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Turning One-sided Illumination into Two-sided Illumination by Target-enclosing Interferometric Redatuming

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SUMMARY

We present a novel method to transform seismic data with sources at the surface and receivers above and below a selected target zone in the subsurface into virtual data with sources and receivers located at the initial receiver locations. The method is based on inverting a series of multidimensional equations of the convolution- and the correlation-type. The required input data can be computed from surface seismic data with a new iterative scheme that is currently being developed. The output data contains virtual sources that illuminate the target not only from above (as in the original data), but also from below, facilitating the needs of seismic imaging and inversion in an optimal way. The method is nonlinear in the sense that all internal multiples are correctly accounted for and true amplitude in the sense that the virtual sources are forced to inherit uniform radiation patterns even though the overburden is strongly heterogeneous.
Introduction

When seismic data are available with sources at the surface and receivers at a particular level $\Lambda_x$ in the subsurface, sources can be redatumed to this level with interferometric redatuming (IR) by multidimensional deconvolution (Wapenaar et al., 2008; van der Neut, 2012). Unlike conventional redatuming methods (Berryhill, 1984), IR does address both primary and multiple reflections and provides a response as if the medium above $\Lambda_x$ were homogeneous. Recently, it has been shown that Green's functions to any level $\Lambda_x$ can be retrieved from surface seismic data, using a novel iterative scheme (Broggini et al., 2012; Wapenaar et al., 2012). With IR, these Green's functions can be transformed into virtual responses as if sources and receivers were both located at $\Lambda_x$. These virtual responses can be used for accurate imaging and characterization below this level, since all internal multiple reflections from heterogeneities above this level have been effectively eliminated from the data.

If we are able to generate full Green's functions from the surface to any level in the subsurface, we might go beyond conventional IR. In this abstract we introduce a new alternative redatuming scheme that we refer to as target-enclosing interferometric redatuming (TEIR). In TEIR, we transform Green's functions at two subsurface levels $\Lambda_1$ and $\Lambda_2$ into virtual data as if sources and receivers were located at both $\Lambda_1$ and $\Lambda_2$ in a reference medium. This reference medium is identical to the physical medium in the volume $V$ between $\Lambda_1$ and $\Lambda_2$ and homogeneous outside this volume. In conjunction with the scheme of Wapenaar et al. (2012), TEIR allows us to turn conventional surface seismic data with one-sided illumination into virtual data with two-sided illumination.

Theory

Wapenaar et al. (2008) presented a forward model for IR, starting with a reciprocity theorem of the convolution type for flux-normalized one-way wavefields in the frequency-space domain. In this model, the upgoing (superscript $-$) field $p^-(x_{y_1}, x, \omega)$ at receiver $x_{y_1}$ at level $\Lambda_1$ in the subsurface is expressed in terms of the downgoing (superscript $+$) field $p^+(x, x_{y_1}, \omega)$ at the same level and the upgoing field at a deeper level $\Lambda_2$, according to

$$p^-(x_{y_1}, x, \omega) = \int_{\Lambda_1} R^0_\omega (x_{y_1}, x, \omega) p^+(x, x_{y_1}, \omega) d^2 x + \int_{\Lambda_2} T^0_\omega (x_{y_1}, x, \omega) p^- (x, x_{y_1}, \omega) d^2 x.$$

(1)
In this representation, that is schematically illustrated in Figure 1, \(x_s\) is the source location at the surface and \(\omega\) is the angular frequency. The field \(R_0^+\) is the reflection response at level \(\Lambda_1\) in the reference medium (indicated by the subscript 0 ) and \(T_0^-\) is the transmission response from level \(\Lambda_2\) to \(\Lambda_1\) in the same medium. By choosing level \(\Lambda_2\) infinitely deep in the subsurface and assuming the medium below this level to be homogeneous, it can be shown that the second integral vanishes (Wapenaar et al., 2008). The remaining integral can be inverted for \(R_0^+\), being the essence of conventional IR (Van der Neut, 2012). However, by having chosen the lower level infinitely deep, the retrieved response contains reflections from the entire medium below the upper level.

In TEIR we make a different choice by selecting \(\Lambda_2\) below a target zone in the subsurface (and \(\Lambda_1\) above this zone). As a consequence, we have a single equation with two unknowns, leading to an ill-posed inverse problem. To constrain this problem, we introduce a second forward model. Starting with a reciprocity theorem of the correlation type and repeating the steps in the derivation of Wapenaar et al. (2008), we find

\[
p^+ (x_{y_1}, x_s, \omega) = \int_{\Lambda_1} (R_0^+ (x_{y_1}, x, \omega))^* p^- (x, x_s, \omega) d^2x + \int_{\Lambda_2} (T_0^- (x_{y_1}, x, \omega))^* p^+ (x, x_s, \omega) d^2x, \tag{2}
\]

where superscript * denotes complex conjugation. Equation 1 describes the field that leaves \(V\) at the upper level in terms of all fields that enter \(V\), whereas equation 2 describes the field that enters \(V\) at the upper level in terms of all fields that leave \(V\). Two similar expressions can be derived for the fields that leave and enter \(V\) at the lower level:

\[
p^+ (x_{y_2}, x_s, \omega) = \int_{\Lambda_1} T_0^+ (x_{y_2}, x, \omega) p^+ (x, x_s, \omega) d^2x + \int_{\Lambda_2} R_0^- (x_{y_2}, x, \omega) p^- (x, x_s, \omega) d^2x, \tag{3}
\]

\[
p^- (x_{y_2}, x_s, \omega) = \int_{\Lambda_1} T_0^- (x_{y_2}, x, \omega) p^- (x, x_s, \omega) d^2x + \int_{\Lambda_2} R_0^+ (x_{y_2}, x, \omega) p^+ (x, x_s, \omega) d^2x, \tag{4}
\]

where \(R_0^-\) is the reflection response from level \(\Lambda_2\) to level \(\Lambda_1\) (illuminating the target from below) and \(T_0^+\) is the transmission response from level \(\Lambda_1\) to \(\Lambda_2\). We discretize equations 1, 3 and the complex conjugate of equations 2 and 4, using the notation of Berkhout et al. (2012). In this notation, row \(i\) and column \(j\) contain data from receiver \(i\) and source \(j\). All four equations can be cast into a single matrix inverse problem, which can be inverted for \(R_0^+\), \(T_0^+\), \(R_0^-\) and \(T_0^-\), being discrete representations of \(R_0^+\), \(T_0^+\), \(R_0^-\) and \(T_0^-\), respectively:

\[
\begin{pmatrix}
  p_1^+ \\
  p_2^+
\end{pmatrix} =
\begin{pmatrix}
  R_0^+ & T_0^- \\
  T_0^+ & R_0^-
\end{pmatrix}
\begin{pmatrix}
  p_1^- \\
  p_2^-
\end{pmatrix}, \tag{5}
\]

where \(p_1^+\) and \(p_2^+\) are the up- and downgoing fields at levels \(\Lambda_1\) and \(\Lambda_2\), respectively. By solving this problem through (least-squares) inversion, we enforce that all virtual sources have uniform radiation patterns. To retrieve the fields \(R_0^-\) (illuminating the target from below) and \(T_0^+\), it is required that sufficient multiple scattering is generated below \(\Lambda_2\). Therefore, the inverse problem can only be solved if the medium is sufficiently heterogeneous below this array.
Figure 2  a) Velocity model. Sources are located at the red array at the surface. Receivers are located at the green arrays $\Lambda_1$ and $\Lambda_2$. b) Downgoing field from a source at $\Lambda_1$. c) Retrieved record $R_0^+$ by conventional IR for a virtual source in the center of $\Lambda_1$.

Figure 3  a) Retrieved record $R_0^+$ by TEIR for a virtual source in the center of $\Lambda_1$ with receivers at $\Lambda_1$ versus b) its directly modelled equivalent. c) Retrieved record $T_0^-$ by TEIR for a virtual source in the center of $\Lambda_2$ with receivers at $\Lambda_1$ versus d) its directly modelled equivalent.

Example

Synthetic data have been generated in the velocity model that is shown in Figure 2a. At the surface, 161 sources are located and 81 receivers are positioned at two arrays $\Lambda_1$ and $\Lambda_2$. Strong salt reflections can be found below $\Lambda_1$ and above $\Lambda_2$, providing rich illumination to constrain the inverse problem. Up- and downgoing wavefields are separated at both depth levels. In Figure 2b, we show the downgoing field of an arbitrary shot as observed at array $\Lambda_1$ for illustration. Aim is to redatum the sources to the receiver levels and to eliminate reflections outside the volume $V$ enclosed by these levels. Conventional IR is applied with the data at the upper array only (van der Neut, 2012), yielding a response $R_0^+$ as if there were a virtual source at level $\Lambda_1$, see Figure 2c. Besides the target reflections from $V$ at ~0.5s and ~0.7s, we also observe a strong reflection at ~1.4s, coming from the part of the salt body below $\Lambda_2$. This event is retrieved because the second integral in equation 1 has not been evaluated by IR, leading to the reflection response of a reference medium that is homogeneous above $\Lambda_1$ but identical to the physical medium below $\Lambda_2$.

Next, we apply the new algorithm TEIR. Input data are the decomposed data at both depth levels. Output data are $R_0^+$, $T_0^-$, $R_0$ and $T_0^+$, obtained by least-squares inversion of equation 5. In Figure 3a, we show a shot from $R_0^+$, which can be interpreted as if there were a virtual source and receivers at $\Lambda_1$. Note the similarity with Figure 2c, except for some artefacts that are subject of current investigation. The reflector at ~1.4s has been eliminated by TEIR, as we have defined the reference medium to be homogeneous below $\Lambda_2$ in the forward models. TEIR also provides $T_0^-$, having a virtual source at $\Lambda_2$ and receivers at $\Lambda_1$. This response and its directly modelled equivalent are shown in Figure 3c-d.
Besides these responses, we have also retrieved $R^-$, which can be interpreted as if there were virtual sources and receivers at $\Lambda_2$, illuminating the target zone from below. A retrieved record from this response and its directly modelled equivalent are shown in Figure 4a-b. Note that these responses are constructed solely from multiply scattered waves, providing the required illumination from below. Finally we find $T^+$, having a virtual source at $\Lambda_1$ and receivers at $\Lambda_2$. This response and its directly modelled equivalent are shown in Figure 4c-d. For all responses, we report a proper match, except that the amplitude-versus-offset characteristics are not properly retrieved (yet), which is subject of current investigation.

Discussion and conclusion

Ideally, we want to retrieve Green's functions to levels $\Lambda_1$ and $\Lambda_2$ from surface seismic data with the scheme of Wapenaar et al. (2012), followed by TEIR. Such an approach would allow us to turn conventional seismic data with one-sided illumination (from above only) into virtual data with two-sided illumination (both from above and below). The additional illumination from below, being constructed solely from multiply scattered waves, could complement the primary seismic illumination in shadow zones and provide additional input to constrain the inversion for rock and fluid properties. Note also that the data volume is reduced significantly by TEIR, since we have datumed the acquisition levels close to the target zone. We emphasize that the results we have shown in this abstract are preliminary and that more could probably be achieved with additional (reciprocity and sparsity) constraints and filters to make the inversion more robust, which is currently being investigated. The method should also still be tested in different environments with less favorable scattering and additional complications such as elastodynamic wave conversions, diffractions, intrinsic losses and noise.

References