



**Quantification of the
influence of
preferential flow on
slope stability**

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This discussion paper is/has been under review for the journal Hydrology and Earth System Sciences (HESS). Please refer to the corresponding final paper in HESS if available.

Quantification of the influence of preferential flow on slope stability using a numerical modeling approach

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Received: 7 October 2014 – Accepted: 13 October 2014 – Published: 26 November 2014

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Published by Copernicus Publications on behalf of the European Geosciences Union.

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Abstract

The effect of preferential flow on the stability of landslides is studied through numerical simulation of two types of rainfall events on a hypothetical hillslope. A model is developed that consists of two parts. The first part is a model for combined saturated/unsaturated subsurface flow and is used to compute the spatial and temporal water pressure response to rainfall. Preferential flow is simulated with a dual-permeability continuum model consisting of a matrix domain coupled to a preferential flow domain. The second part is a soil mechanics model and is used to compute the spatial and temporal distribution of the local factor of safety based on the water pressure distribution computed with the subsurface flow model. Two types of rainfall events were considered: long duration, low-intensity rainfall, and short duration, high-intensity rainfall. The effect of preferential flow on slope stability is assessed through comparison of the failure area when subsurface flow is simulated with the dual-permeability model as compared to a single-permeability model (no preferential flow). For the low-intensity rainfall case, preferential flow has a positive effect on the slope stability as it drains the water from the matrix domain resulting in a smaller failure area. For the high-intensity rainfall case, preferential flow has a negative effect on the slope stability as the majority of rainfall infiltrates into the preferential flow domain when rainfall intensity exceeds the infiltration capacity of the matrix domain, resulting in larger water pressure and a larger failure area.

1 Introduction

Landslides are commonly triggered by rainfall events. Hydrological models may be integrated with slope stability analysis methods to calculate the factor of safety and predict the time and magnitude of landslides (Crosta and Frattini, 2008; Shuin et al., 2012; Aleotti and Chowdhury, 1999; Westen et al., 2006). Combined hydro-mechanical models can roughly be divided into two types: simplified conceptual models (Montra-

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but can not effectively simulate preferential flow resulting in rapid infiltration (Nieber and Sidle, 2010; Beven and Germann, 2013).

In highly heterogeneous slopes, preferential flow and transport can fundamentally influence subsurface flow (Jarvis, 2007; Hendrickx and Flury, 2001) and contaminant transport (Köhne et al., 2009; Allaire et al., 2009; Debieche et al., 2012; Zehe et al., 2001). A chain of connected macropores is commonly found in various types of soils, including forest soil and semiarid land (Uchida et al., 2001; Jarvis, 2007; Flury et al., 1994). For example, an earthworm burrow can extend from the surface deep into the soil, as can decayed plant roots or soil cracks (Jarvis, 2007; Beven and Germann, 1982; Hendrickx and Flury, 2001). The self-organizing preferential flow network will become active and hydraulically connected with an increase in soil saturation (Nieber and Sidle, 2010). The saturated hydraulic conductivity of preferential flow paths is significantly larger than that of the soil matrix (Beven and Germann, 1982; Köhne et al., 2009). A significant portion of subsurface stormflow (Uchida et al., 2004; Zhang et al., 2006; Beven, 1981) is transmitted via preferential flow paths (Nieber and Sidle, 2010). Preferential flow through macropores, fractures, and other local high-permeability zones is extremely rapid, and contributes instantly to high pore-water pressures in deep soils (Jarvis, 2007).

Quantification of landslide triggering mechanisms is an essential step in landslide forecasting. Field studies have shown that preferential flow is one of the major mechanisms affecting the timing and location of landslides (Sharma and Nakagawa, 2010). In forested hillslopes, minor preferential flow paths, such as soil pipes and macropores, are clearly associated with slope failure (Hencher, 2010; McDonnell, 1990; Uchida et al., 2001; Krzeminska et al., 2012; Debieche et al., 2012). Besides the fact that internal erosion in preferential flow paths deteriorates the slope mass and reduces the soil shear strength, the occurrence of preferential flow can give rapid access to the deeper soil and groundwater system, reduce soil shear strength (due to pore pressure changes), and influence the timing and frequency of landslides (Köhne et al., 2009; Hendrickx and Flury, 2001).

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Preferential flow and solute transport have been simulated at various scales including the scales of pores, soil columns, hillslopes, and catchments (Šimůnek et al., 2003; Gerke, 2006; Köhne et al., 2009) using increasingly sophisticated models such as the dual-porosity/dual-permeability model (Gerke and van Genuchten, 1993a; Jarvis et al., 1991; Larsbo and Jarvis, 2003), the multi-permeability model (Wu et al., 2004; Greco, 2002; Gwo et al., 1995), and the empirical model (Armstrong et al., 2000; Weiler, 2005; Vrugt et al., 2004; Mulungu et al., 2005). The dual-permeability model is widely used because of its clear physical concept and powerful simulating ability (Roulier and Jarvis, 2003; Kodešová et al., 2005; Gerke and Köhne, 2004; Köhne et al., 2006; Christiansen et al., 2004; Weiler, 2005; Therrien and Sudicky, 2005; Vogel et al., 2000). The dual-permeability model assumes that the soil consists of two interacting, overlapping pore domains. The matrix domain with relatively low permeability represents the soil micropores where flow is governed by Richards' equation. The preferential flow domain represents the highly permeable preferential flow paths, such as macropores, fractures, cracks, or large pores between soil aggregate. Preferential flow is described by Richards' equation (Šimůnek et al., 2008; Gerke and van Genuchten, 1993a) or the gravity-driven kinematic wave equation (Larsbo and Jarvis, 2003; Jarvis et al., 1991; Greco, 2002). The water exchange between the two domains is driven by the pressure head difference between the two domains (Pirastru and Niedda, 2010; Gerke and van Genuchten, 1993b). Dual-permeability models have proven to be effective for preferential flow simulation, but have not been incorporated into slope stability models.

The objective of this study is to quantify the temporal and the spatial effect of preferential flow on slope stability, and to analyze its underlying hydrological mechanisms using numerical experiments of rainfall-induced shallow landslides. This paper is organized as follows. First the subsurface dual-permeability hydrological model is described. The subsurface hydrological model is sequentially coupled with a soil mechanics model and a stress-field-based local factor of safety slope stability method (Sect. 2.2). The numerical experiments and parameterization are discussed in Sect. 3. The hydrological and geotechnical results are given in Sect. 4. The influence of

preferential flow on subsurface hydrological processes and consequent slope stability is discussed in Sect. 5 by comparing the results of single and dual permeability models.

2 Methods

2.1 Subsurface flow model

The single-permeability model is described by one Richards' equation to represent flow in a homogenous soil. The dual-permeability model divides the flow domain into two overlapping and interacting continua, where two coupled Richards' equations are used to describe the matrix flow and preferential flow (Gerke and van Genuchten, 1993a):

$$[C_f + \Theta_f S_s] \frac{\partial h_f}{\partial t} = \nabla[K_f(\nabla h_f + \nabla z)] - \frac{\Gamma_w}{w_f} \quad (1)$$

$$[C_m + \Theta_m S_s] \frac{\partial h_m}{\partial t} = \nabla[K_m(\nabla h_m + \nabla z)] + \frac{\Gamma_w}{w_m} \quad (2)$$

where the subscript f indicates the preferential flow domain and the subscript m indicates the matrix domain. C is the differential water capacity ($d\theta/dh$) (L^{-1}), Θ is the effective saturation (-), h is the pressure head (L), t is time (T), z is the vertical coordinate (positive upward), K is the isotropic hydraulic conductivity (LT^{-1}), S_s is the specific storage (L^{-1}), w is the volumetric ratio of the preferential flow domain or the matrix domain over the total soil volume (-), and Γ_w is the water exchange term (T^{-1}) between the two domains.

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The Brooks–Corey function is used to describe the hydraulic properties of both the matrix and preferential flow domains (Brooks and Corey, 1964):

$$\Theta = \frac{\theta - \theta_r}{\theta_s - \theta_r} = \begin{cases} |\alpha_{BC} h|^{n_{BC}}, & \alpha_{BC} h < -1 \\ 1, & \alpha_{BC} h \geq -1 \end{cases} \quad (3)$$

$$K = K_s \Theta^{2/n_{BC} + l_{BC} + 2} = K_s |\alpha_{BC} h|^{-2 - n_{BC}(l_{BC} + 2)} \quad (4)$$

$$C = -\frac{d\theta}{d|h|} = \begin{cases} \alpha_{BC} n_{BC} (\theta_s - \theta_r) |\alpha_{BC} h|^{-n_{BC}-1}, & \alpha_{BC} h < -1 \\ 0, & \alpha_{BC} h \geq -1 \end{cases} \quad (5)$$

where θ is the water content ($L^3 L^{-3}$), subscripts s and r denote saturation and residual state, K_s is the saturated hydraulic conductivity (LT^{-1}), and α_{BC} , l_{BC} , n_{BC} , are fitting parameters.

Γ_w is the water exchange rate between the two domains (Ray et al., 1997):

$$\Gamma_w = \alpha_w K_a (h_f - h_m) \quad (6)$$

where α_w (L^{-2}) is the effective water transfer coefficient, and the relative hydraulic conductivity K_a (LT^{-1}) is calculated by averaging the hydraulic conductivities of the two pore domains (Arora et al., 2011; Laine-Kaulio et al., 2014):

$$K_a = \frac{K_f + K_m}{2}. \quad (7)$$

The volumetric ratio of the preferential flow domain and matrix domain sum up to one:

$$w_f + w_m = 1. \quad (8)$$

The total water content of the soil is the weighted average of the water contents of the two domains:

$$\theta = w_f \theta_f + w_m \theta_m. \quad (9)$$

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The same holds for the total saturated hydraulic conductivity of the soil:

$$K_s = w_f K_{sf} + w_m K_{sm}. \quad (10)$$

Boundary conditions may be specified pressure head, specified flux, or mixed (Chui and Freyberg, 2009). In the case of a dual-permeability model, specified flux i (infiltration from rainfall) is divided between the matrix and preferential flow domains:

$$i = w_f i_f + w_m i_m \quad (11)$$

where i_m and i_f are boundary fluxes to the matrix and the preferential flow domains (LT^{-1}), respectively. The two domains have an equal opportunity to receive rainfall and are initially equal to rainfall intensity R (Dusek et al., 2008):

$$R = i = i_f = i_m. \quad (12)$$

As the matrix domain has a larger volumetric ratio ($w_m > w_f$), the infiltration process is initially dominated by the matrix domain. Once the specified flux into the matrix is larger than its infiltration capacity, the boundary condition changes to specified pressure head and the specified flux for the preferential flow domain is increased to:

$$i_f = \frac{R - w_m i_m}{w_f}. \quad (13)$$

Once the specified flux into the preferential flow domain is also larger than the infiltration capacity, the boundary conditions of both the matrix and the preferential flow domain are changed to a specified pressure head of zero and overland flow occurs.

2.2 Slope stability analysis method

The slope stability analysis is based on the local factor of safety approach (Lu et al., 2012). The plane-strain linear elasticity model is used to calculate the stress (Abramson, 2002), which is governed by a momentum balance equation:

$$\nabla(\sigma) + \gamma b = 0 \quad (14)$$

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where σ is a stress tensor ($\text{ML}^{-1}\text{T}^{-2}$) with three independent stress variables in two-dimensional space, γ is the bulk unit weight of the slope material ($\text{ML}^{-2}\text{T}^{-2}$), and \mathbf{b} is the unit vector of body forces with two components. Neglecting the relative air pressure in Bishop's equation, the effective stress equation is:

$$5 \quad \sigma' = \sigma - \chi \rho_w \quad (15)$$

where σ' is the effective stress, p_w ($\text{ML}^{-1}\text{T}^{-2}$) is the pore water pressure, and χ (–) is the matrix suction coefficient, which is usually approximated by the effective saturation (Lu et al., 2010).

10 The local factor of safety F_{LFS} is defined as the “ratio of the Coulomb stress at the current state of stress to the Coulomb stress of the potential failure state under the Mohr–Coulomb criterion” (Lu et al., 2012):

$$F_{\text{LFS}} = \frac{\tau^*}{\tau} \quad (16)$$

where τ^* is the limit Coulomb stress and τ is the actual shear stress ($\text{ML}^{-1}\text{T}^{-2}$). Application of the Mohr–Coulomb failure criterion gives:

$$15 \quad F_{\text{LFS}} = \frac{2 \cos \phi'}{\sigma'_1 - \sigma'_3} \left[c' + \frac{\sigma'_1 + \sigma'_3}{2} \tan \phi' \right] \quad (17)$$

where c' is the effective cohesion ($\text{ML}^{-1}\text{T}^{-2}$), ϕ' is the friction angle, σ'_1 and σ'_3 are the first and the third effective stress for the variably saturated soil ($\text{ML}^{-1}\text{T}^{-2}$).

20 The influence of hydrology on slope stability is manifested in two ways. First, the unit weight function depends on the water content (Eq. 9). Second, the effective stress depends on the pore water pressure. In the dual-permeability model, the pore water pressure of the preferential flow domain is used in the computation of the effective stress.

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Figure 1 summarizes the structure of coupled dual-permeability and slope stability model. Two Richards' equations are coupled by the water exchange function. The hydrological results are sequentially coupled with a soil mechanics model without considering possible feedback of soil deformation on soil properties and the hydrological process.

3 Setup of the numerical experiments

3.1 Slope geometry

Consider a slope of 23° consisting of fine-grained lithology such as clay shales with a more permeable weathered top soil layer (Bogaard, 2002; Berti and Simoni, 2012; Picarelli et al., 2006); this is a typical slope that is vulnerable to failure. The slope is 6 m high and 15 m long and consists of two layers with a 2 m thick homogeneous upper soil layer (see Fig. 2).

The model domain is 42 m by 25 m to reduce the influence of boundary effects on hydrological and slope stability results. The computational mesh and the boundary conditions are shown in Fig. 2. The boundary conditions of the subsurface flow model are atmospheric at the surface; the left-hand and bottom sides are no-flux boundaries. The right-hand side consists of a seepage boundary condition for the upper soil layer and a specified pressure head to mimic a constant groundwater table for the lower layer. For the soil mechanics model, the surface is a free boundary, the bottom boundary (only horizontal displacements) and the left- and right-hand sides (only vertical displacements) are all roller boundaries.

Since the pressure head in the surface area can change drastically during rainfall, a very dense mesh was used near the surface to accurately model the transient hydrological conditions. The mesh density of the upper layer is approximately 0.25 m (vertical) by 0.5 m (horizontal). A coarser mesh was defined in the lower part of the slope as a less dynamic condition will occur here.

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3.2 Parameterization

The volumetric ratio of the preferential flow domain w_f is 0.1; a typical range is 0.025 ~ 0.2 (Köhne et al., 2002). The pore-size distribution of the preferential flow domain allows unsaturated infiltration before the matrix domain is saturated (Dusek et al., 2008). A comparison is made between the hydrological results of the single-permeability and the dual-permeability models. The total weighted saturated hydraulic conductivity of the dual-permeability model is equal to the saturated hydraulic conductivity of the single-permeability model. The water exchange between the matrix and preferential flow domains depends on the hydraulic conductivity between the two domains K_a and the water exchange coefficient α_w (Eq. 6). Equilibrium between the preferential flow and matrix domains is reached quicker for smaller values (closer to 1) of K_{sf}/K_{sm} and larger values of α_w . Moderate values are used for K_{sf}/K_{sm} (100 in the upper layer and 5 in the lower layer) and for α_w (0.2 m^{-2}).

The soil hydraulic parameters are presented in Table 1. Preferential flow plays an important role in the upper soil layer where there is an abundance of macropores, but less so in the lower soil layer where macropores are almost non-existent (Bogner et al., 2013). In other words, the volumetric percentage of preferential flow domain is still the same, but in the lower layer the saturated hydraulic conductivity of macropores are more similar to the pores of the matrix. The more permeable top layer is sandy loam and the fine-grained lower layer is clay; the soil hydraulic parameters are taken from the UNSODA database (Nemes et al., 2001; Leij, 1996).

Current laboratory practice for soil hydraulic testing cannot measure the parameters for two hydraulic functions and two water retention curves for one soil sample (Arora et al., 2011; Köhne et al., 2009). There are two approaches to parameterize a dual-permeability model. The first approach determines the parameters from an infiltration experiment and inverse modeling, which results in a non-unique parameter set (Dusek et al., 2008; Köhne et al., 2002; Arora et al., 2011). The second approach, which is

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part of the rainfall turns into overland flow. For the single-permeability model and low-intensity rainfall overland flow starts after 95 h (or 190 mm of rainfall) while for the high intensity rainfall overland flow starts after 8.5 h (or 170 mm of rainfall) (Fig. 4a and b).

In the dual-permeability model (Fig. 5), the rainfall infiltration is divided over the two domains and additional rainfall infiltrates into the preferential flow domain when the matrix domain reaches infiltration capacity. Recall that the matrix domains is 90 % of the domain, and the preferential flow domain is 10 % of the domain. A smaller fraction of rainfall infiltrates into the preferential flow domain for the case of low-intensity rainfall (10–15 %) than for the case of high-intensity rainfall (50–85 %). Overland flow starts after 80 h (or 160 mm of rainfall) for the low-intensity case and after 60 h (or 120 mm of rainfall) for the high-intensity case.

The seepage outflow increases along all three boundaries during the rainfall event (Figs. 4c, d, and 5c, d) and is smaller than the infiltration rate (storage is increasing). In the dual-permeability model and the low-intensity rain, outflow along the surface boundary starts after 115 h (or 230 mm of rainfall) while for the high-intensity rain outflow starts after 9 h (or 180 mm rainfall). The outflow rate along the surface boundary depends on the groundwater level in the upper layer. In the dual-permeability model, the outflow along the right boundary is approximately 10 times larger for the preferential flow domain than for the matrix domain, which is consistent with their volumetric ratio and their saturated hydraulic conductivity ratio. The water exchange between the two domains in the dual-permeability model is shown in Fig. 5e and f. For the low-intensity rainfall case, the water exchange from the preferential flow domain to the matrix domain increases during the first 100 h and then decreases, while the water exchange from the matrix domain to the preferential flow domain is almost always increasing (more negative). For the high-intensity rainfall case, the water exchange from the matrix to the preferential flow domain is negligible, while the water exchange from the preferential flow domain to the matrix domain reaches more than $0.3 \text{ m}^2 \text{ h}^{-1}$, which is similar to the infiltration into the preferential flow domain. After five hours, approximately 75 % of

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infiltration into the matrix domain is water exchange from the preferential flow domain (Fig. 5f) and 25 % infiltration from the surface boundary (Fig. 5b).

4.2 Water balance

The integrated rainfall and water storage for the study area are shown for both models in Fig. 6. The water balance is obtained by integrating all flow components along the boundaries of the study area. The numerical water balance errors are between 2 and 3 %.

For all cases, the storage increase flattens out when the inflow decreases (Figs. 4 and 5). For the high-intensity rainfall, the dual-permeability model stores 8 % less water than the single-permeability model. The total storage after 150 h of low-intensity rainfall is less than after 15 h of high-intensity rainfall, probably caused by the longer time that water can drain from the study area under low-intensity rain.

For the dual-permeability model, the water exchange has a significant influence on the storage change in each domain. For the low-intensity rainfall, the storage in the preferential flow domain does not increase much after 6 h (Fig. 6). For the high-intensity rainfall, the storage in the preferential flow domain increases rapidly over the first 3 h as very little water infiltrates into the matrix domain due to the low infiltration capacity of the matrix. After 3 h, the preferential flow domain has almost reached full saturation and the large pressure difference between the preferential flow domain and matrix domain causes extensive water exchange (Fig. 5f).

4.3 Water content

The water content distribution in the study area is shown in Fig. 7 for both the single-permeability model (left-hand panels) and the dual-permeability model (center and right-hand panels). The water exchange rate between the matrix and preferential flow domains of the dual-permeability model is shown in Fig. 8. The infiltration process

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of the dual-permeability model differs significantly from that of the single-permeability model.

The initial water content distribution in the matrix and preferential flow domains is similar for both models. During the rainfall events, the wetting front in the single-permeability model develops parallel to the surface and propagates downward. This holds for both low and high rainfall intensities (Fig. 7 left-hand column). The wetting front generally reaches the groundwater table at the toe of the slope first, after which the infiltrated water continuously enlarges the saturated area.

In the dual-permeability model, the combined effects of the preferential flow and the matrix flow show a more complicated response. For the low-intensity rainfall, infiltration is dominated by matrix flow, as 90 % of the subsurface consists of the matrix. Because the rainfall intensity is lower than the saturated conductivity of the matrix domain, rainfall never exceeds infiltration capacity (Fig. 5a), so that 90 % of the rainfall infiltrates into the matrix domain and 10 % of the rainfall infiltrates into the preferential flow domain. The pressure are different between domains, that directly cause the water exchange at the matrix wetting front (Figs. 5e and 8a). At first, water quickly reaches the soil layer interface by preferential flow where it transmits to the matrix, although this exchange flux is very small (Figs. 5e and 8a). After sufficient time (70 h), a much stronger matrix flow (taking about 80 % of the infiltrated rainfall) reaches the soil layer interface and generally reverses the water exchange direction (Fig. 5e). Overall, water exchange during low-intensity rainfall in the study area is dominated by flow from the matrix to the preferential flow domain (Fig. 8a and b).

For the high-intensity rainfall, the rainfall intensity is 8.4 times the matrix saturated hydraulic conductivity. The percentage of infiltration into the matrix domain decreases from 90 to 50 % within the first half hour, and continues to decrease to less than 20 % after 1.5 h. In contrast, the percentage of rainfall that infiltrates into the preferential flow domain increases from 10 to over 80 % after 2 h. Water in the preferential flow domain quickly reaches the deeper soil layer and forms a perched groundwater table (Fig. 7), where a significant amount of water infiltrates into the matrix (Fig. 5f).

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4.4 Slope stability

The local factor of safety is computed based on the computed water pressure distribution (Fig. 7). The distribution of the local factor of safety is shown in Fig. 9 for the initial condition and after 150 h (low-intensity rainfall) and 15 h (high-intensity rainfall) for both the single-permeability model and the dual-permeability model and for the case with different cohesion values for the upper and lower layers. The case with equal cohesion values is not shown because the potential failure areas are very small.

A local factor of safety below 1 indicates a potential failure area. The area with a F_{LFS} below 1 was determined every time interval (5 h in case 1, and 0.5 h in case 2) and is shown by the black line in Fig. 9. Slope stability is related to both the specific weight of the wet soil and the pore water pressure in the soil. The specific weight changes due to changes in water storage are relatively small, but changes in water pressure have a significant effect on slope stability, especially in the area of the perched water table.

The size of the potential failure area is plotted vs. the cumulative rainfall in Fig. 10 for the two different rainfall events and two sets of cohesion values. The results for the same cohesion values ($c'_1 = c'_2 = 5$ kPa) are shown in Fig. 10a. For the low-intensity rainfall, the failure area is very small and is approximately the same for both permeability models. For the high-intensity rainfall, the failure area in the single-permeability model is larger than for the low-intensity rainfall, but the trend is similar. The failure area in the dual-permeability model is significantly larger. Failure starts after 60 mm rainfall, and the failure area continues to grow during the rainfall infiltration process.

The results for different cohesion values ($c'_1 = 3$ kPa, $c'_2 = 6$ kPa) are shown in Fig. 10b. For the low-intensity rainfall, the failure area is 0.7 m^2 in the single-permeability model after 20 mm of cumulative rainfall. The size of this area shows almost no increase until approximately 220 mm of cumulative rainfall, when the ground-water table starts to rise (Fig. 7). The failure area of the dual-permeability model is 40 % smaller than that of the single-permeability model as the preferential flow domain drains more water into the matrix domain. For the high-intensity rainfall, the failure area of the

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dual-permeability model is larger than of the single-permeability model, as for the case with equal cohesion values. The failure areas of both models increase fairly quickly to 2 m^2 , or 5% of the upper layer in the study area. The failure area increases to 5 m^2 in the dual-permeability model and to 3 m^2 in the single-permeability model.

The slope stability result are directly related with subsurface hydrological results. For the low-intensity rainfall, the failure area for the single-permeability model is very similar in size and location to the dual-permeability model as the location of the water table is very similar in both models (Fig. 7). The initial condition of the dual-permeability model is slightly more stable than that of the single-permeability model, since the preferential flow domain has a higher drainage capacity and, consequently, a lower pore water pressure. In the case of low-intensity rainfall, the matrix flow dominates the ground-water recharge and, consequently, the slope instability. Furthermore, the pore water pressure in the preferential flow domain is very low due to its strong drainage capacity. As a result, the failure area calculated by the dual-permeability model under low-intensity rainfall is slightly smaller than that calculated by the single-permeability model (Fig. 10a). The location of the failure area is similar in the single- and the dual-permeability domain (Fig. 9).

For the high-intensity rainfall, the failure area is significantly larger for the dual-permeability model than for the single-permeability model as the perched water table in the preferential flow domain is much more extensive in the dual-permeability model as compared to the single-permeability model (Fig. 7). The regular wetting front of the single-permeability model does not reach the interface between soil layers, and the failure area is limited to the toe of the slope. For the dual-permeability model, the high-intensity rainfall results in a rapid infiltration through preferential flow, which quickly reaches the interface between soil layers, and increases the degree of saturation and pressure head of the deeper soil. Positive pore water pressure occurs in the preferential flow domain before the entire slope is fully saturated, and produces a larger failure area than in the equivalent single-permeability model.

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The role of preferential flow in hydrology focuses mainly on the rapid vertical infiltration of water and contaminant (Christiansen et al., 2004; Kodešová et al., 2005; Laine-Kaulio et al., 2014), or the rapid discharge in hillslope and catchment hydrological studies on discharge generation (Zhang et al., 2006; Mulungu et al., 2005). A physically based numerical model can be used to investigate the hydrologic response under predefined conditions. Prior to this study, a systematic evaluation of the influence of preferential flow on slope stability using a fully coupled dual-permeability and slope stability model has not been carried out. In this section, the underlying approximations of the numerical model are explored and the influence of the chosen parameter sets on the model outcome is discussed. The numerical experimental results are compared with field studies and other published numerical experiments.

5.1 Continuum model

Soil heterogeneity is one of the most difficult problems in both hydrology and soil mechanics studies. As an alternative to the continuum approach used here, preferential flow may be simulated by explicitly including fissures, pipes, or fracture networks in discrete (or discontinuous) model. Several field studies (Hencher, 2010) and numerical experiments (Tsutsumi and Fujita, 2008; Chang et al., 2014) have focused on the investigation and simulation of pipe flow (in soil) and fracture flow (in rock). In order to accurately describe the geometry of the preferential flow paths, the high-resolution macropore image reconstruction approach (Hu et al., 2014) or the statistical approach (Köhne et al., 2009) may be applied. Numerical simulation of these natural macropore networks require large amounts of geometry information (Nieber and Sidle, 2010) and computational time and are consequently limited to small-scale studies with a limited number of pipes (Tsutsumi and Fujita, 2008) or cracks (Moonen et al., 2008).

The dual-permeability model is a useful tool to simulate subsurface stormflow and solute transport in a forested hillslope when the parameterization is able to capture

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the hydraulic characteristics of each domain (Laine-Kaulio, 2011; Laine-Kaulio et al., 2014). As the dual-permeability model describes the subsurface as a continuum of two linked domains, it is suitable for heterogeneous slopes with a high density of preferential flow paths and not for slopes with only a few large fissures or cracks (van der Spek et al., 2013).

In this paper, flow in both domains is described with the Darcy–Richards' equation, which is valid when the macropores have a relatively small size, and the macropore flow is still viscous (Köhne and Mohanty, 2005; Laine-Kaulio et al., 2014). When fluid velocities are high and flow becomes turbulent, Darcy's equation is not valid (Beven and Germann, 2013) as may be the case in large cracks or fissures under near-saturated or ponded infiltration (Beven and Germann, 1982). The existence of pore necks and dead ends in preferential flow paths reduce the occurrence of turbulent flow (Jarvis, 2007).

5.2 Coupling term in dual-permeability model

In the dual-permeability model, the two domains are in general not at equilibrium. The water exchange is governed by two parameters: the water exchange coefficient and the average hydraulic conductivity between the two domains (Eq. 6). The average hydraulic conductivity in turn is a function of the hydraulic conductivities of the two domains, which are a function of the pressure head. The larger the product, the quicker the two domains equilibrate. Estimation of the water exchange coefficient from physical measurements is very difficult. The most widely used equation is (Gerke and van Genuchten, 1993b):

$$\alpha_w = \frac{\beta}{d^2} \gamma_w \quad (18)$$

where β is a scaling factor, d is half the representative distance between two macropores, and γ_w is a geometry-dependent shape factor that equals 3 for rectangular slabs and 15 for spheres (Ray et al., 1997). Parameter values for the water exchange term

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model. A 23° slope consisting of two soil layers was used in the study. The upper layer is sandy loam and the bottom layer is clay. Both the case where the cohesion of the two layers are equal, and the case where the cohesion of the upper layer is smaller than the lower layer were simulated. Two types of rainfall events were considered low-intensity, long duration rainfall, and high-intensity short duration rainfall; the total amount of water of both rainfall events was equal. The effect of preferential flow on slope stability was studied by comparing the failure area obtained with a single-permeability model and a dual-permeability model for the same rainfall event.

For low intensity rainfall, the failure area of both models is similar when the cohesion of the upper and lower layers is equal, but the failure area is significantly larger in the single-permeability model as compared to the dual-permeability model when the cohesion of the upper layer is lower than the cohesion of the lower layer. During low intensity rainfall, preferential flow has a positive effect on slope stability as it drains water from the matrix domain and decreases the water pressure.

For high intensity rainfall, the failure area of the dual-permeability model is significantly larger than the single-permeability model whether the cohesion values of the two layers are equal or not. During high intensity rainfall, the rainfall intensity is larger than the infiltration capacity of the matrix domain so that most of the rainfall infiltrates into the preferential flow domain. As a result, the water pressure increases very quickly in the preferential flow domain resulting in a much larger failure area than is the case for the single-permeability model.

In summary, the coupled dual-permeability and slope stability model is an effective tool to better understand the influence of preferential flow on slope stability. Preferential flow has a positive effect on slope stability during low intensity rainfall and a negative effect on slope stability during high intensity rainfall.

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Table 2. Parameters setting of water exchange coefficients in different literatures.

Parameter	Vogel et al. (2000) 2-D virtual Numerical study	Gerke and Köhne (2004) Bokhorst clay soil	Köhne and Mohanty (2005) Macropore: coarse sand Matrix: sandy loam	Arora et al. (2011) Multiple-macropore in sandy loam soil	This study
w_f	0.05	0.05	0.009675	0.00033	0.1
K_{sf}/K_{sm}	1000	100	3878	63.6	100
K_{sa}/K_{sm}	0.01	0.001	1	32.1	50.5
K_{sm} (cm h ⁻¹)	0.012	0.1	0.056	0.13	0.238
d (cm)	1	1	11	1.89	–
γ_w	0.4	0.4	0.4	0.001	–
β	3	15	1.0685	0.67	–
α_w (cm ⁻²)	1.2	6	2.5×10^{-3}	1.8×10^{-4}	2.0×10^{-5}
$\alpha_w K_{sa}$ (cm ⁻¹ h ⁻¹)	1.4×10^{-4}	6.0×10^{-4}	2.0×10^{-4}	7.8×10^{-4}	2.4×10^{-4}

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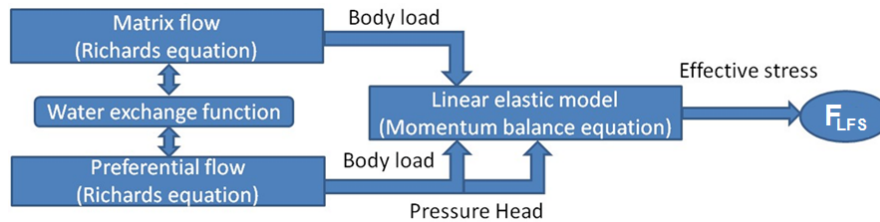


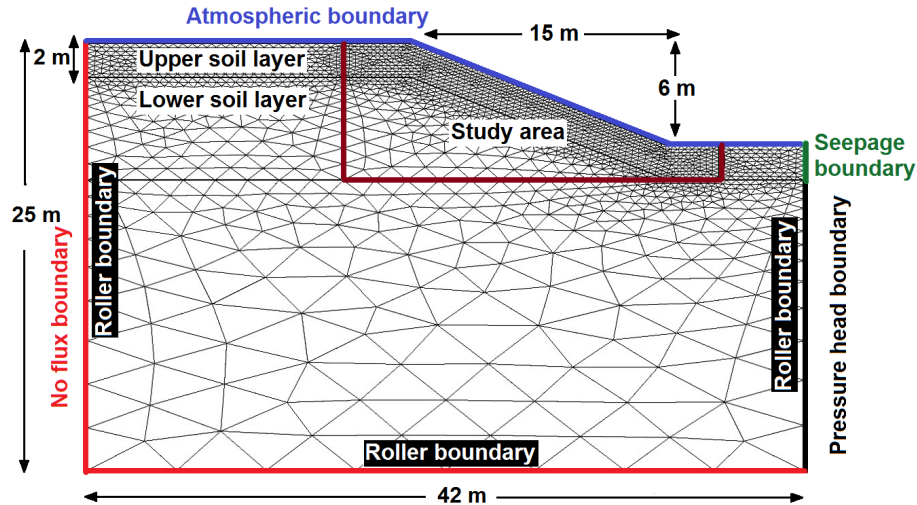
Figure 1. Structure of coupled dual-permeability model and soil mechanics model.

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**Figure 2.** Computational mesh and boundary conditions.[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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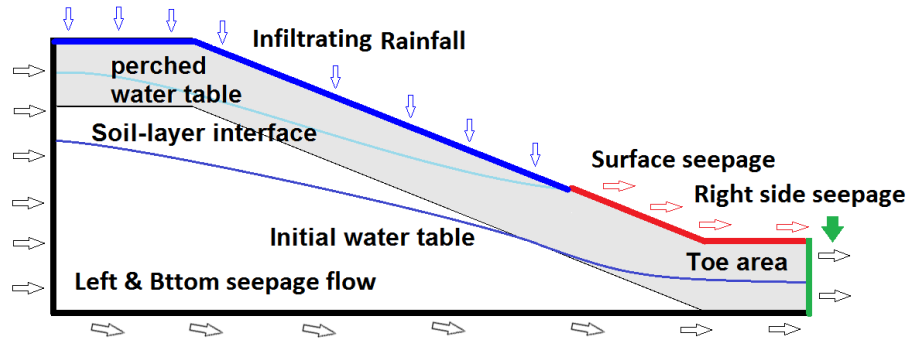


Figure 3. Flow component and water balance of study area.

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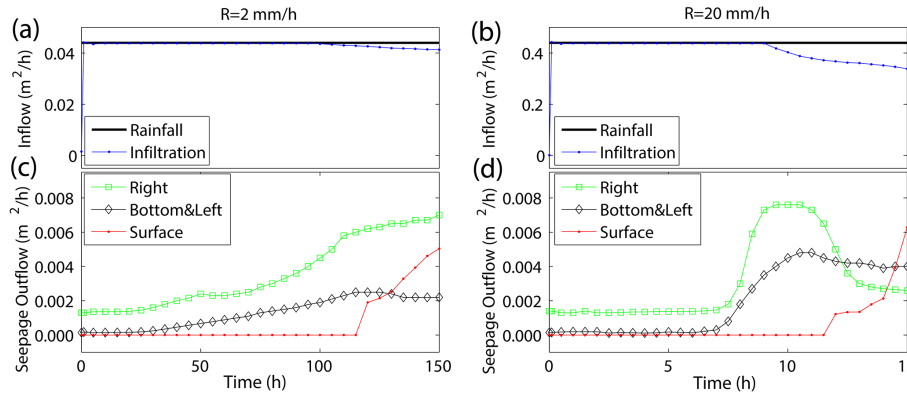


Figure 4. Integrated fluxes for single-permeability model and 2 mm h^{-1} (left panels) and 20 mm h^{-1} (right panels) rainfall. Rainfall and infiltration (a, b), and outflow at the right, outflow at the left and bottom, and outflow at the surface (c, d).

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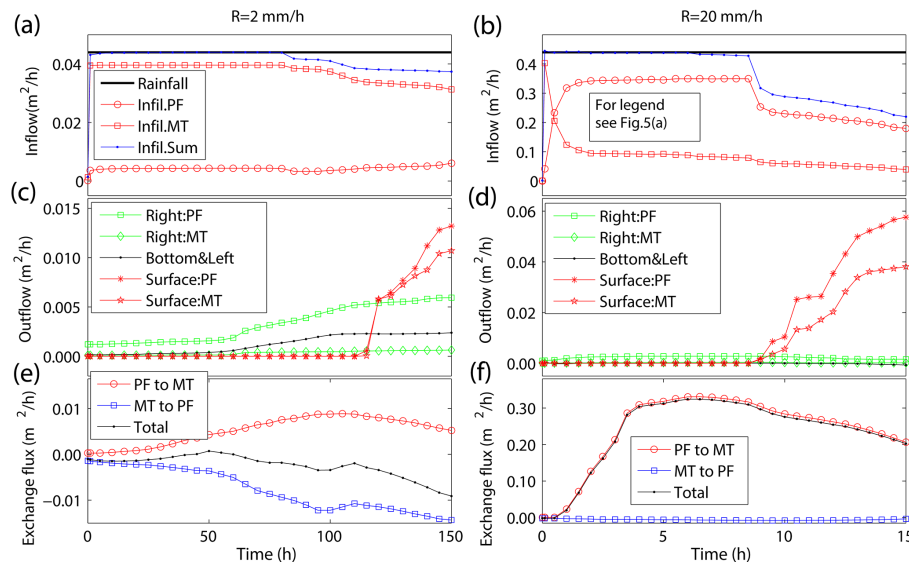


Figure 5. Integrated fluxes for dual-permeability model and 2 mm h^{-1} (left panels) and 20 mm h^{-1} (right panels) rainfall. Rainfall and infiltration (**a**, **b**), and outflow at the right, outflow at the left and bottom, and outflow at the surface (**c**, **d**), exchange between matrix domain (MT) and preferential flow domain (PF) (**e**, **f**) positive for flow from PF to MT and negative for flow from MT to PF.

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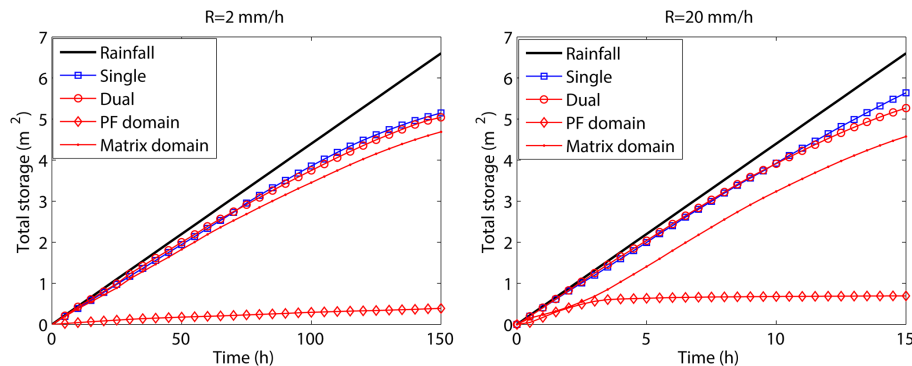
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**Figure 6.** Storage increase of single-permeability model and dual-permeability model.[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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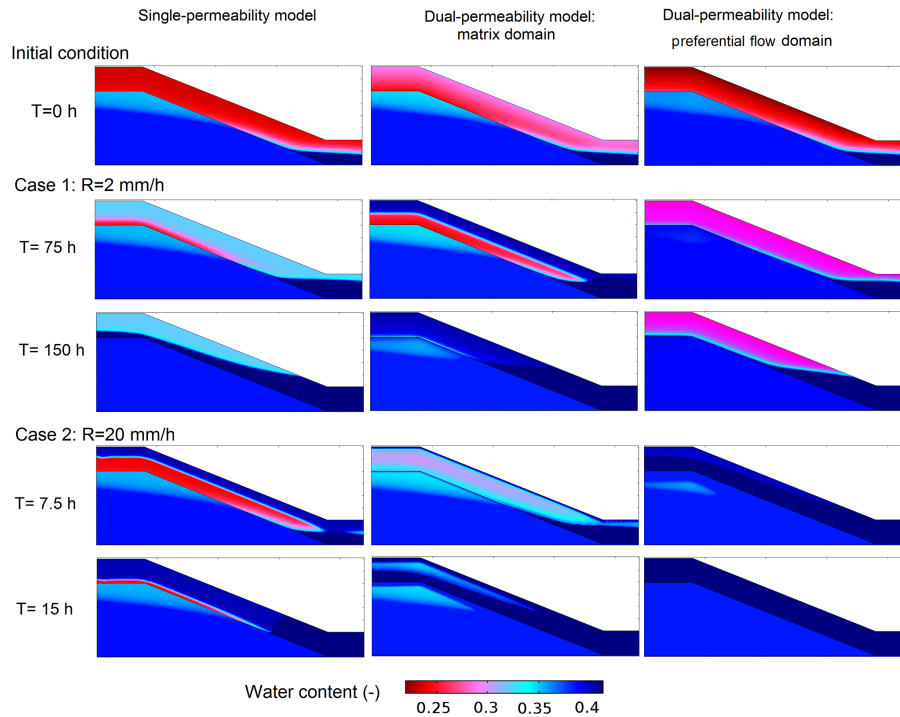


Figure 7. Water content distribution.

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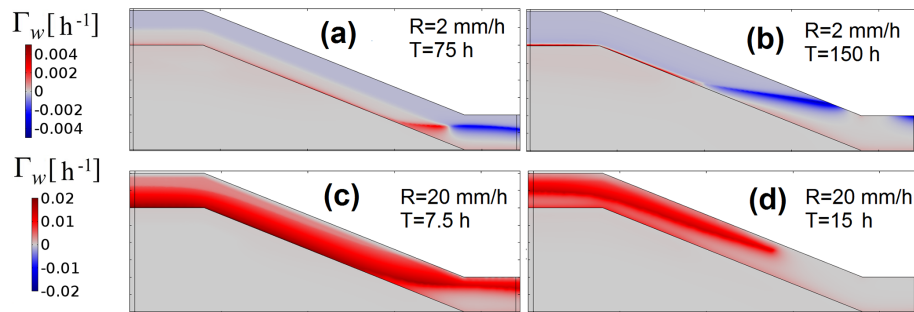


Figure 8. Water exchange rate distribution. Positive values (red) mean water exchange from preferential flow domain to matrix, negative values (blue) mean water exchange from matrix to preferential flow domain.

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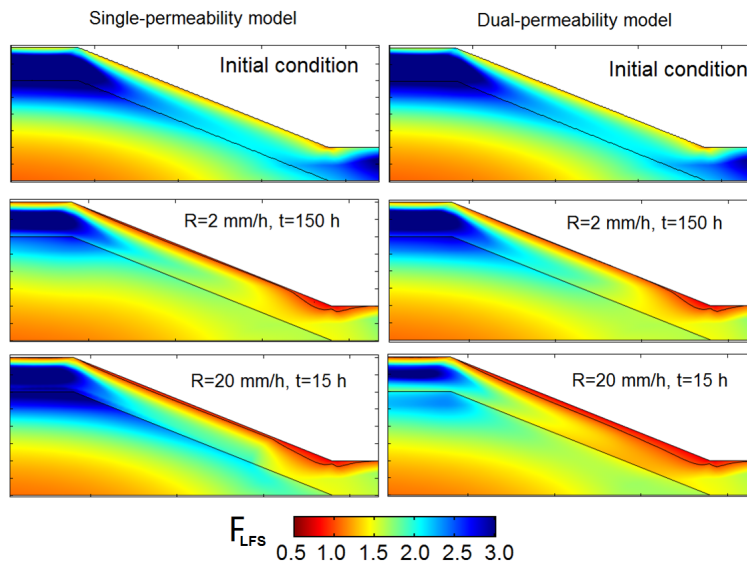


Figure 9. Final slope stability after the rainfall event ($c'_1 = 3$ kPa, $c'_2 = 6$ kPa). The black line delineates potential unstable area ($F_{LFS} < 1$).

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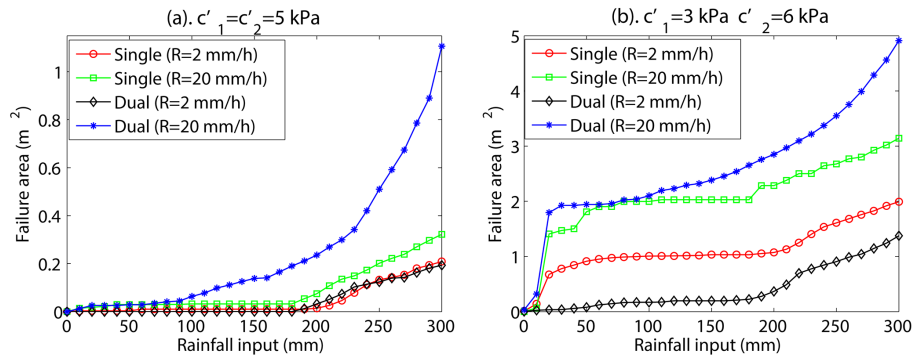


Figure 10. Development of the failure area under different rainfall intensities and soil cohesions.

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