and induces new crack formation in the critically stressed coreplugs. Continuing pore pressure increase increases the rate of axial deformation and in moments preceding failure, the axial strain exponentially increases. V_p and V_s concurrently decreases rapidly when the axial strain increases fast, signifying axial failure of the specimen. In general, change in V_p is more pronounced with pore pressure increase compared to change in V_s . For K15-12-5 V, the failure occurred after minimal CO₂ pressure increase and therefore the number of datapoints for that experiment are fairly less than the other samples.

Similar trend can be seen for Pamp and Samp in different stages of the experiments. Both of them show high sensitivity with pore pressure increase and CO₂-brine displacement compared to stress change ([Fig. 4](#)). Interestingly in stage e, the amplitudes behavior differently from the velocities. The amplitudes, especially Samp increase as the CO₂ pressure increases with only exception being in low-porosity samples (K15-15A-8

V and K5-12-5 V).

The increasing trend for Samp and in some cases Pamp in stage e continues as long as there is minimal change in axial strain. As soon as the coreplug approaches failure and axial strain increases exponentially, both Pamp and Samp drops.

4. Discussion

For ease of understanding we subdivide the experiments in four stages (b-e) to discuss the key finding in each stage as follows. We skip discussing stage A, where only hydrostatic loading is performed to bring the axial and confining stress to 40 MPa.

4.1. Stage B—approaching critical stress conditions

Starting from hydrostatic stress condition, axial stress increment also causes increase in V_p , V_s and their corresponding amplitudes ([Fig. 3, 4](#)), barring samples belonging to L09 well, where there is negligible increase in V_s and especially V_p . The rate of increase in V_p and V_s is also strongly dependent on porosity as validated by [Chandra et al., \(2025\)](#). Interestingly, the change in V_p and V_s or their amplitude does not happen at the same rate even within the same coreplug. Up to 100 MPa, all of the coreplugs go through elastic deformation and in that period Samp/Pamp drops sharply ([Fig. 5a](#)) (except K15-12-5 V). A linear rescaling method was adopted to normalize the Samp/Pamp and V_p/V_s values. For example, minimum Samp/Pamp value for each experiment is considered as 0, whereas the maximum value is considered to be 1. This method preserves the relative spacing of the values and makes comparing parameters across different samples easier. Similar normalizing scheme is used in all following instances.

During this phase the V_p/V_s ratio also drops ([Fig. 5b](#)) which indicates relatively faster increase in Pamp compared to Samp but slower

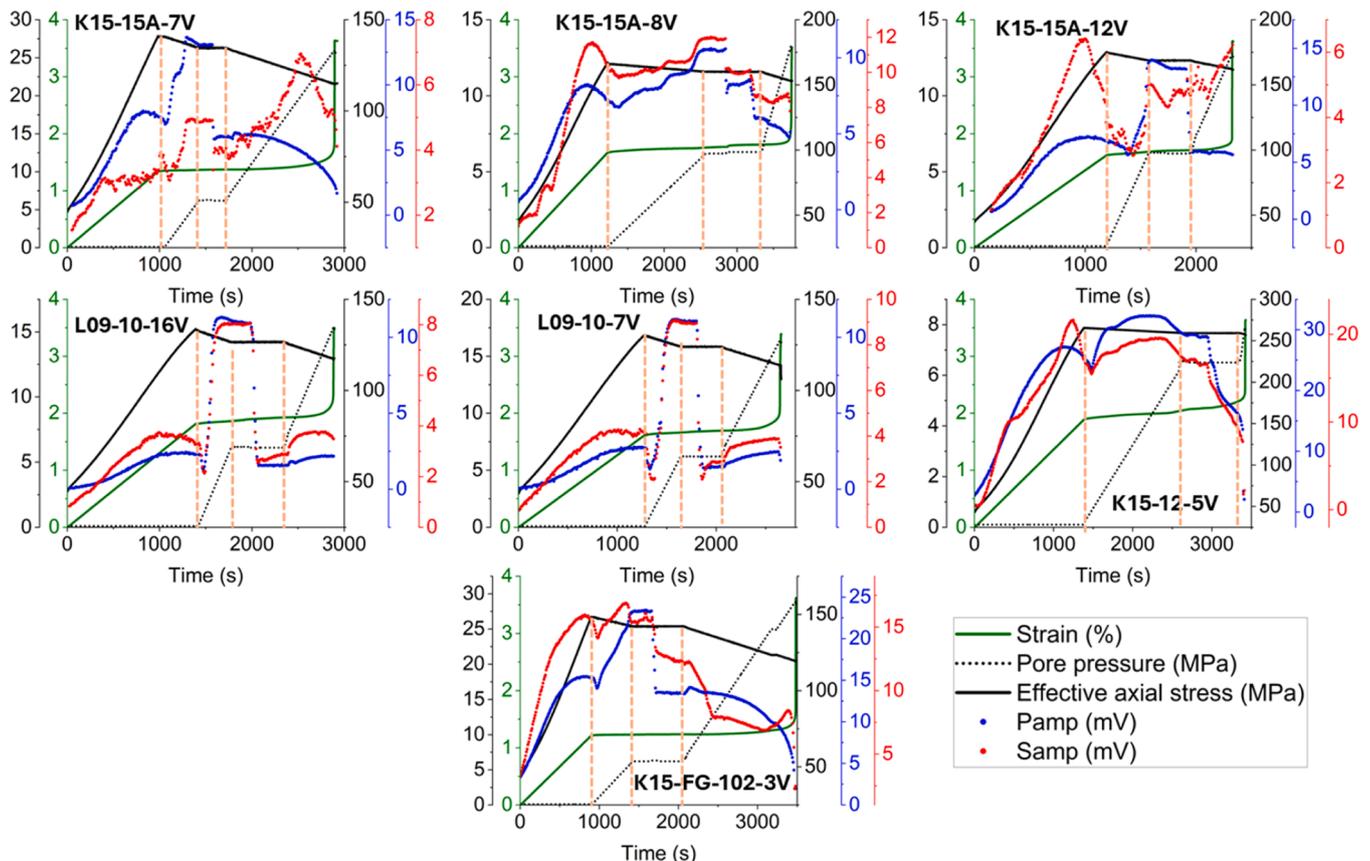


Fig. 4. Change in Pamp and Samp as a function of stress change and fluid displacement until the point of failure during the triaxial tests.

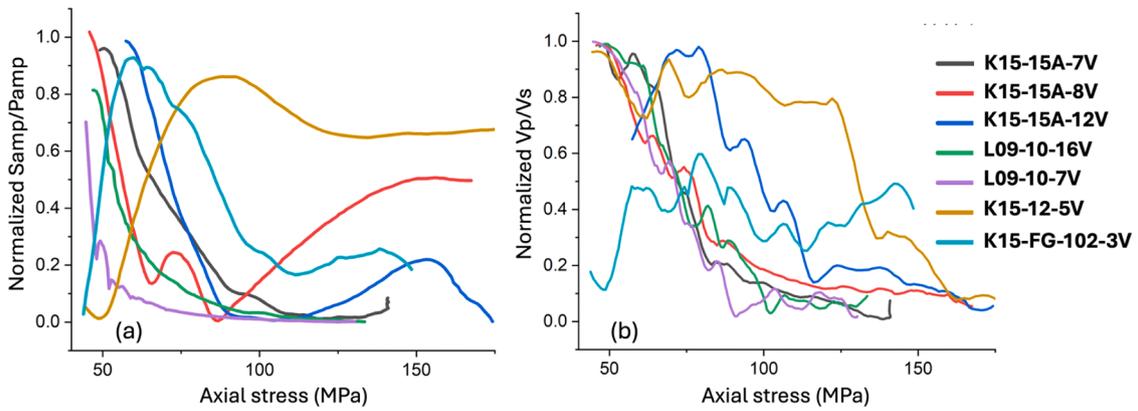


Fig. 5. (a) Normalized Samp and Pamp ratio and (b) Normalized ratio of Vp and Vs as a function of axial stress during the axial loading phase of the experiment.

increase in Vp compared to Vs. While transitioning into plastic deformation, there is a subtle trend reversal in Samp/Pamp, where increase in Samp happens faster than Pamp (Fig. 5a). However, there is no such trend reversal in the Vp/Vs ratio. Transitioning into plastic deformation, Vp/Vs maintains the decreasing trend, however the rate of decrease is

slowed down significantly.

4.2. Stage C-approaching depleted reservoir condition

In this stage we start increasing the pore pressure in the coreplug

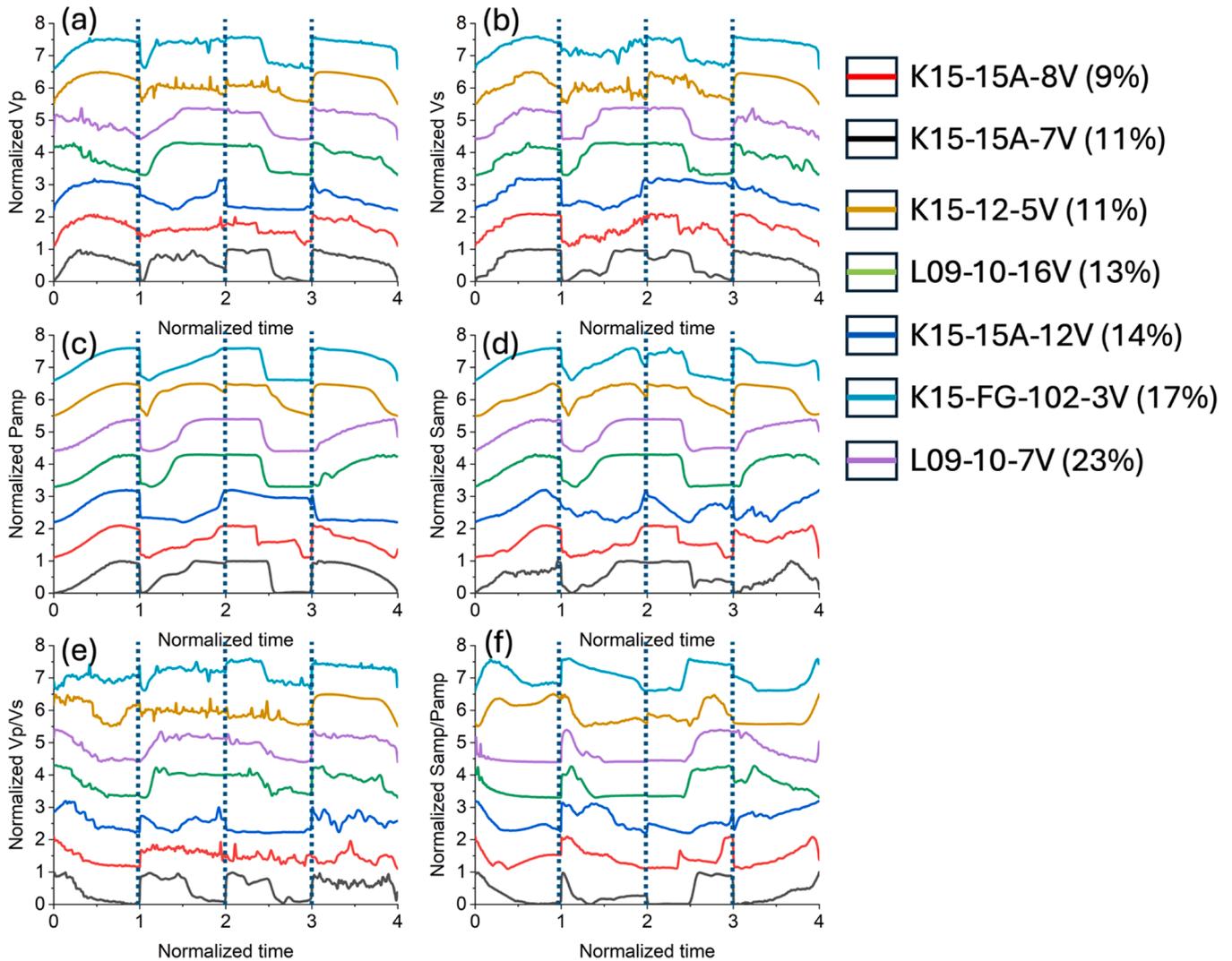


Fig. 6. Change in different ultrasonic attributes throughout the experiments (a) Vp, (b) Vs, (c) Pamp, (d) Samp, (e) Vp/Vs, (f) Samp/Pamp. Each experiment is divided into 4 parts following the protocol mentioned in Fig. 2. Values of sonic attributes and time are normalized within each part of the experiment for better comparison across all samples.

with NaCl solution while the axial and radial stresses are kept constant. The samples are allowed to deform freely as long given the constant stress. In theory, increase of the pore pressure within the elastic limit of the coreplug will lead to an increase in V_p and V_s due to increasing density of pore fluid contained within the pores. A higher pore pressure will also allow filling of more pore spaces, which might have remained unsaturated during the vacuum saturation stage. Interestingly, despite a constant increase in pore pressure, the change in velocity in this stage is abrupt. Up to almost 3 MPa pore pressure, the velocity increase is slow, and suddenly halfway through, there is a sudden increase in V_p , V_s and their corresponding amplitudes (Fig. 6). This abrupt change in velocities is validated by earlier studies on anisotropic geological samples with complex pore structure (Berryman, 2004; Knight et al., 2012; Lebedev et al., 2009; Zhang et al., 2017) described as patchy saturation model which doesn't quite follow the Gassman-Wood or Gassman-Hill trends. This phenomenon occurs due to differential saturation of the porous rock when fluid pressure starts to increase at the upstream, depending on the permeability, the downstream end is still at a lower pressure and therefore not completely saturated. Due to dispersion of waves at the interface of partially and fully saturated part of the coreplug, the velocities and amplitudes show minimal change in the beginning. When the fluid pressure increases from one end, the fluid front progresses further, saturating more pores, which results in an increase in velocity and amplitudes. Depending on the complexities of flow paths, the increase in velocity and amplitude is stepped or gradual. Since the measurements are conducted with high frequency waves, the delayed velocity response can also be explained by capillary pressure equilibrium theory (Dasgupta et al., 2024; Qi et al., 2014b; Sen and Dvorkin, 2013), which follows similar principle as squirt flow (Gurevich et al., 2010; Ógúnsàmi et al., 2025; Sun, 2021; Sun et al., 2020). The capillary pressure-based model assumes a heterogeneous media, which is composed of multiple patches with different porosity, permeability and capillary threshold in each patch. In each patch, capillary forces resist rapid fluid movement, which delays the equilibrium. Until then, the velocity and amplitude responses remain in transitional state and do not reflect the new fluid configuration. Both capillary pressure-based model and squirt flow justify the hysteresis in the velocity-saturation trends, causing decoupling between stress state and immediate acoustic response, which is evident in stage B and C for all experiments (Fig. 6). However, in order to quantitatively discern the effect of capillary pressure-based model, degree of heterogeneity needs to be measured across the coreplug at a resolution smaller than the characteristic wavelength (smaller than ~ 3.3 mm for these set of samples), and such microstructural analysis is beyond the current scope of study. It is worthwhile to mention that there is a small drop in the velocity and more observable drop in amplitude when the brine injection starts (Fig. 4, 5). Given that the pressure front moves faster than the fluid front in complex porous materials (Brantut and Aben, 2021; Chapman et al., 2021), and considering that the rock is already in plastic deformation zone, even slight increase in pore pressure induces more microcracks in the coreplug, which further attenuates compressional and shear waves through the media. The minimal change in V_p/V_s ratio during this stage indicates that the wave velocities increase by the same proportion. However, the $Samp/Pamp$ ratio decreases steadily during the pore pressure buildup. There is also considerable research performed on the evolution of velocities as a function of injection rate (Lopes et al., 2014; Lopes and Lebedev, 2012; Qi et al., 2014a), however in our present study, we set control on the rate of fluid pressure while the injection rate was allowed to change freely.

4.3. Stage D-effect of fluid displacement on wave propagation

In this stage of the experiment, we start pushing liquid CO_2 (lCO_2) through the coreplug at a constant upstream pressure and by applying 5

ml/min drawdown rate with the backpressure pump. As expected, due to the lower density of liquid CO_2 than brine, the velocity and amplitudes decrease during the fluid displacement phase. In a reservoir scale CO_2 injection operation, we expect this situation in the beginning, where CO_2 injection only displaces the brine partially out of the pores (depends on the wettability of the pores) with minimal influence on pressure. Similar to stage C, in this case also, the velocity and amplitude response is a bit delayed (Fig. 6). Generally, in CCS projects, depending on the gradient between injection pressure and initial pore pressure, CO_2 injected in liquid state gets converted to supercritical state. However, to achieve experimental simplicity, we chose the pressure conditions in such a way to avoid phase change or freezing effect involved with liquid to supercritical transition of CO_2 , and also at the experimental timescale, the reactivity of CO_2 with brine or host rock is assumed to be minimal. Since the pressure is kept constant during the fluid displacement, we assume no further deformation in the coreplug takes place. The late and sudden drop in velocity and amplitude can be attributed to the patchy saturation model discussed in the previous subsection. Other attributes like V_p/V_s and $Samp/Pamp$ show opposite trend of the previous stage, i.e. the V_p/V_s ratio drops and $Samp/Pamp$ ratio increases synchronously to the change in velocities.

4.4. Stage E-pore pressure increase leading to failure

In this final stage, we increase the pore pressure keeping the axial and radial stresses constant and this is where the deformation and velocity has a complex relationship. On one hand increase in liquid CO_2 pressure increases the density of pore fluid, resulting in faster wave propagation (Janssen et al., 2023), whereas the increasing plastic deformation in the coreplug due to increasing pore pressure attenuates wave propagation. Risk of failure due to pore pressure injection is well documented in previous studies using experimental and numerical approach (Damani et al., 2018; Frempong et al., 2007; Komijani et al., 2019; Mayr et al., 2011; Rafieepour et al., 2021). In these experiments the start of pore pressure increase immediately leads to a decrease in V_p and V_s , signifying that inelastic deformation plays a major role in governing the velocity change during pore fluid injection. Similar velocity change due to inelastic deformation was also observed in Stage B, however, the deformation was caused only due to axial loading. Upon comparing the change in velocities due to inelastic deformation in both Stage B and E, we found generally higher magnitude of change in Stage E compared to Stage B (Fig. 7a, b), which is intuitive, since more inelastic deformation close to failure occurs in the last stage. However, this also means that presence of liquid CO_2 in the pore spaces doesn't slow down the rate of decrease in velocities. Magnitude of change in V_s for K15-12-5 V was higher in Stage B already, due to more inelastic deformation incurred (as discussed in section 3), hence we did not consider those datapoints for linear regression fit. Interestingly, with increasing porosity, the magnitude of velocity drop also increases for both V_p and V_s . Conversely, $Pamp$ and $Samp$ shows a reverse trend, where higher porosity reduces the magnitude of change in amplitudes (Fig. 7c, d). We can get a clue for such behavior also from Fig. 4, which indicates that if inelastic deformation attenuates wave amplitudes, the increasing pressure and density of pore fluid tries to increase the amplitudes. Finally, close to failure, the amplitudes decrease rapidly as it cannot keep up with the exponentially increasing crack volume (indicated by the axial strains in Fig. 4). Coming now to Fig. 7c, higher porosity indicates higher volume of liquid CO_2 in the pore space, which contributes to a subtler change in $Pamp$. On the contrary, in low porosity samples, where waves travel mostly through the matrix, increasing volume of cracks considerably bring down the amplitudes, and the pore fluid density has minimal role to play.

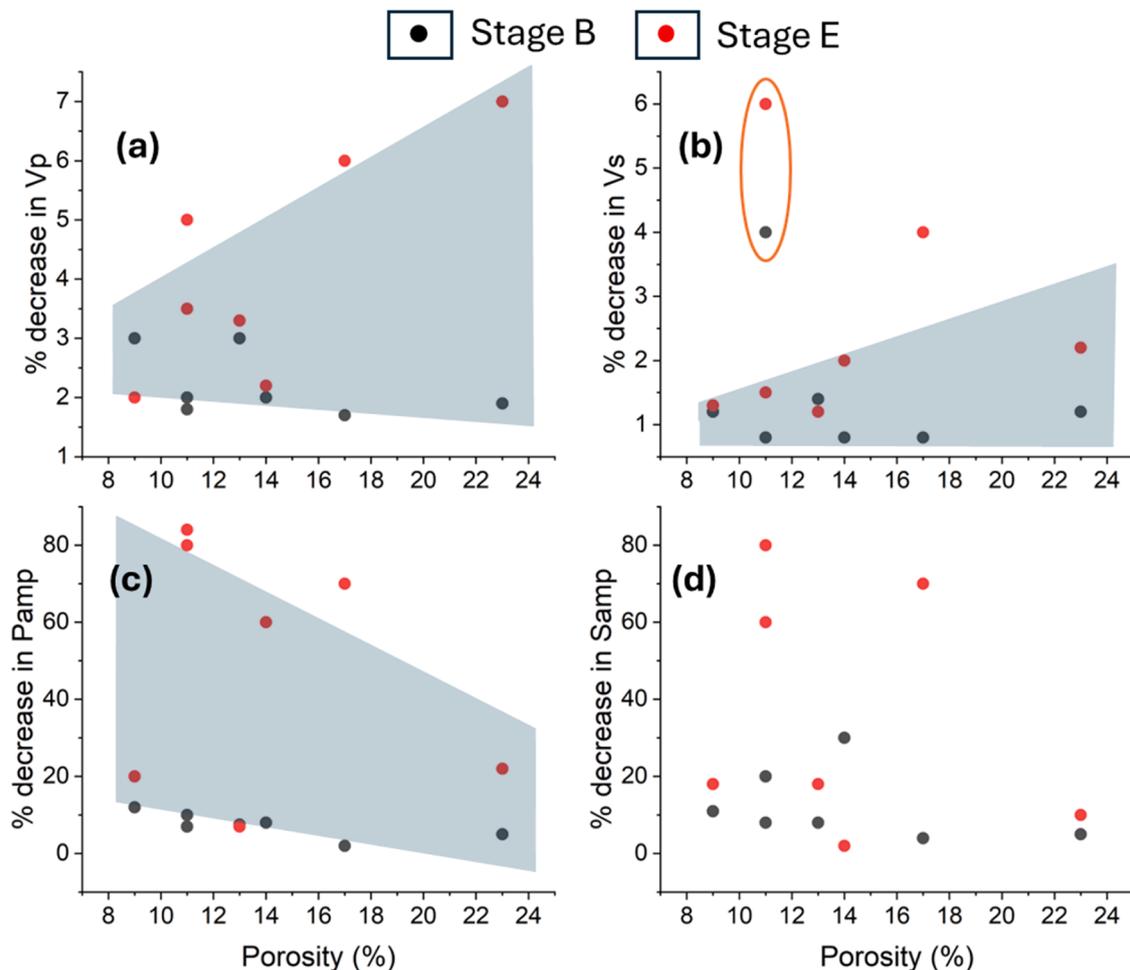


Fig. 7. Comparison of the percentage decrease in sonic attributes (a) Vp, (b) Vs, (c) Pamp and (d) Samp in Stage B and Stage E of the experiments. The circled dataset in (b) belongs to K15-12-5 V. The color-filled region represents the area between the linear regression fits of Stage B and E datasets. Due to lack of strong linear fit, there is no such region marked in (d).

4.5. Applicability for reservoir monitoring at seismic frequency

Different forms of seismic monitoring are widely used for monitoring plume migration in CO₂ storage complexes. Leveraging the difference in acoustic impedance of reservoir brine and injected CO₂, migration of CO₂ in the subsurface can be traced. Although combining other forms of monitoring (e.g resistivity, electromagnetic) provides better resolution (Fawad and Mondol, 2021), but might drive up the monitoring significantly (Mathieson et al., 2011). In a unique field-scale controlled CO₂ leakage experiment (Michael et al., 2020) found that the change in CO₂ plume is clearly detectable in downhole seismic measurements. Vertical seismic profiling (VSP) along with surface seismic measurement is currently used extensively for monitoring plume migration near injection well. Effectiveness of surface seismic to monitor CO₂ plume migration was demonstrated by Arts et al., (2004) in Sleipner. Timelapse VSP was also used in the Frio Pilot well to monitor CO₂ plume near injection well (Daley et al., 2007, 2008a), whereas a more cost effective permanent and semi-permanent deployment was setup in Penn-West and Otway Project respectively (Chalaturnyk et al., 2006; Daley et al., 2008b; Lawton et al., 2006), which additionally also provides opportunity for microseismic monitoring. 2D and 3D VSP monitoring is also successfully implemented in Weyburn, Decatur, and SACROC project (White et al., 2021). Techniques like crosswell seismic monitoring is deployed to get around the lack of resolution in traditional VSP monitoring. Permanent or semi-permanent continuous crosswell monitoring has also been used successfully in Cranfield, Aquistore, Frio Pilot and

Nagaoka Pilot, which could monitor changes in CO₂ plume at a high spatial and temporal resolution based on velocity change in the sub-surface. Using a combined reservoir simulation, rock physics and synthetic seismic workflow, Eid et al. (2015) was able to robustly detect CO₂ plume growth and overall migration, especially when there are structural baffles creating accumulations. However, their approach had considerable uncertainty while inverting for exact saturation due to the uncertainty in extent of heterogeneity. Many other studies (L. Hu et al., 2017; Huang et al., 2018; Ivandic et al., 2018; Kolkman-Quinn et al., 2023; Picotti et al., 2012; Wang et al., 2020; Zhu et al., 2019) have expressed similar uncertainty while inverting for exact CO₂ saturation using different approaches including timelapse 3D seismic, borehole seismic, and synthetic seismic simulation using direct arrivals, AVO (amplitude versus offset) or coda waves. The degree of uncertainty rises when the reservoir is highly heterogeneous and depends on sensor coverage, their proximity to plume and repeatability of the active seismic sources.

Although several instances of plume migration monitoring are well documented with long-term field scale application, there is no evidence of direct stress monitoring using calibrated velocity data. Prasad et al. (2021) highlighted that with information of crack density, soft grain contact and experimental data on stress-velocity relationship, pressure change within the reservoir can be quantified. Chandra et al. (2025) already demonstrated quantification of crack density from ultrasonic P and S waves in laboratory scale. Several experimental studies exist that attempts to bridge the gap between ultrasonic and seismic scale stress

and saturation-dependent velocity changes.

For saturated porous sedimentary rocks, previous studies have pointed out considerable acoustic dispersion when comparing ultrasonic and seismic frequency measurements (Chapman et al., 2019, 2016; Nur and Simmons, 1969; Pimienta et al., 2015). It is postulated that in a global scale, flow dispersion in saturated media arises due to the forced motion of the fluid caused by induced waves. This flow dispersion is usually caused by low-frequency waves and in larger pores or channels. In a pore scale, high-frequency waves may cause pore fluid to exhibit stiffer elastic behavior, which is commonly known as ‘squirt flow’ (Batzle et al., 2006; Müller and Gurevich, 2004; Tisato et al., 2021). It is observed that the changes in velocity at different stress conditions caused by dispersion cannot be justified uniformly with the squirt-flow model (Ba et al., 2024; Tisato et al., 2021). During an injection scenario, when the pore pressure is significantly low and starts increasing gradually, the decrease in V_p can be explained by a modified acousto-elasticity model incorporating Mori-Tanaka scheme and crack density model (David and Zimmerman, 2012; Mori and Tanaka, 1973). This closely follows a stress-dependent linear change in velocity. At seismic frequency, the physics remains the same, however the effect of squirt flow is less pronounced. Absolute magnitude of V_s shows minimal deviation with dispersion, however, V_p is much higher for saturated porous rocks at higher frequencies. When pore pressure approaches the confining pressure, the rate of velocity drop accelerates can be better explained by the squirt-flow model, since more channels and pore spaces are filled with pore fluids and there is higher chance of the pore fluid exhibiting a ‘stiff’ behavior (Ba et al., 2024). As expected, their experimental results indicate stronger influence of squirt-flow at ultrasonic frequency and vice versa. Additionally, the increase in pore pressure results in a proportionally faster decrease in velocity in both seismic and ultrasonic V_p , which is also evident in our ultrasonic experimental findings (Fig. 8a). Changes in V_s is also pore pressure-dependent, but is less affected by squirt-flow mechanism (Fig. 8b).

On the other hand, multiple studies reported contrasting response in wave velocities due to imbibition and drainage of non-wetting phases (Knight and Nolen-Hoeksema, 1990; Li et al., 2001; Mavko and Nolen-Hoeksema, 2012b). Fluid imbibition and drainage processes, especially of different wetting behavior is often complicated and hard to fit in a generalized physics. Various research indicates that the velocity change during imbibition of non-wetting phase is often sudden. There is

negligible change in velocity in low saturation and a sudden change at a smaller saturation interval. This behavior is often attributed to the residual fluid distribution in sandstone (varies between 10–40 %) which depends on particle size, pore size and structure, grain contact etc. (Wilson and Conrad, 1984). Bacri and Salin, (1986) also pointed out that the change in velocity is more gradual when miscible phases (vapor against water) are displaced. In our experiments, liquid CO_2 is a non-wetting phase in brine saturated sandstone in stage D, which causes similar change.

Si et al. (2016) demonstrated that increase in saturation in sandstone cause pulsed and exponential increase in V_p and comparatively nominal increase in V_s . Gassman-Biot and Gassman-Hill theories also indicate a decrease in V_p at low saturation, followed by rapid increase which could also be observed in some of our coreplugs (Fig. 9). More pronounced changes in velocity and amplitudes could be seen in high porosity coreplugs, however, other petrophysical properties (grain size, permeability, mineralogy) of these specimens were different, which led to an inconclusive correlation between porosity and magnitude of velocity change. In Stage C of our experiments, we increase brine injection pressure, which causes first a quick increase in V_p , V_s and then remains constant (Fig. 10), possibly due to reduction in effective stresses.

4.6. Key indicators for reservoir-scale CO_2 injection monitoring

As mentioned before, the objective of this study was to find out the velocity and amplitude change pattern during different stages of CO_2 injection in a depleted saline aquifer. Displacement of brine by CO_2 which is expected during CO_2 injection, causes decrease in velocity and amplitudes, whereas irreversible deformation in a reservoir can also give similar indications. This study is also aimed to refine the traffic-light protocol proposed in Chandra et al. (2025) and give more robust information on the difference between stress induced velocity changes and fluid injection induced velocity changes. In order to avoid false indicator in monitoring, the traffic light system needs to be more robust so that we can differentiate effect of fluid and effect of stress. In the experiments discussed above, increase in V_s in Stage B is caused by elastic deformation of the coreplugs, where pore pressure remains constant. In Stage C, stress and corresponding velocity change is caused by pore pressure. Finally in Stage D, there is no stress change, but displacement of brine by liquid CO_2 . We observed very different velocity and amplitude response

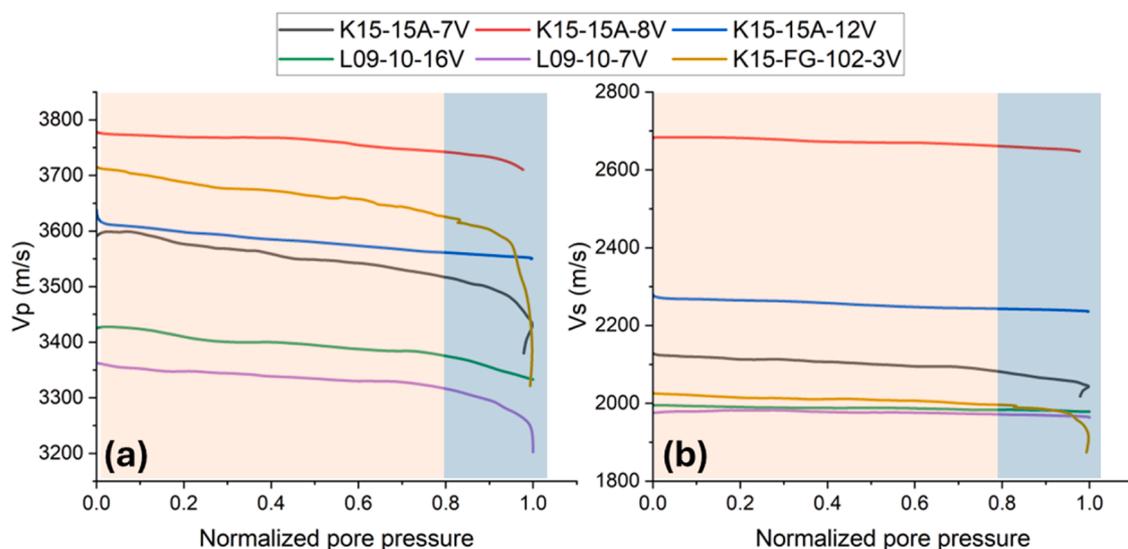


Fig. 8. Stage E of the triaxial experiments showing (a) V_p and (b) V_s drop with increase in pore pressure leading to failure. Since the pore pressure required for failure varies for different coreplugs, we normalize the values to make them comparable. The orange shaded zone represent velocity drop following acoustoelastic principle, whereas the blue shaded zone represent squirt-flow dominated velocity drop. K15-12-5 V is excluded since there were very few representative datapoints in Stage E.

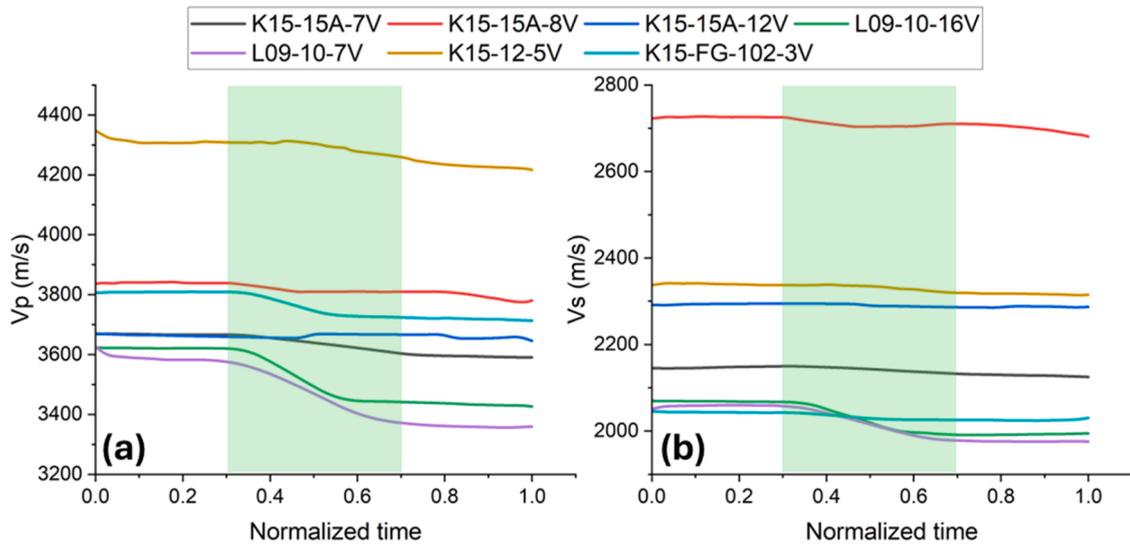


Fig. 9. The drop in (a) Vp and (b) Vs during liquid CO₂ flushing in brine saturated sandstone (Stage D). The time axis is normalized to make the data comparable across different samples. The green shaded zone indicates the period where a rapid drop in velocity was observed.

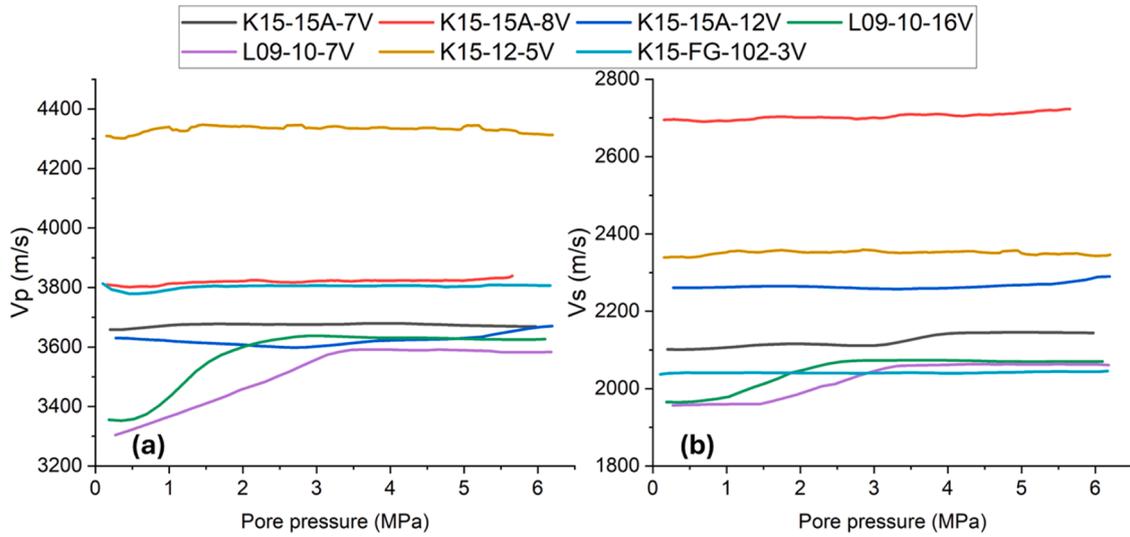


Fig. 10. Increase in (a) Vp and (b) Vs during Stage C of the triaxial tests due to increase in pore pressure with brine.

Table 2

Metrics indicating change in Vs, Samp and their ratio across different stages of the experiments. =Pp and +Pp indicates pore pressure unchanged and pore pressure increasing respectively.

Sample ID		K15-15A-7V	K15-15A-8V	K15-15A-12V	L09-10-16V	L09-10-7V	K15-12-5V	K15-FG-102-3V
Porosity (%)		11	9	14	13	23	11	17
Stage B (=Pp)	Vs*	6.36	8.39	3.64	1.85	1.97	1.71	1.99
	Samp*	4.7	5.59	4.85	4.76	6.12	12.2	18.3
	Samp*/Vs*	0.74	0.67	1.33	2.57	3.11	7.13	9.2
Stage C (+Pp)	Vs*	8.04	6.52	5.87	14.64	14.33	5.84	1.53
	Samp*	30.26	34	36	100	118	67	45
	Samp*/Vs*	3.76	5.21	6.13	6.83	8.23	11.47	29.41
Stage D (=Pp)	Vs*	-1.34 %	-1.87 %	-2.24 %	-4.06 %	-4.41 %	-1.39 %	-1.2 %
	Samp*	-45.72 %	-39.7 %	-20.14 %	-215.69 %	-327.01 %	-81.54 %	-31.08 %
	Samp*/Vs*	34.04	21.28	9	53.13	74.12	58.48	25.81
Stage E (+Pp)	Vs*	3.84	5.42	6.81	1.93	2.56	-	2.71
	Samp*	-1.66	-1.42	-2.54	-1.00	-1.06	-	-2.50
	Samp*/Vs*	-0.43	-0.26	-0.37	-0.52	-0.41	-	-0.92

in all three cases with shear waves (Table 2). Here we calculate the change in Vs and Samp per unit pressure change across all stages. The new terms Vs* and Samp* are expressed as follows (eq 1 and 2):

$$V_s^* = \frac{Vs_{max} - Vs_{min}}{\sigma^{eff}_{Vs_{max}} - \sigma^{eff}_{Vs_{min}}} \quad (1)$$

$$Samp^* = \frac{Samp_{max} - Samp_{min}}{\sigma^{eff}_{Vs_{max}} - \sigma^{eff}_{Vs_{min}}} \times 10 \quad (2)$$

For convenience in comparing, a multiplication factor of 10 is included in Samp*. Here Samp values are considered in mV. $\sigma^{eff}_{Vs_{max/min}}$ indicates the effective axial stress at Vs_{max} and Vs_{min} . It is worth mentioning that in Stage D, there was no stress change, so the changes in Vs and Samp are calculated as percentage.

The rate of increase in Vs due to pore pressure increase, even while the coreplugs are in plastic deformation regime, overpowers the rate of increase in Vs during elastic compression. The difference is even more significant for Samp, which is significantly affected by pore pressure increase. Intuitively the rate of change of Samp Vs is faster when there is influence of pore fluid. Whereas in Stage D, where there was no stress change, Vs showed minor decrease due to CO₂ imbibition, however, Samp decreases significantly compared to previous stages. In the final stage leading to failure, Vs decreases, but Samp shows minor increase until close to failure. We excluded the last few datapoint from this calculation where axial strain increases exponentially, since in context of CO₂ injection, failure (or situations close to failure) is rare. Here due to the opposite trend in Vs* and Samp*, their ratio is negative, which could act as a major indicator for CO₂ pressure increase leading to plastic deformation during field scale operation. It has been demonstrated that modern cross-well active seismic monitoring can detect extremely small variation (order of nanoseconds) in seismic travel times, making minor stress or fluid induced velocity or attenuation change easy to detect (Crampin, 2001; Silver et al., 2007). In Frio II seismic monitoring test, despite having noisy data due to fluid sampling, massive stacking of continuous waveforms enabled clear detection of 2–8 % changes in travel time every 15 mins due to CO₂ plume movement (Daley et al., 2007). The benefit of stacking and good background VSP data to identify plume movement was also iterated in the Citronelle Alabama CO₂ storage demonstration (Trautz et al., 2020). Marchesini et al. (2017) conducted continuous active source monitoring in Cranfield DAS site to measure injection pressure at different depth intervals with high accuracy using 10 tubing-deployed hydrophones and 2 piezoelectric seismic

sources. They concluded that there is strong correlation between travel time and pore pressure with a uncertainty of 26 m in pore pressure distribution across depth. Crosswell seismic monitoring in the Aquistore CO₂ store project reported an even higher accuracy of CO₂ plume movement (5–10 m) and emphasized on velocity and amplitude changes as well as scattering due to plume migration (White et al., 2014). Many related studies have demonstrated excellent resolution in velocity and spectral change stress change and plume migration during CO₂ injection using crosswell active seismic (using geophone or fiber optics) and compared its efficiency with electromagnetic and 3D seismic surveys (Ajo-Franklin et al., 2013; Angerer et al., 2000; Daley et al., 2008a, 2017; Um et al., 2020). Therefore indicators like Vs* and Samp* proposed in this study can be determined even based on small changes in pore pressure and fluid displacement (e.g in the early stage of CO₂ injection, it should be easy to distinguish the magnitude and rate at which the pressure front and fluid front is moving).

To assess the correlation between stress change and acoustic properties more critically across different samples, we determined the Spearman rank correlation across different stages. Spearman rank correlation is a non-parametric measure of the strength and direction of the monotonic relationship between two variables, in this case percentage change in each indicator within each stage (Spearman, 1904). The correlation coefficient varies from -1 to +1 for perfectly negative and perfectly positive correlations respectively. Fig. 11a shows the parameter matrix of Stage B for K15-15A-7 V, and similar matrices were made for each stage and each experiment (Fig. S1). However, the first parameter varies across the stages. For stage B the first parameter is percentage change in axial stress, for stage C and E percentage change in pore pressure. However, in stage D, there was no apparent changes in effective stress, so no stress-term was considered for the ranking. In this stage, liquid CO₂ was injected at a constant flow rate, so percentage change in time was considered as a replacement for stress. In stage B, only the elastic loading part was considered, as the linear relationship between stress and acoustic indicators reverse after the onset of plasticity. Similarly, in stage E, data beyond the point of exponential strain increase is not considered for ranking. Fig. 11b indicates the average Spearman correlation rank determined within each stage (e.g the absolute average of all ranks in Fig. 11a). Any rank above 0.6 (dotted line) is assumed to be a good correlation and vice versa. We can see a consistent high rank in stage B, where acoustic indicators are only a function of axial loading. Except K15-15A-8 V and K15-12-5 V, all samples show a high rank in stage C and D and interestingly these two

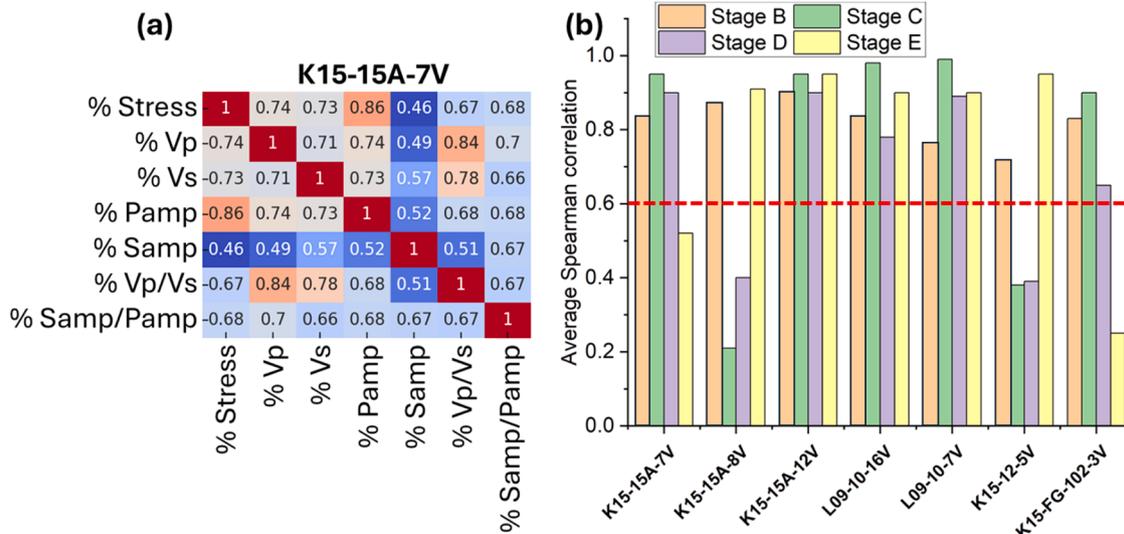


Fig. 11. (a) Spearman rank correlation of stress change and acoustic parameters in stage B for K15-15A-7 V. Stronger correlation is indicated by warmer colors and vice versa. (b) Average Spearman rank correlation for all stages in all experiments with the dotted line acting as a threshold for categorizing good and bad correlation.

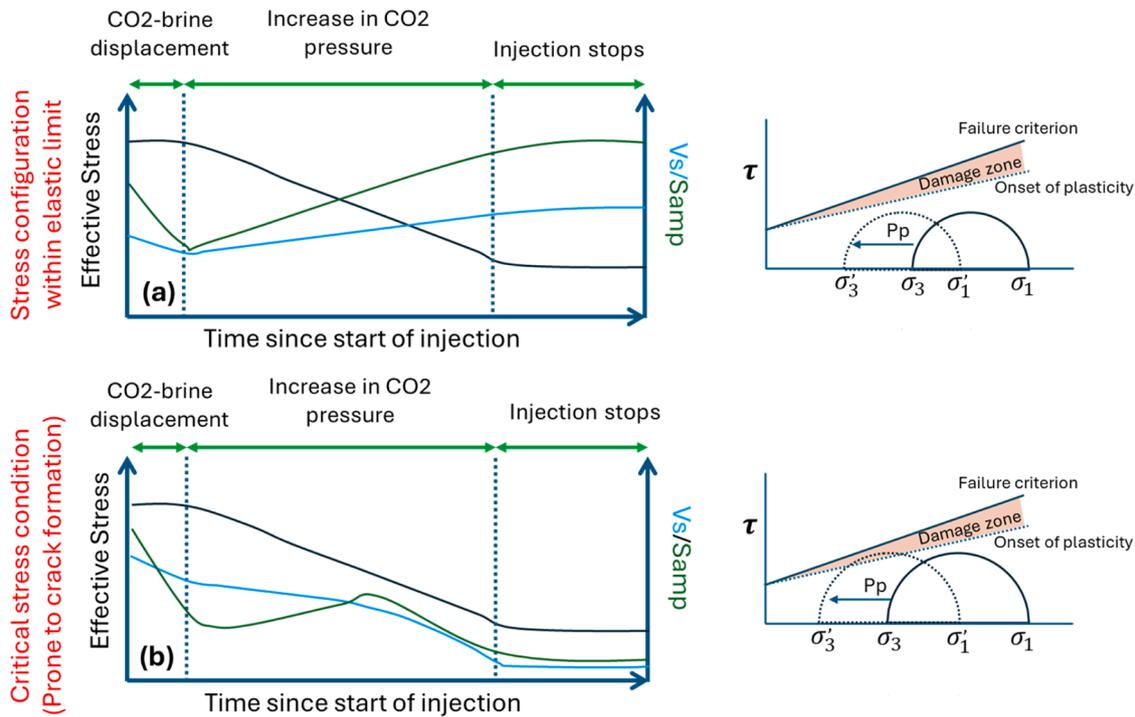


Fig. 12. A simple schematic of expected shear wave velocity and amplitude change during different timeframe of CO₂ injection in a (a) reservoir within elastic limit and (b) reservoir in critical stress condition. The Mohr circles on the right represent the stress configuration in each case.

are samples with the least initial permeability and comparatively low porosity (Table 1).

Based on our observations, the reservoir stress configuration, Vs and Samp can be linked and can act as strong monitoring tool for assessing safe CO₂ injection (Fig. 12). Having a clear idea about the reservoir stress configuration is crucial before making such assessments. In section 4.5 we also discussed how observations from seismic frequency and ultrasonic frequency are similar in many cases, albeit of different magnitude. Therefore, we present the monitoring protocol focusing more on the trend rather than the absolute values. Assuming the monitoring is taking place close to the injection well, and considering the reservoir stress conditions are well within the elastic limit, brine displacement by CO₂ in the starting phase of injection will cause minimal increase in pore pressure, but significant decrease in Samp and a gentler decrease in Vs (Fig. 12a). In a depleted saline aquifer, where CO₂ is stored in supercritical form, the drop in Samp and Vs will be greater during displacement of brine with CO₂ as density of liquid CO₂ > supercritical CO₂. A similar observation has been reported by Zhao et al. (2025) where they investigated the effect of fluid displacement and pore pressure on ultrasonic wave velocities and amplitudes using supercritical CO₂. After some time, depending on injection rate and permeability, pore pressure will start to increase, causing increase in CO₂ density, and a decrease in effective stresses- resulting in an increase in velocity and faster increase in amplitude. When we compare this to a critically stressed reservoir (i.e prone to irreversible deformation), the stress-velocity-amplitude correlations are a little more complicated (Fig. 12b). In the CO₂-brine displacement phase, Vs-Samp behaves similar to Fig. 12a, since the stress change is not applicable in that phase. However, when pore pressure starts increasing, the Mohr-Coulomb circle shifts more towards the failure criterion, going into the damage zone (characterized by the onset of plastic deformation) leading to more microcrack formation in the reservoir. This causes a slow decrease in Vs, however, Samp increases slowly initially due to the increase in density of pore fluid. After a certain point, when the rate of crack formation increases significantly, both Vs and Samp decreases at a faster rate. Compressibility of the pore fluids also significantly contributes to the

density and therefore the amplitudes. Although we did not measure the compressibility of the brine used for these experiments, we know that water has negligible change in isothermal compressibility ($\beta_{\text{water}} = 0.005 \text{ MPa}^{-1}$) across the experimental pressure range (Fine and Millero, 1973). Therefore, the brine used for this study is assumed to have a constant density with pore pressure. Liquid CO₂ on the other hand, has a non-linear decrease in isothermal compressibility ($\beta_{\text{liquid CO}_2} = \frac{0.01}{P^{0.6}}$) within experimental pore pressure range (Aakre et al., 2021) (Fig. S2) where P is pore pressure. In a reservoir, assuming a 90°C isothermal condition, supercritical CO₂ will show a similar trend in compressibility ($\beta_{\text{sc CO}_2} = \frac{0.015}{P^{0.75}}$), albeit with lower density (Aakre et al., 2021; Romei and Persico, 2021). Therefore, the increase in density of CO₂ is rapid at lower pore pressure, counterbalancing the possible decrease in amplitudes due to microcrack formation. At higher pore pressure, the increase in density is slower, therefore the impact of crack volume increase is more pronounced. The impact of density will be lower when the injected phase is brine/water (for geothermal projects) and more pronounced for compressible fluids like H₂.

5. Conclusions

The objective of this study was to demonstrate the applicability of compressional and shear waves for stress and plume migration monitoring in CO₂ injection operations, where at different stages of the project lifetime, stress change or fluid displacement or a both their coupled occurrence could be seen. However, more lab-scale and field-scale research needs be conducted on reservoirs having varying mineralogy and petrophysical properties, in order to improve confidence in the extent to which the results are representative. A critical assessment of differences in scale (reservoir instead of coreplugs), heterogeneity and wave attributes (seismic instead of ultrasonic) can help establish a link between experimental outcome and real-time monitoring. Based on our lab-scale experiments using ultrasonic wave attributes, we could derive the following key points which could be applicable for field scale monitoring:

1. Ultrasonic velocity and amplitudes increase linearly during elastic compaction. The rate of increase of shear wave velocity is higher compared to compressional waves, however, the trends in amplitude are opposite.
2. Fluid injection (imbibition of wetting and non-wetting fluids) causes pulsed increase (during pore pressure increase) or decrease (during CO₂-brine displacement) in ultrasonic velocity and amplitudes. This is caused by a combination of difference in capillary number, fluid wettability, porosity, pore connectivity and heterogeneity of the specimen. In low-frequency measurements, phenomena like wave induced flow and heterogeneities in the reservoir needs to be accounted for to calculate saturation from velocity change.
3. An increase in pore pressure under critically stressed conditions promotes crack formation, leading to a gradual decline in both P-wave (V_p) and S-wave (V_s) velocities, along with an exponential reduction in P-wave amplitude. In contrast, S-wave amplitude either remains constant or increases until reaching the failure threshold, after which it decreases sharply. Pore pressure changes in critical stress condition continuously changes the patch size and permeability due to pore collapse and crack formation, which needs to be studied in detail to upscale ultrasonic measurements to seismic scale.
4. The relative rate of increase in wave velocity and amplitude serves as a key indicator for distinguishing different scenarios, including plastic deformation independent of pore pressure, pore pressure-induced effective stress changes, CO₂-brine displacement, and pore pressure-induced plastic deformation. Monitoring these indicators over time enables operators to evaluate the reservoir's stress state and safety factor through cross-well or downhole seismic measurements. Since seismic waves are sensitive to microcracks and their directional anisotropy, velocity change due to inelastic deformation would be identifiable, however the exact magnitude of change would vary from our ultrasonic measurement and still needs to be investigated.
5. Applicability of this study extends beyond the scope of fluid injection operation, including subsurface CO₂, H₂ storage or geothermal applications, provided good background (before CO₂ injection) active seismic measurements are performed, and should be accompanied with other forms of monitoring (borehole pressure, distributed strain, geochemical) to achieve improved confidence on shape and extent of injected fluid plume.

CRedit authorship contribution statement

Debanjan Chandra: Writing – original draft, Visualization, Validation, Methodology, Investigation, Formal analysis, Data curation, Conceptualization. **Auke Barnhoorn:** Writing – review & editing, Validation, Supervision, Resources, Project administration, Funding acquisition.

Declaration of competing interest

The authors hereby declare no conflict of interest and that the work described has not been published previously, nor it is under consideration for publication elsewhere. The article's publication is approved by all authors before submission.

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Supplementary materials

Supplementary material associated with this article can be found, in the online version, at [doi:10.1016/j.ijggc.2025.104498](https://doi.org/10.1016/j.ijggc.2025.104498).

Data availability

Data will be made available on request.

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